Emerging opportunities for enhanced seafloor observation: seismology, geodesy, and EM

James Gaherty Lamont-Doherty Earth Observatory of Columbia University

Rob Evans, Kerry Key, Patty Lin, Scott Nooner, Doug Toomey, Spahr Webb, OBSIP

Seafloor Observations: Outline

- 1. Why do we need observations from the oceans?
- 2. Current state of observational capabilities
 - a) Ocean-bottom seismographs (OBS)
 - b) Seafloor Geodesy
 - c) Marine Electromagnetics
- 3. Frontier Needs and Opportunities
 - a) Data delivery
 - b) Instrument improvement
 - c) Long-term observations
- 4. Obstacles

Why do we need observations from the oceans?



Explicitly called out in

- 7/10 Seismological Grand Challenges
- 3/7 Geodetic Grand Challenges
- 2/8 Ocean Science Board decadal priorities
- (implicit in more)
- "Offshore" comprises 71% of surface, most plate boundaries, all subduction zones, great majority of earthquakes and volcanism
- The shoreline does not define our science. It should not limit our facilities.

Courtesy Chen Ji



Home



Contacts

Data

OCEAN BOTTOM SEISMOGRAPH INSTRUMENT POOL

Experiment Planning

Current OBS Capabilities

About OBSIP

160 broadband systems (standard pool plus Cascadia ARRA instruments) ~1 year duration at sample rates up to ~100 Hz deployment depths: 1000-6000 meters (free-fall, self buoyancy) – 140 systems 0-1000 m (wireline, pop-up buoy or ROV) – 20 systems broadband sensor plus DPG/APG

Instruments

93 short-period systems (all in standard pool) active-source deployments of a few months, sample rates up to 250 hz also used as short-period passive deployments for up to ~8 months deployment depths 0-5x00 m

Maintained and operated by dedicated technical staff (~10 FTE total)

Experiments

Funded out of OCE-MGG (core) by cooperative agreement with IRIS as lead and LDEO, SIO, and WHOI as subcontractors







Lamont-Doherty Earth Observatory Columbia University | Earth Institute



OBSIP success

geoscience

LETTERS PUBLISHED ON LINE: 22 APRIL 2012 | DOI: 10.1038/NGE01454

Variations in earthquake rupture properties along the Gofar transform fault, East Pacific Rise

Jeffrey J. McGuire^{1*}, John A. Collins¹, Pierre Gouédard², Emily Roland^{3†}, Dan Lizarralde¹, Margaret S. Boettcher⁴, Mark D. Behn¹ and Robert D. van der Hilst²

REPORTS

Mantle Shear-Wave Velocity Structure Beneath the Hawaiian Hot Spot

Cecily J. Wolfe,¹* Sean C. Solomon,² Gabi Laske,³ John A. Collins,⁴ Robert S. Detrick,⁴ John A. Orcutt,³ David Bercovici,⁵ Erik H. Hauri²



LETTER

106.4

doi:10.1038/nature14113

105.6

Seismic evidence of effects of water on melt transport in the Lau back-arc mantle

S. Shawn Wei¹, Douglas A. Wiens¹, Yang Zha², Terry Plank², Spahr C. Webb², Donna K. Blackman³, Robert A. Dunn⁴ & James A. Conder⁵

106.2



105.8

Longitude (° W)

OBSIP success



Courtesy of Doug Toomey. Data compiled by Jessica Lodewyk, OBSIP OMO

Recent and ongoing Enhancements





- Trawler resistance
 - Deploy < 1000 m depth</p>
- Shielding
 - Reduce current-induced tilt on horizontals
- Non-glass flotation
 - Deploy > 5000 m
- Chip-scale atomic clocks

 Negligible drift

Seafloor Geodetic Techniques

Vertical Deformation

- Bottom Pressure Recorders (BPRs or APGs)
 - Continuous, easy to deploy, but drift
- Mobile Pressure Recorders (MPRs)
 - More precise, but campaign w/ROV
- Tilt sensors
- Self Calibrating Pressure Recorders (SCPRs)
- Repeat high-res AUV Bathymetry
- Borehole Pressure and Tilt

Horizontal Deformation

- Horizontal Acoustic Ranging
- GPS-Acoustic
- Fiber Optic Seafloor Strainmeter (FOSS)







Acoustic Ranging Systems for Seafloor Geodesy



Burgmann and Chadwell, 2014





Spiess and Chadwell

Chadwell, et al, 1999, Chadwick et al, 1999, McGuire and Collins, 2013

Spanning the deformation spectrum



Burgmann and Chadwell, 2014

Marine Electromagnetics

Controlled-Source Electromagnetic (CSEM) Method



- Deep-towed EM transmitter injects energy into seabed
- EM energy diffuses through the sea and seabed
- Energy decay measured by array of EM receivers





Magnetotelluric (MT) Method



$\mathbf{E}(\omega) = \mathbf{Z}(\omega)\mathbf{H}(\omega)$

Natural-source, low-frequency method for crust and mantle imaging

• Measures induced electric and magnetic fields to estimate impedance (Z)

Marine Electromagnetics



Exceptional sensitivity to fluids – hydration, partial melting
Anisotropic – spatial distribution of heterogeneity, rock fabric
Resolution analogous to active/passive seismic

•Highly complementary with seismic constraints for resolving tradeoffs



MT

Frontier Challenges and Opportunities

- Remote data delivery
 - Episodic (AUV)
 - Quasi-real-time (cable, buoy, waveglider)
- Improve instrument performance to onshore levels
 - Reduce horizontal noise on seismometers
 - Improve vertical and horizontal geodetic precision
- Multi-sensor mini-observatories
 - Seismic and geodetic-quality pressure
 - Seismic and EM
 - Capitalize on logistics investment
 - Capitalize on data-delivery investment
- Long-term observatories
 - Capitalize on logistics
- International Collaboration
 - Pacific Array, SZO





Obstacles

Conclusions of the National Science Board Decadal Survey of Ocean Sciences:

OBS fleet "low relevance" for OCE science

The R/V Marcus G. Langseth

Recommended that NSF consider taking out of service

- •Working in the oceans is expensive
- •Facilities operated by OCE
- •Solid Earth Science is a minor fraction of OCE Science Priorities

Obstacles

Conclusions of the National Science Board Decadal Survey of Ocean Sciences:



The Cascadia Initiative: A Sea Change in Seismology



Douglas R. Toomey, Richard M. Allen, John A. Collins, Robert P. Dziak, Emilie E. Hooft, Dean Livelybrooks, Jeffrey J. McGuire, Susan Y. Schwartz, Maya Tolstoy, Anne M. Trehu, William S. Wilcock



Supported by the National Science Foundation

BBOBS Data Quality



Mobile Pressure Recorders

MPRs

- "Campaign-style" survey using an ROV – analogous to leveling surveys on land.
- Measure water pressure on top of permanent concrete benchmarks.
- Repeat every 1-3 years.



Nooner, Chadwick, Zumberge

Mobile Pressure Recorders

MPRs

- Repeatedly visit
 benchmarks during a
 2-3 day dive (typically).
- Scatter of repeats gives data uncertainty
 -- typically <1 cm.
- Don't capture episodic events.

Nooner, Chadwick, Zumberge



Recent Direct Path Acoustic Ranging





~1 mm/yr interseismic strain signals are detectable over 1-5 km baselines.

Discovery Transform Fault, EPR