

# Seismological constraints on glacial processes

Meredith Nettles

Lamont-Doherty Earth Observatory  
COLUMBIA UNIVERSITY | EARTH INSTITUTE



**KVUG**



**NERC**  
SCIENCE OF THE  
ENVIRONMENT





An aerial photograph of a vast, flat ice sheet under a clear sky. The ice is a mix of white and light blue, with numerous small, dark patches of rock or sediment. A large, dark, rectangular ice island is visible in the center-right, surrounded by a network of channels and ridges. The horizon is flat and distant.

## Seismic constraints for understanding ice-sheet and glacier dynamics:

- Earth structure: mantle/crust; ice-bed interface; within ice
- ice deformation: earthquakes



# Ice loss and sea-level rise

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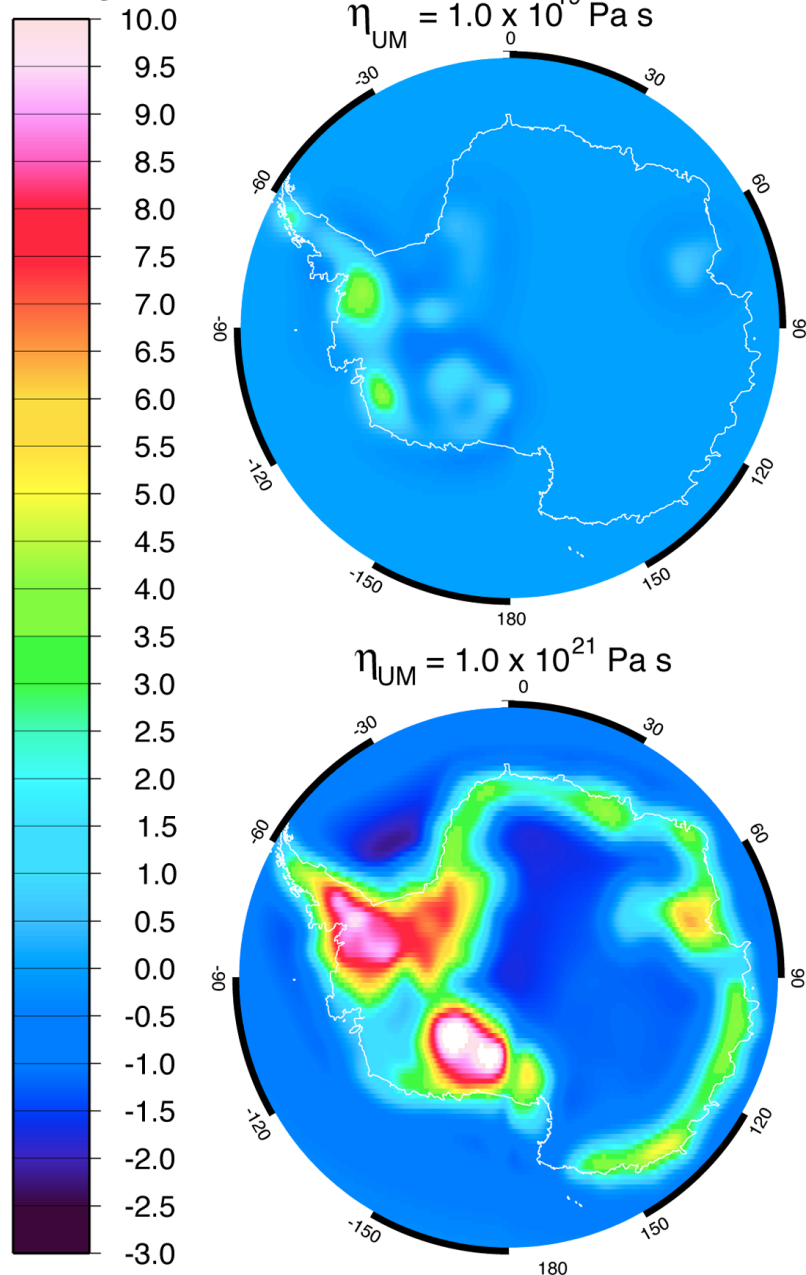
- Probable sea-level rise by 2100: 28-43 cm \* (IPCC, 2007)
- Probable sea-level rise by 2100: 26-98 cm \* (IPCC, 2013)

\* Does not include effects of ice-sheet dynamical changes.

\* “Confidence in projections of global mean sea level rise has increased ... inclusion of ice-sheet dynamical changes.”

# Glacial Isostatic Adjustment Depends on Earth Structure

mm/yr



Low viscosity:  
short-term memory  
Sensitive only to Holocene

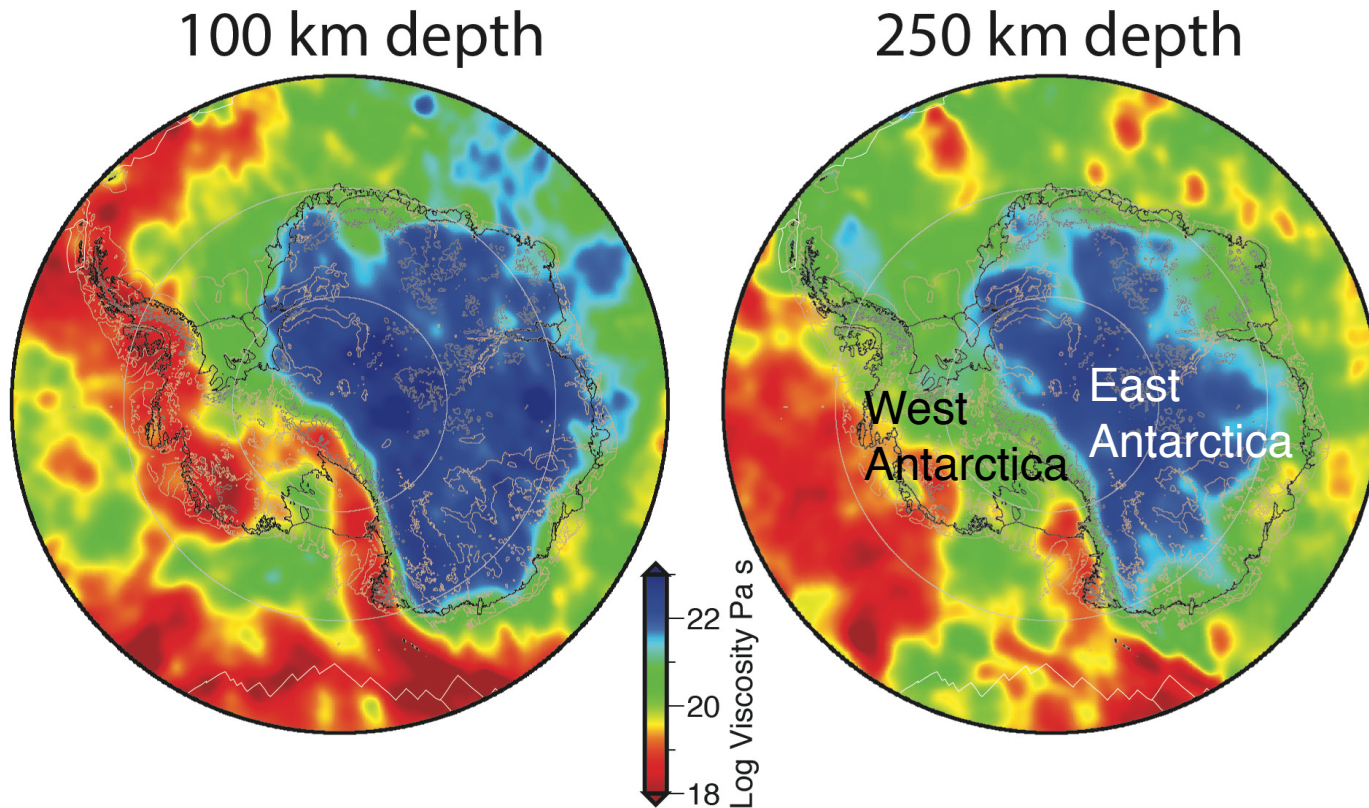
Computed GIA Uplift Rate for  
Different Mantle Viscosities  
(Ivins & James, 2005)

High Viscosity:  
long-term memory  
Sensitive to LGM

*(courtesy D. Wiens)*



# Estimated Mantle Viscosity Maps



isotropic shear velocity (Lloyd et al., in prep) --> viscosity estimate  
method of Wu et al., 2012, and relationships from experimental rock mechanics

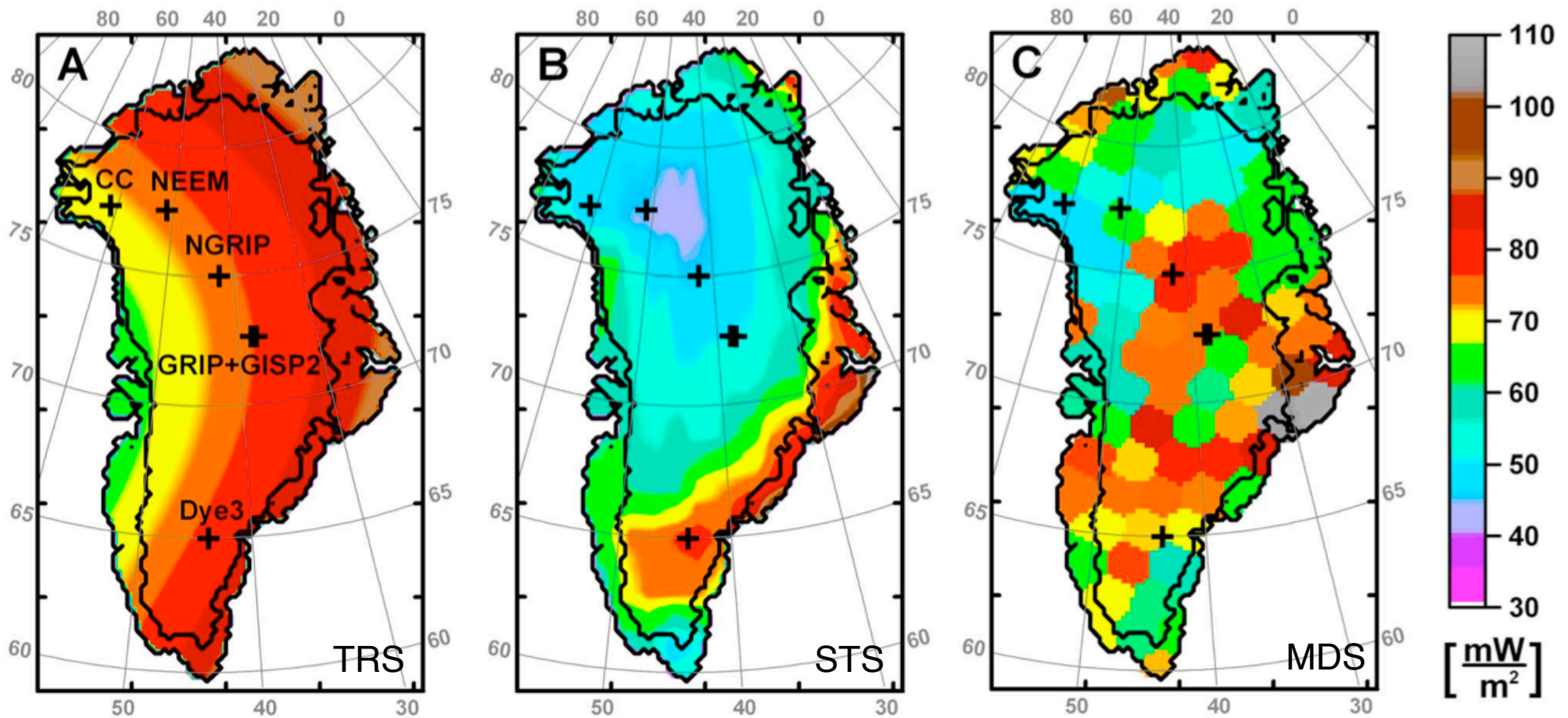
Very low viscosity ( $\sim 10^{18} - 10^{19}$  Pa s) in upper mantle beneath much of West Antarctica --  
GIA should occur over time periods of less than 500 years!

(courtesy D. Wiens)



# Ice-sheet evolution depends on basal heat flux

## Heat-flux models for Greenland (Rogozhina et al., 2012)



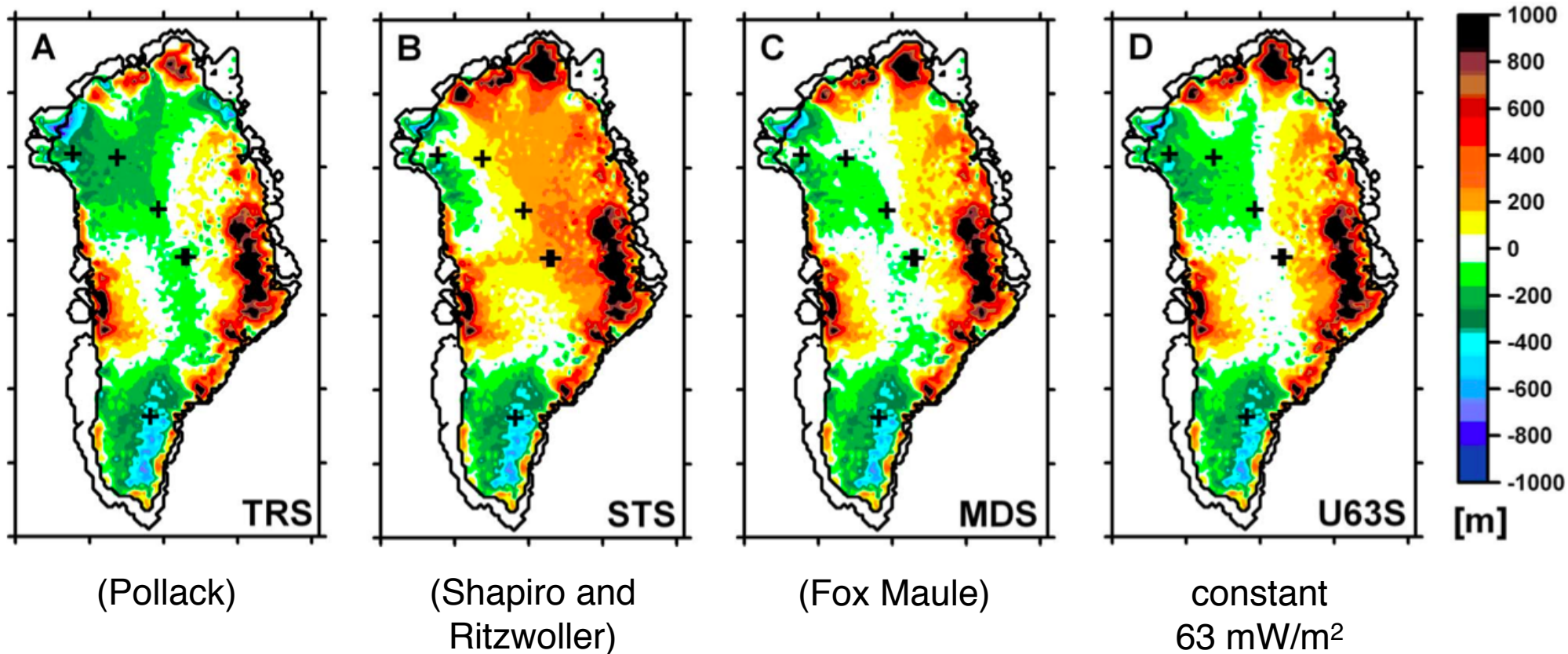
Pollack et al., 1993

Shapiro and  
Ritzwoller, 2004

Fox Maule et al.,  
2009



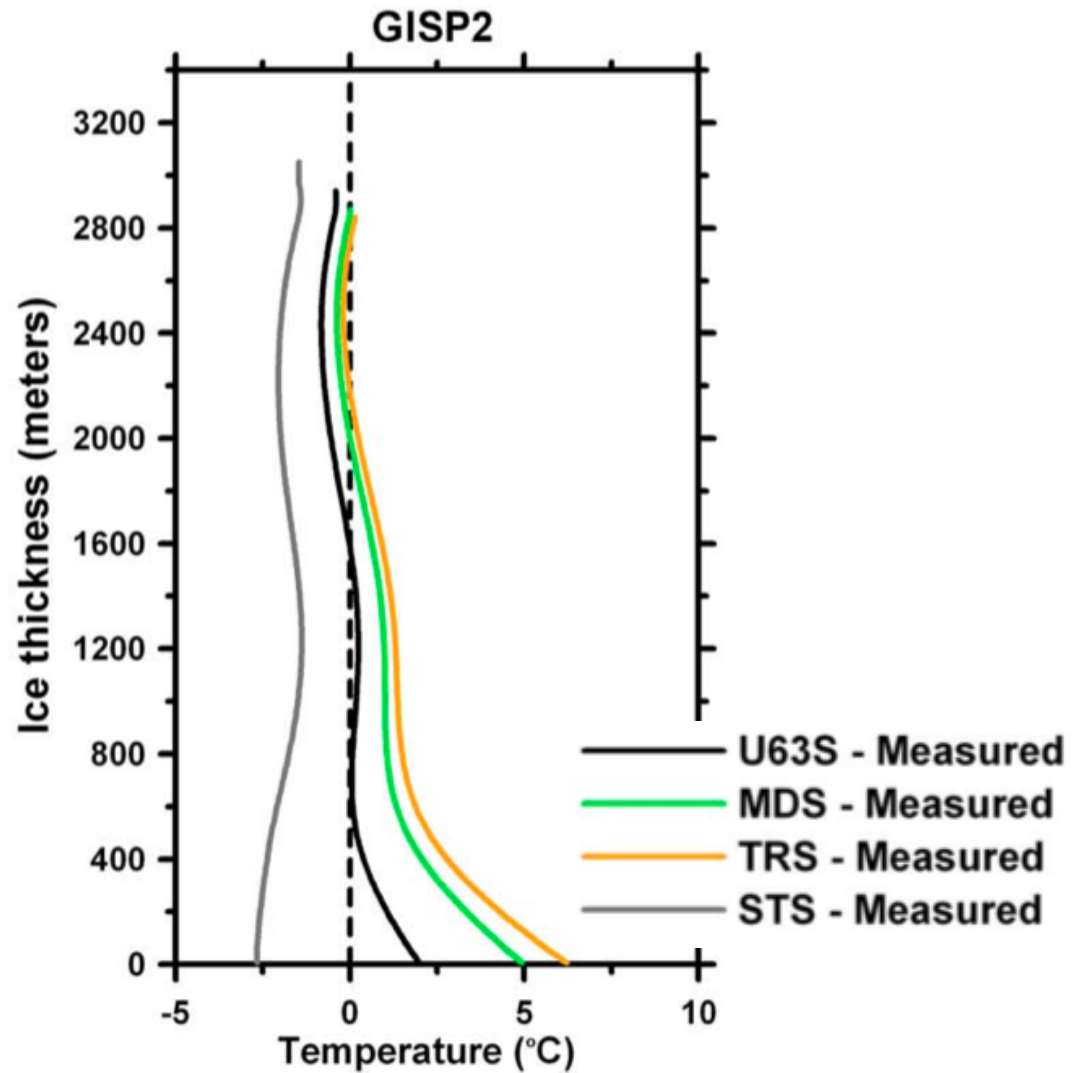
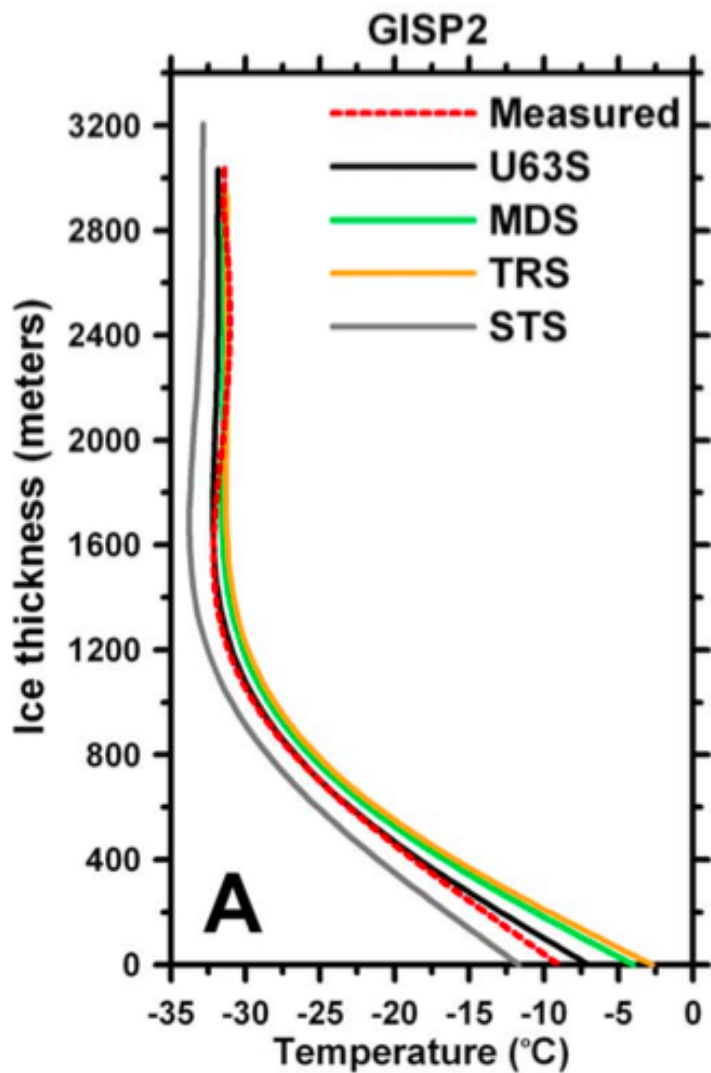
# Predicted modern-day ice thickness (predicted minus observed)



(Rogozhina et al., 2012)



# Predicted temperature profile in ice, and residual

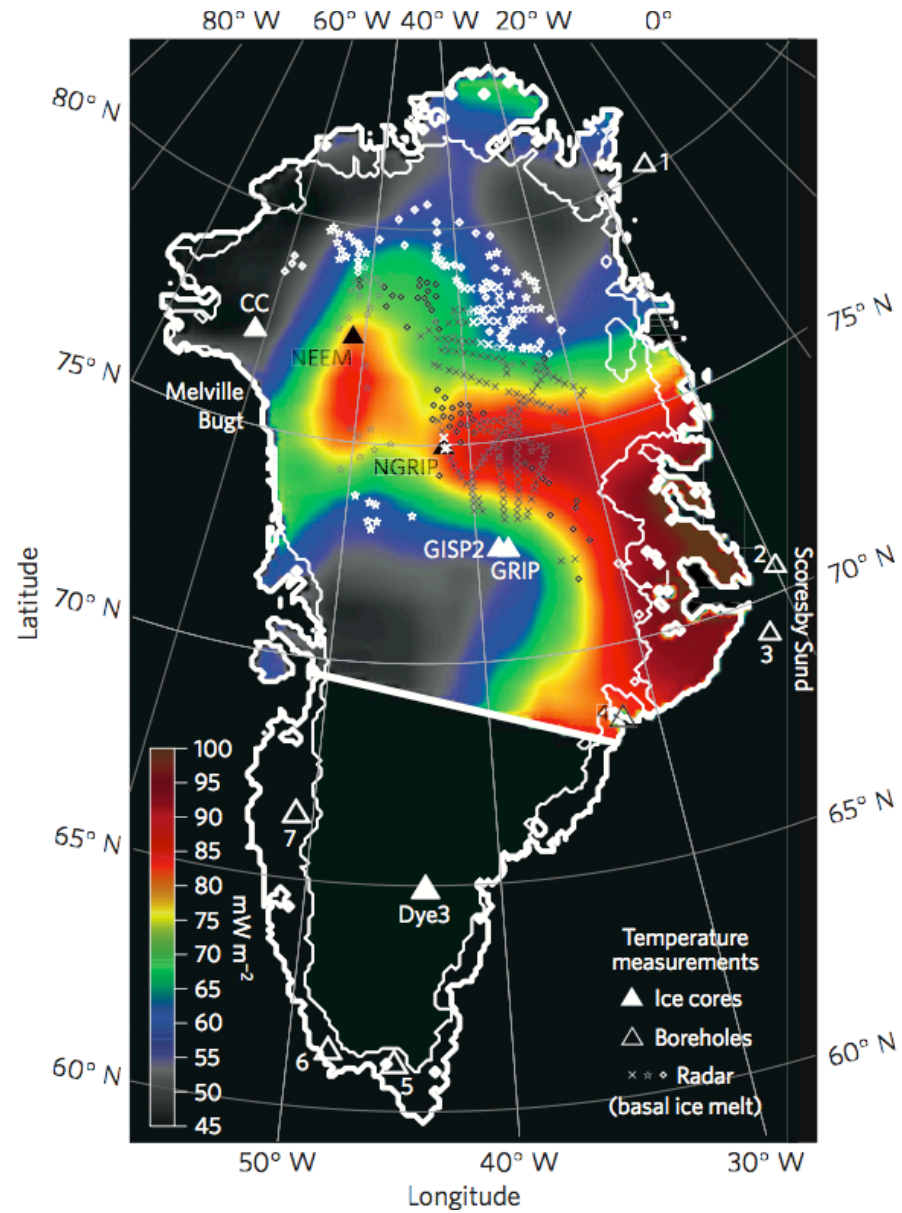


(Rogozhina et al., 2012)



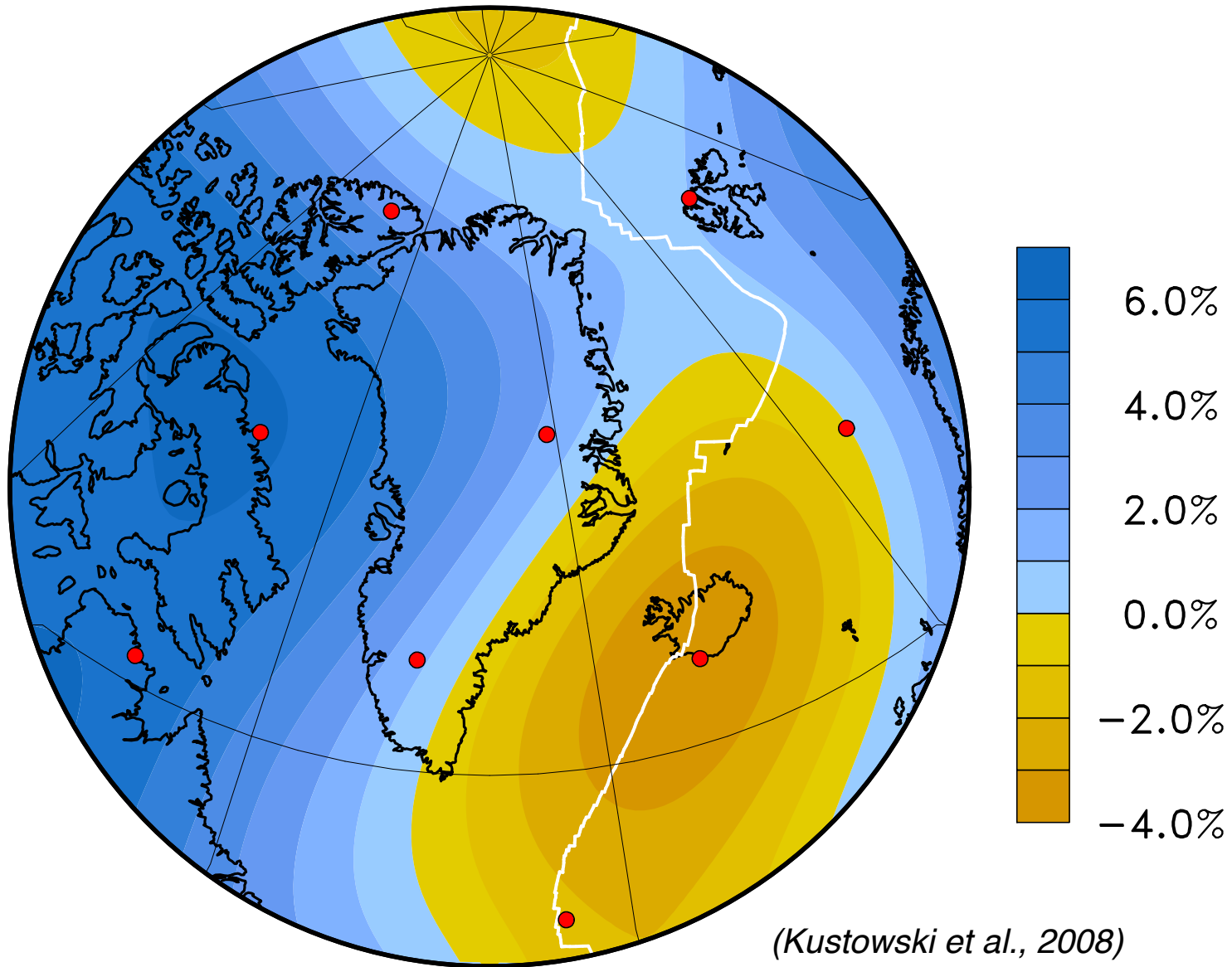
# Heat-flux models for Greenland

(Rogozhina et al., 2016)

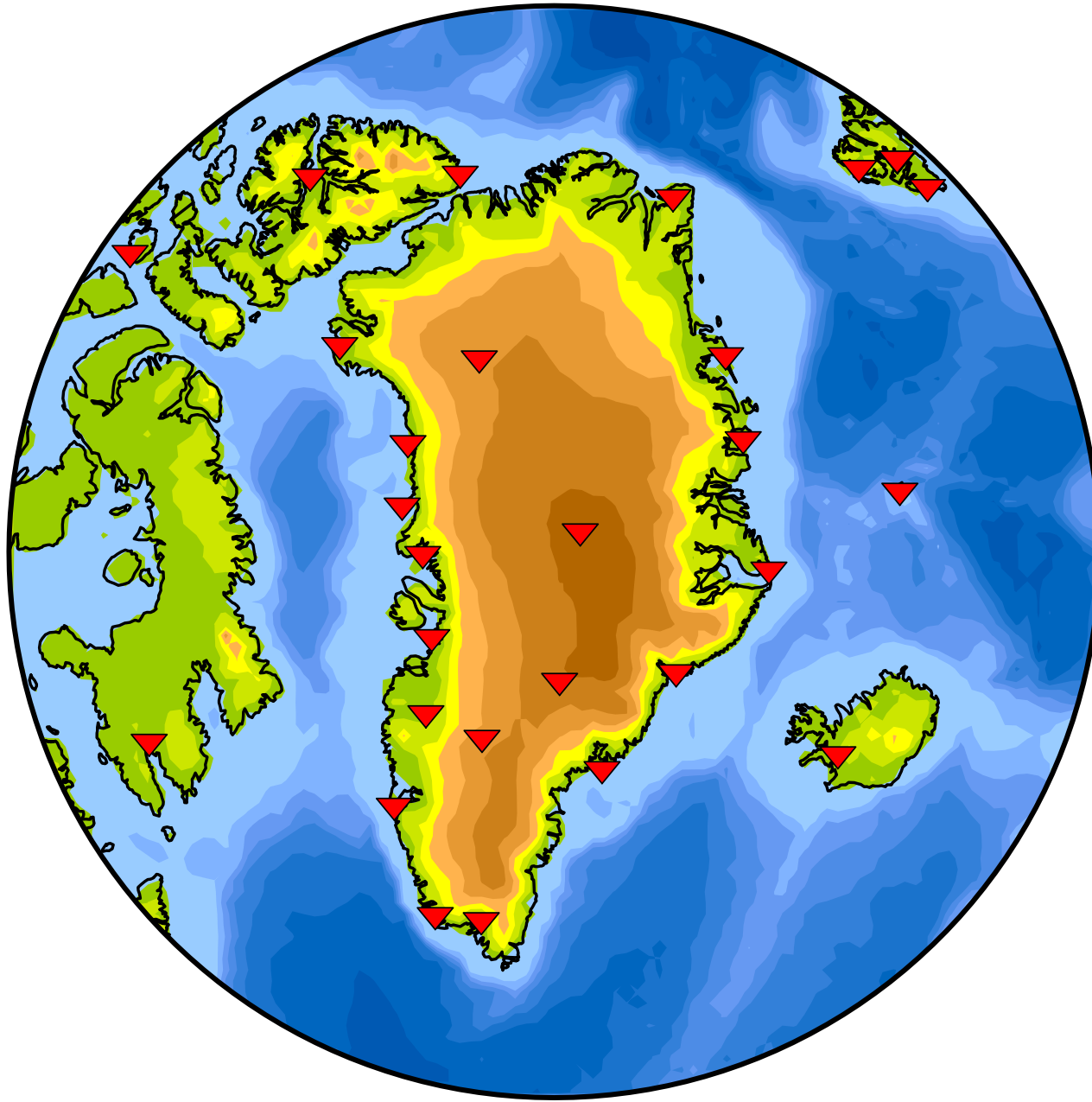




# Global seismic-tomography models: resolving wavelength $\approx$ size of Greenland



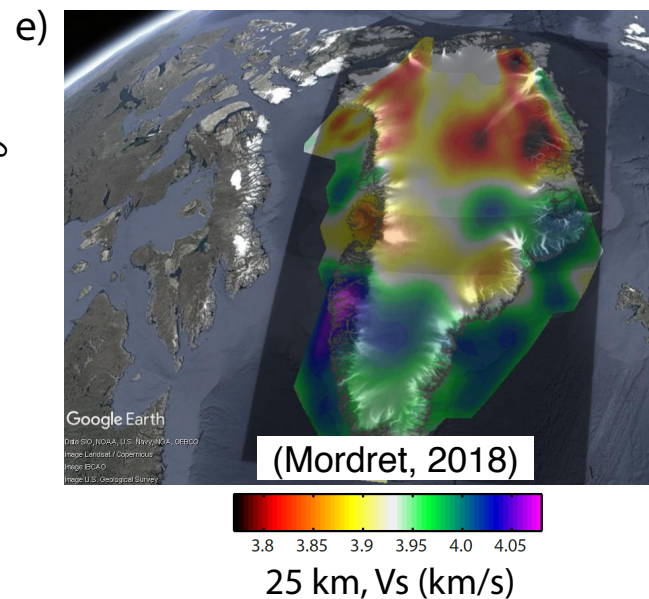
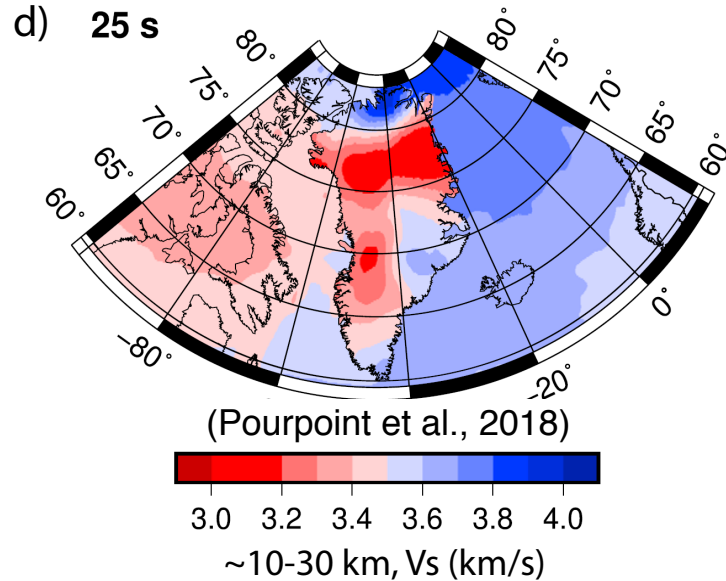
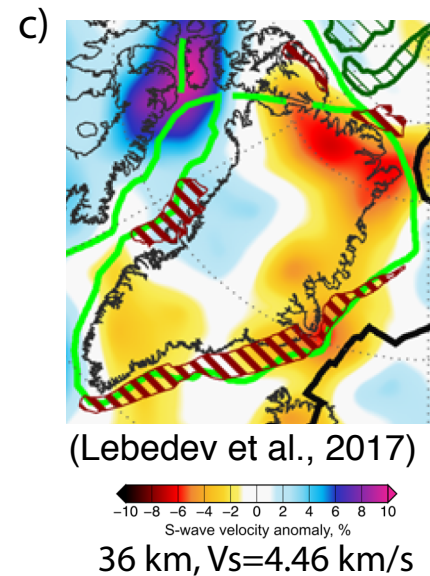
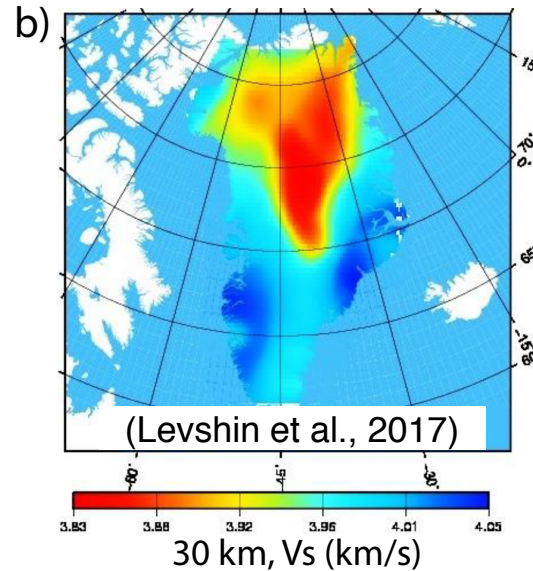
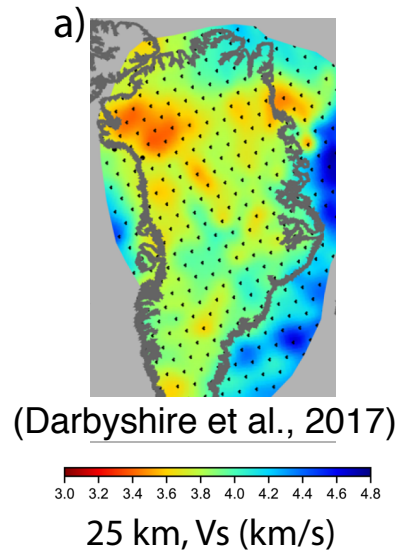
# GLISN stations



International, cooperative effort

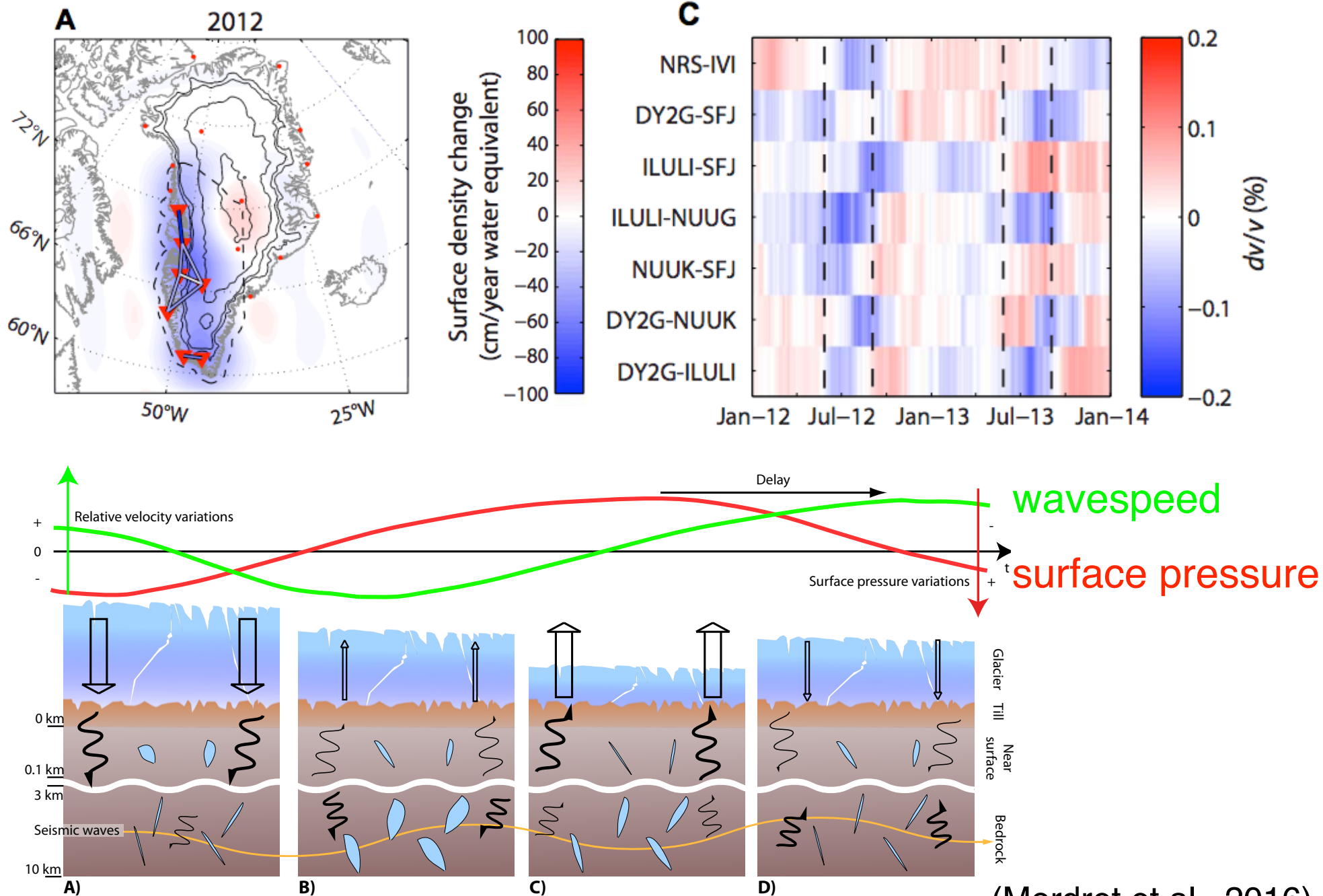


# Recent S-velocity models for Greenland at crustal depths



(compiled by A. Mordret)

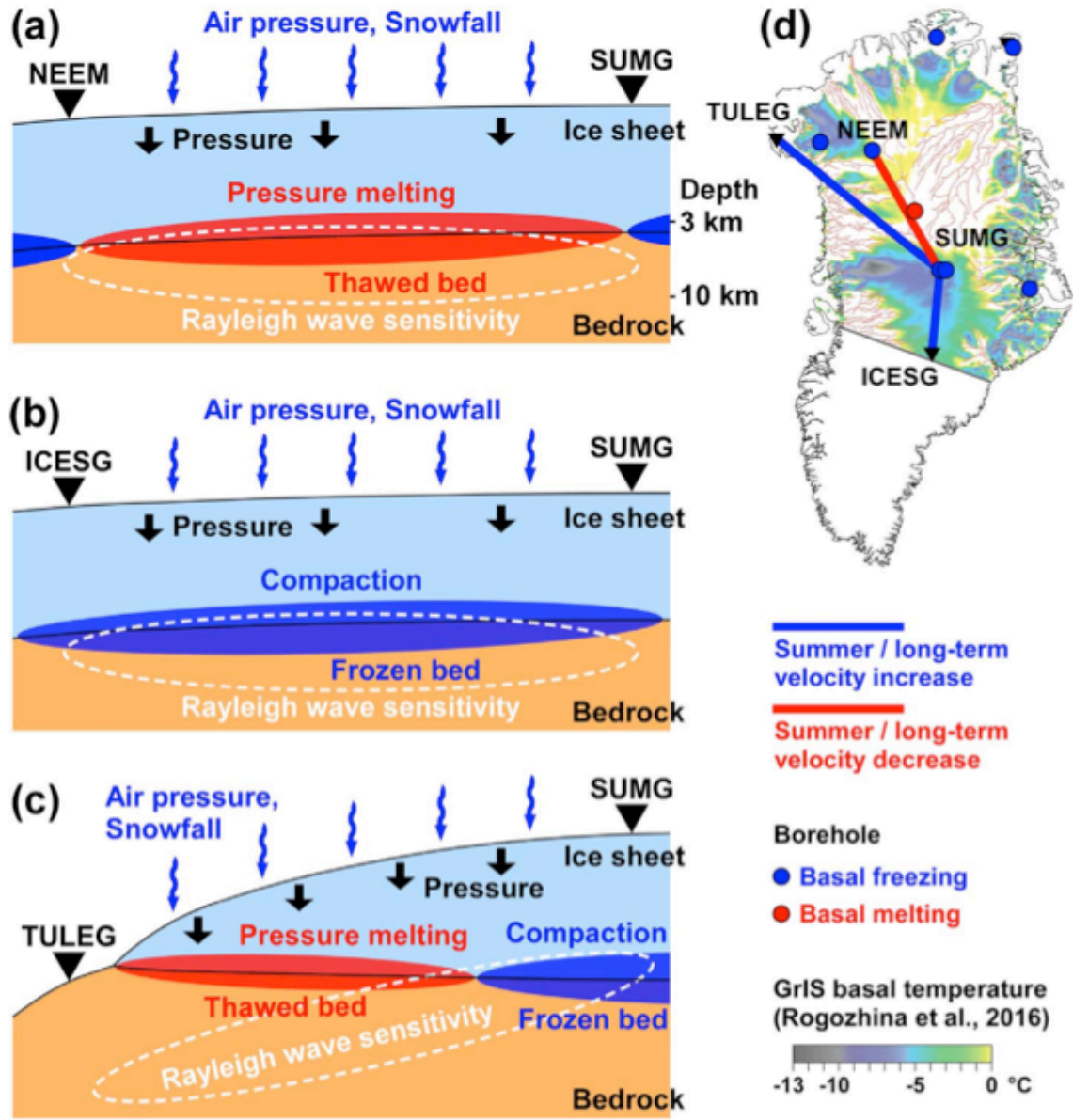
# Ice-bed interface: time-varying seismic wave speeds



(Mordret et al., 2016)

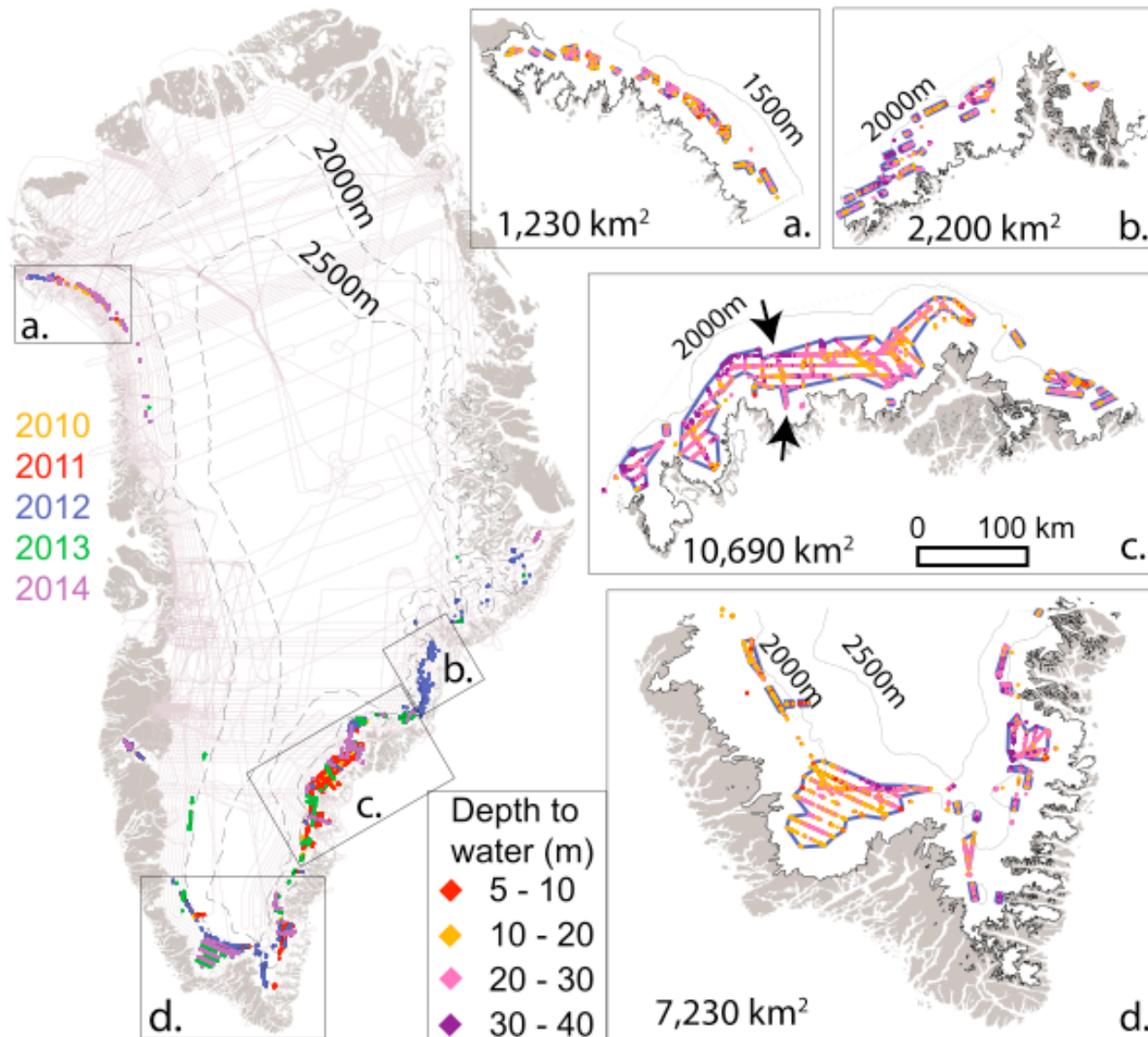


# Ice-bed interface: time-varying seismic wave speeds



(Toyokuni et al., 2018)

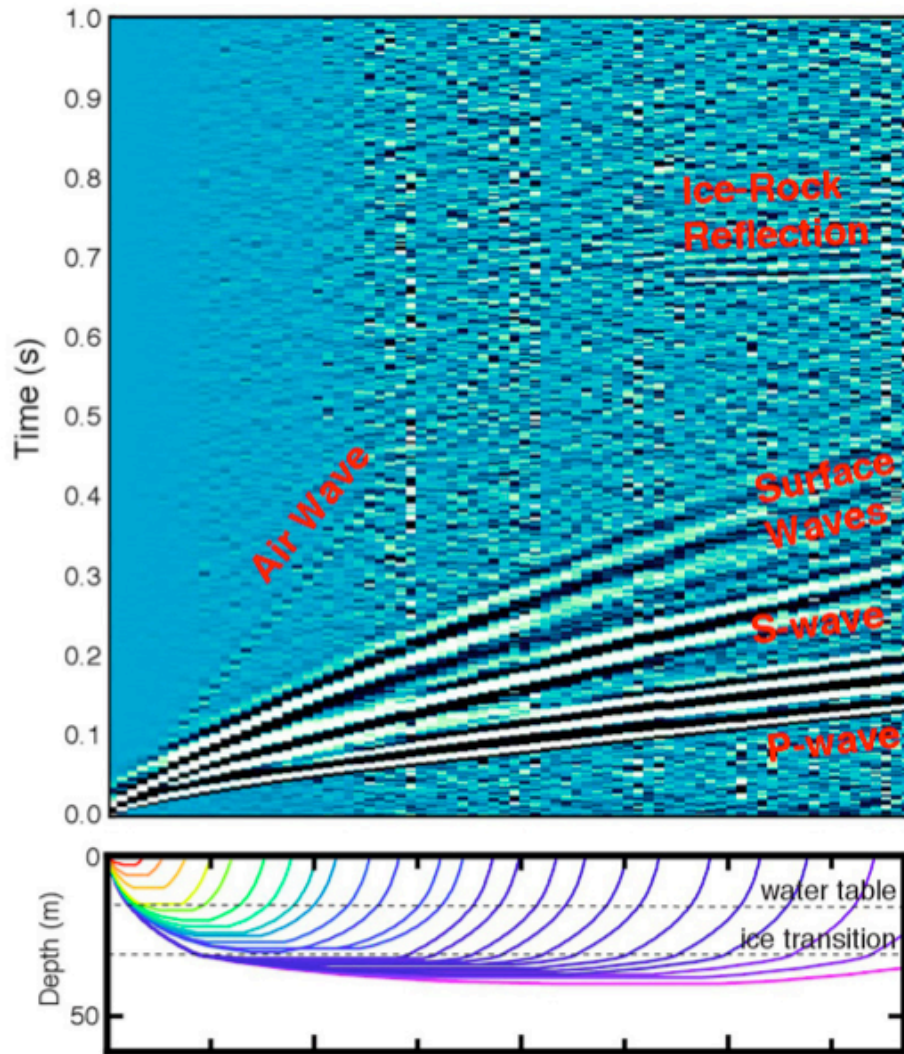
# Firn aquifers in Greenland: storage and discharge of liquid water!



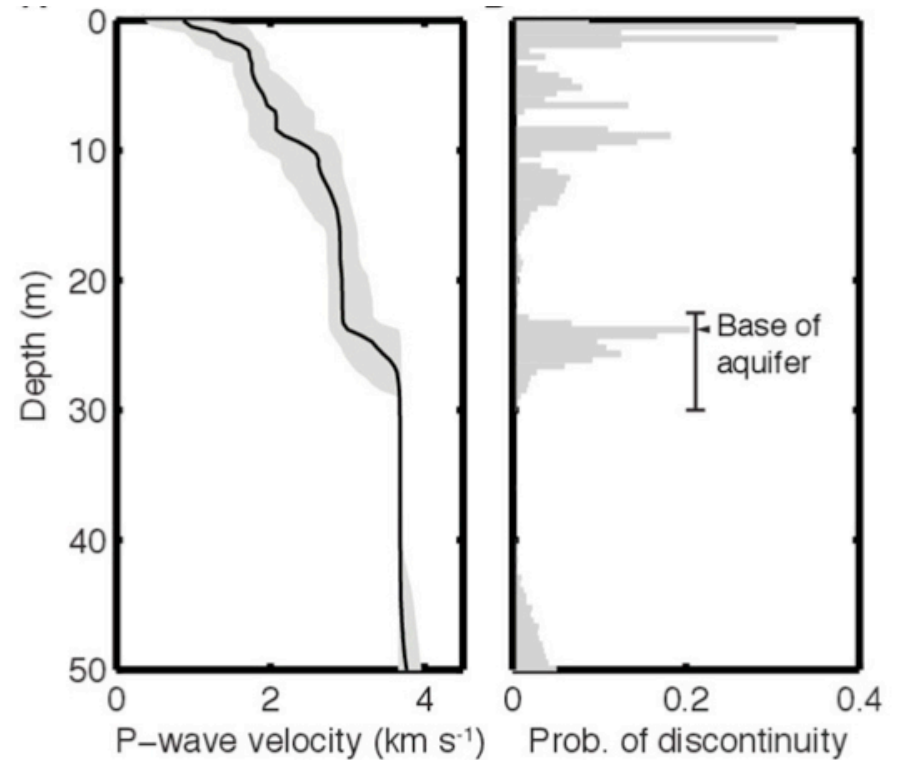
(Miège et al., 2016)



# Surface mapped by radar; base by seismic refraction



(Montgomery et al., 2017)



Probable (but unknown!)  
effect on ice dynamics

# Ice loss and sea-level rise

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- Probable sea-level rise by 2100: 28-43 cm (IPCC, 2007)
- Probable sea-level rise by 2100: 26-98 cm (IPCC, 2013)

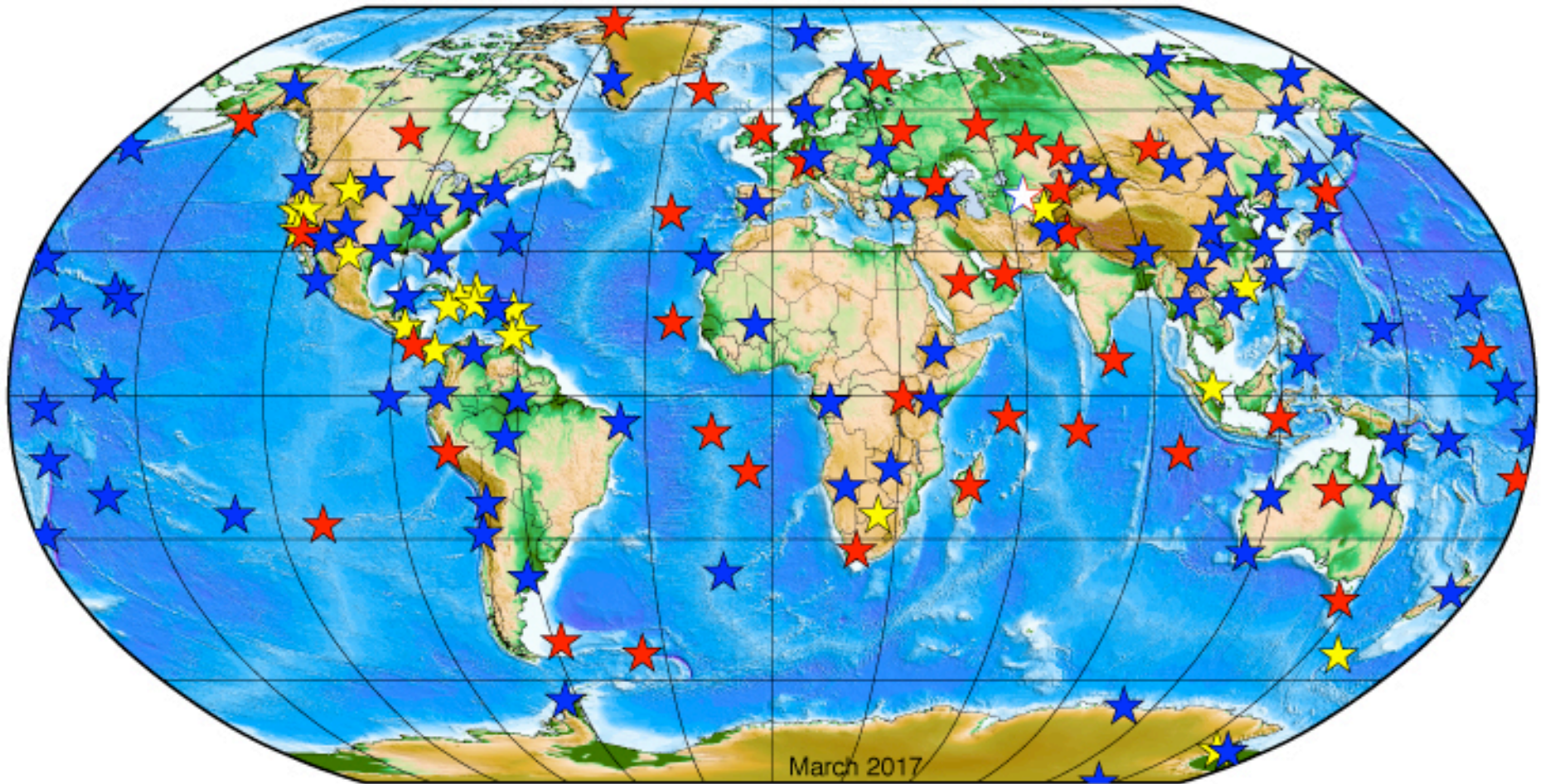
About half of net mass loss in Greenland is “dynamic” (calving).

...

And ~20% of the dynamic loss is observed on GSN stations...



# Global Seismographic Network

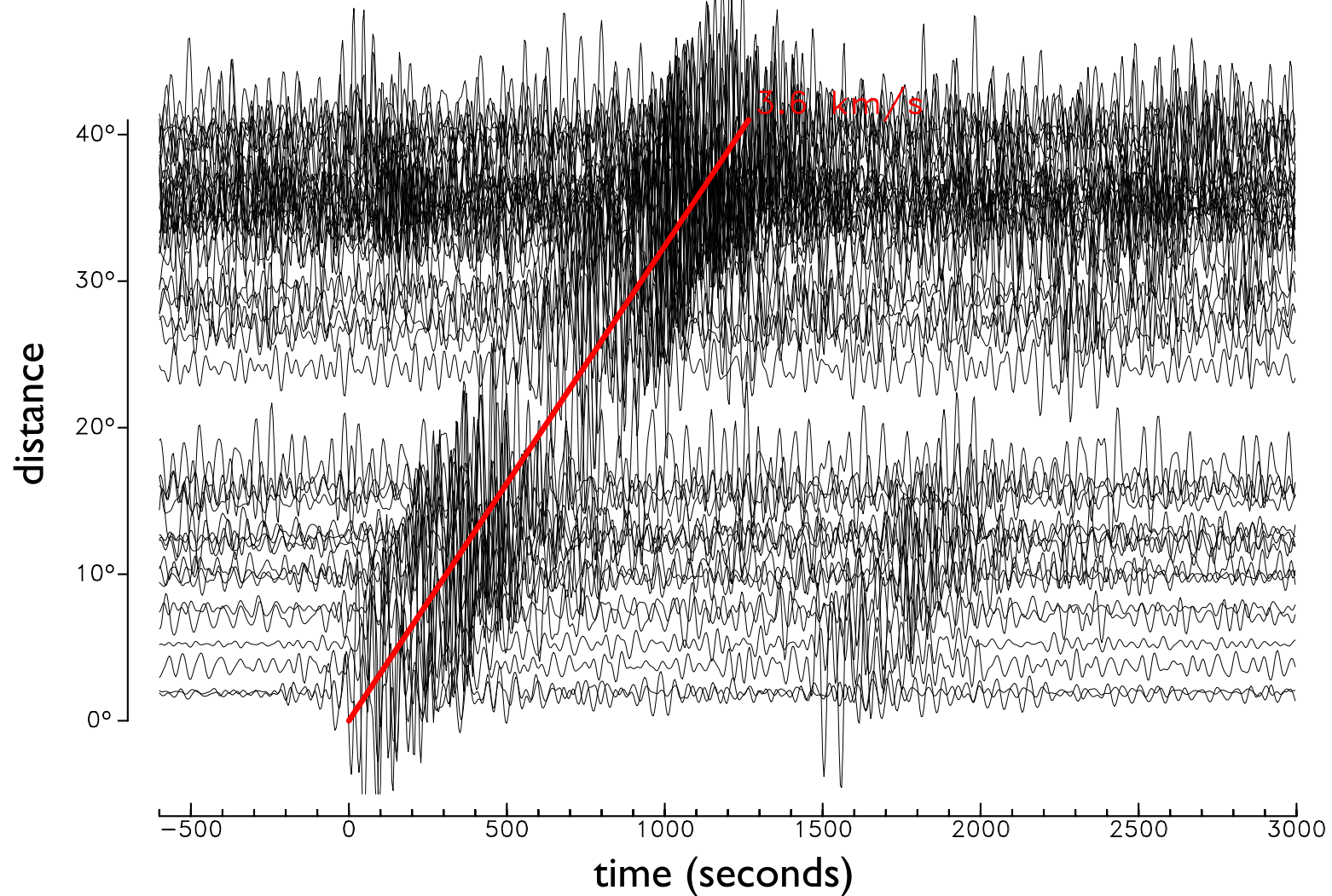


- ★ IRIS/IDA Stations
- ★ IRIS/USGS Stations
- ★ Affiliate Stations
- ★ Planned Stations

# Greenland earthquake: Seismograms from Global Seismographic Network

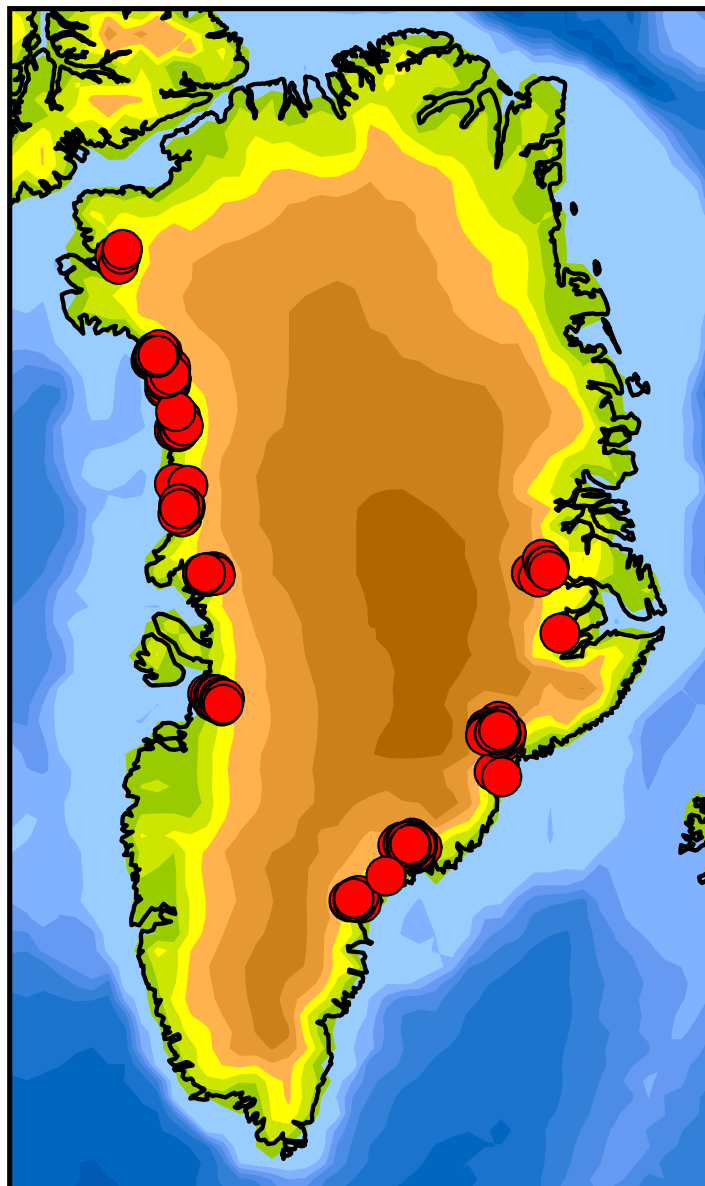
E201404052308A

Event: 2014/04/05, 23:08:24.0, WESTERN GREENLAND  
Hypocenter (SWEQ): Lat= 76.25, Lon= -61.25, h= 10.0, mb=0.0, MS=4.9  
Filter: VEL 75.0 60.0 35.0 25.0, Component: 1





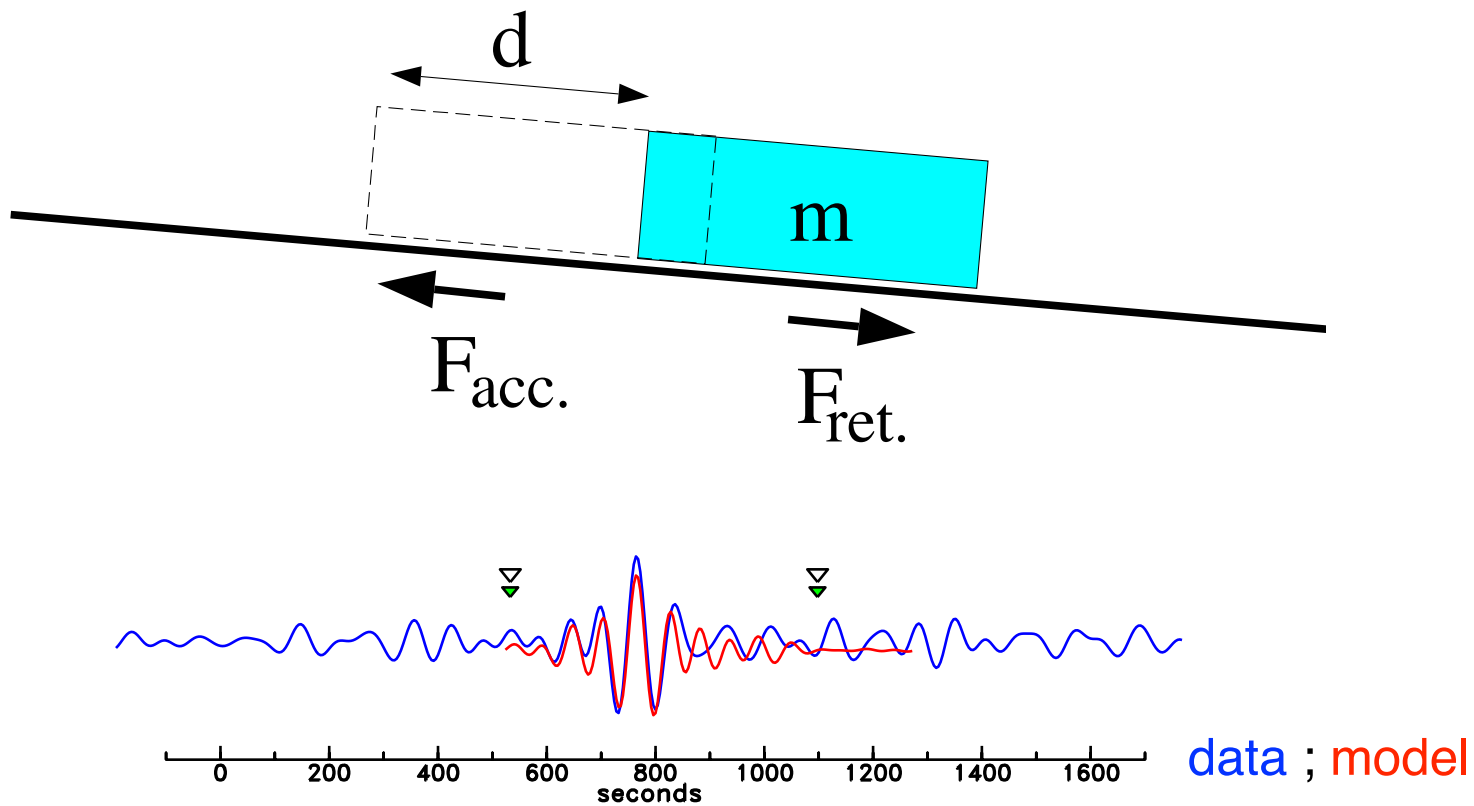
# Glacial earthquakes in Greenland



*(updated from Nettles and Ekström, 2010)*

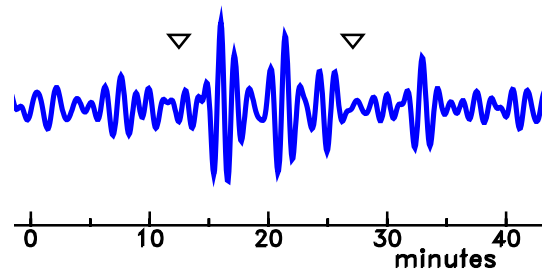
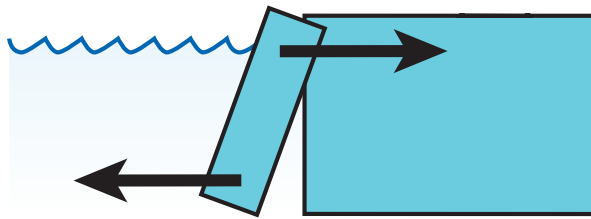


# Momentum-transfer (“landslide”) force model



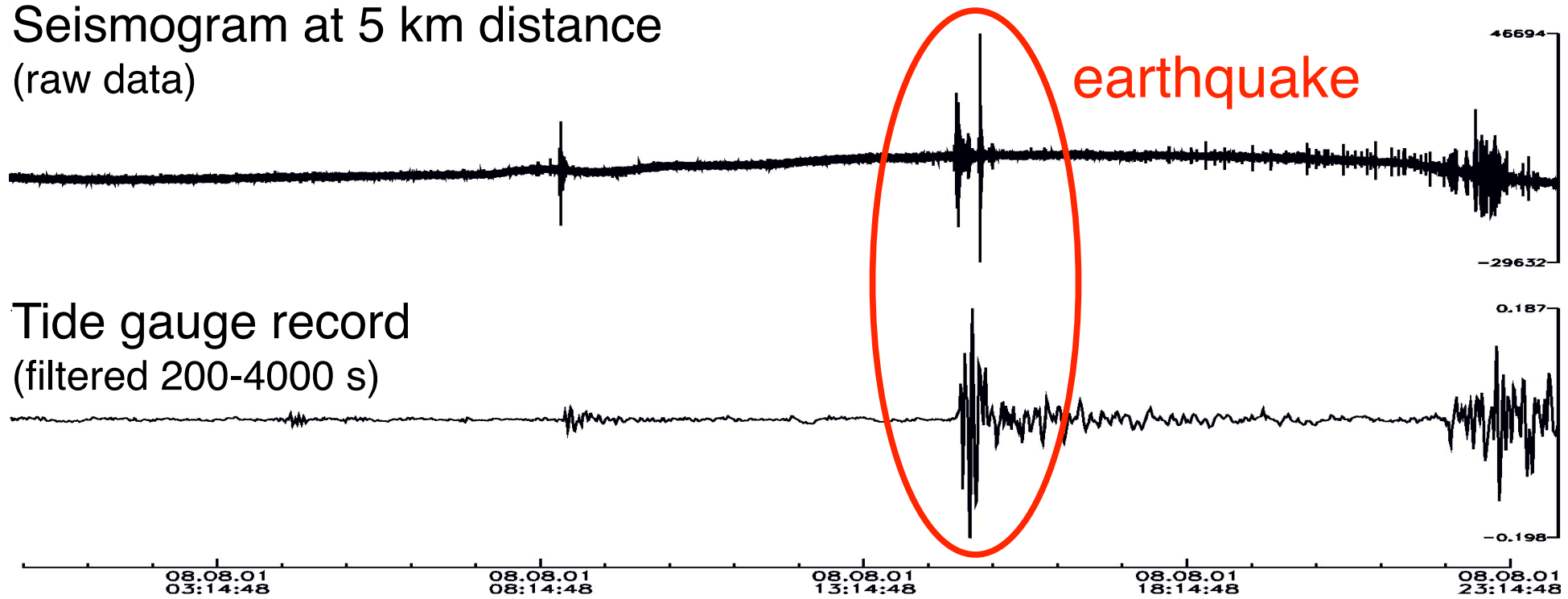
(Ekström et al., 2003; Tsai and Ekström, 2007; Nettles and Ekström, 2010; Veitch and Nettles, 2012; Walter et al., 2012; Sergeant et al., 2016; Olsen and Nettles, 2017)

Seismic signal is generated by reaction force on glacier face as iceberg capsizes

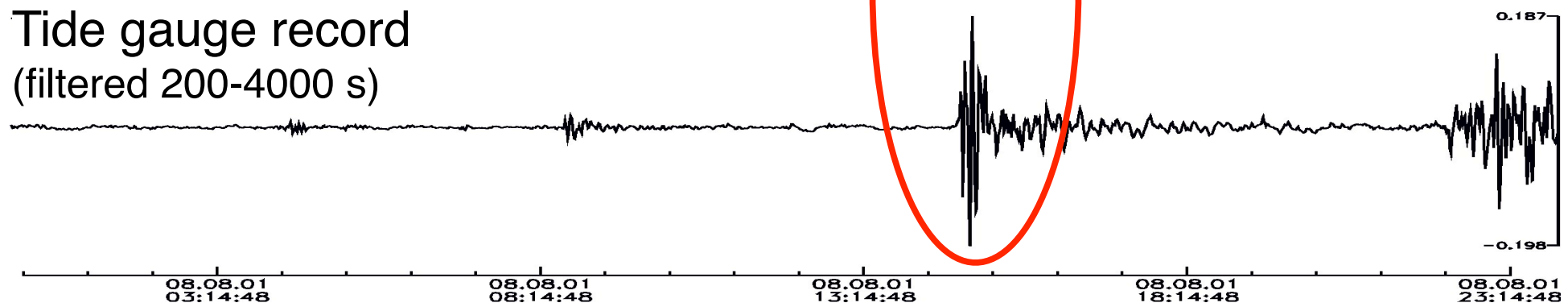




Seismogram at 5 km distance  
(raw data)

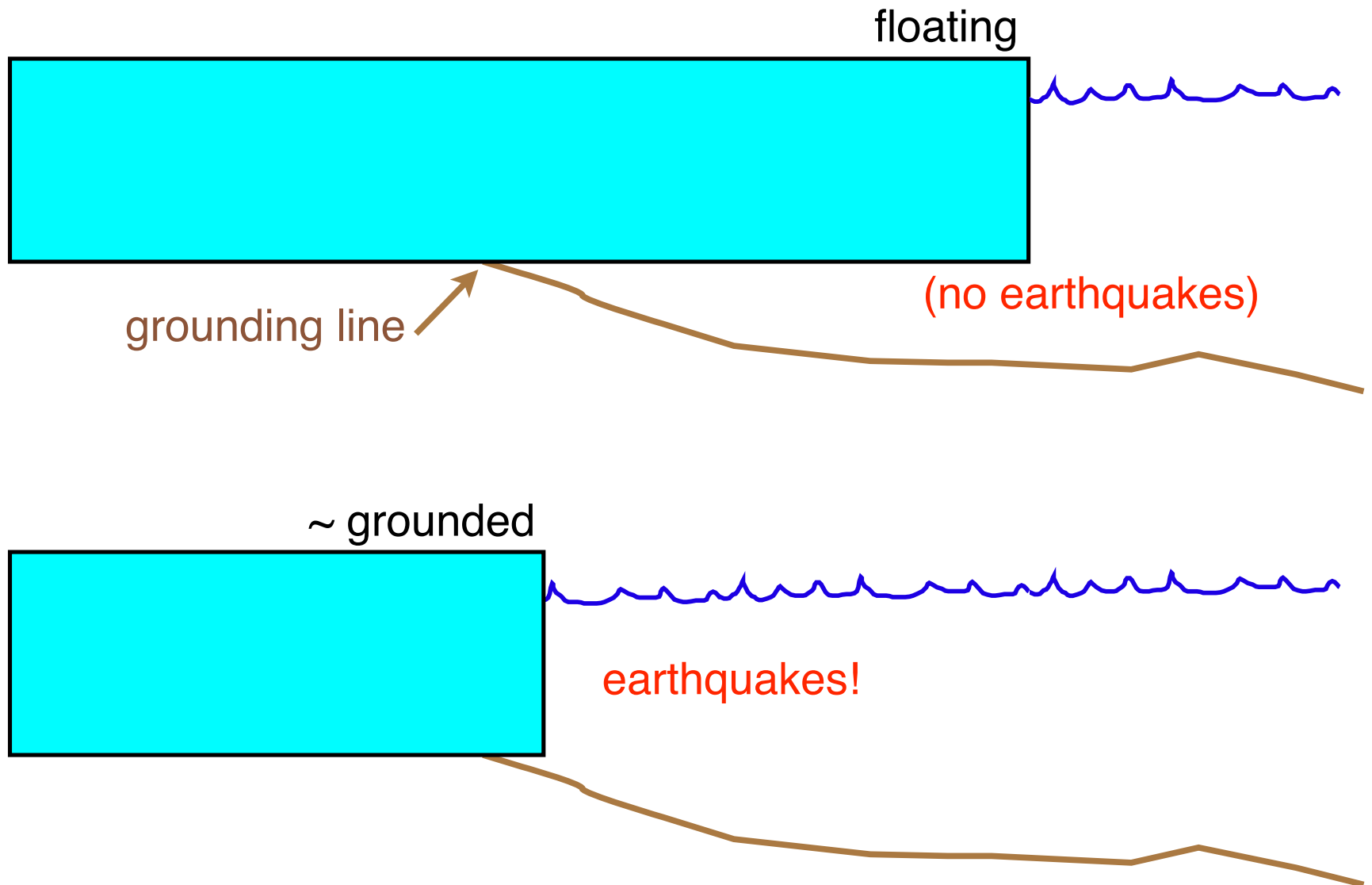


Tide gauge record  
(filtered 200-4000 s)

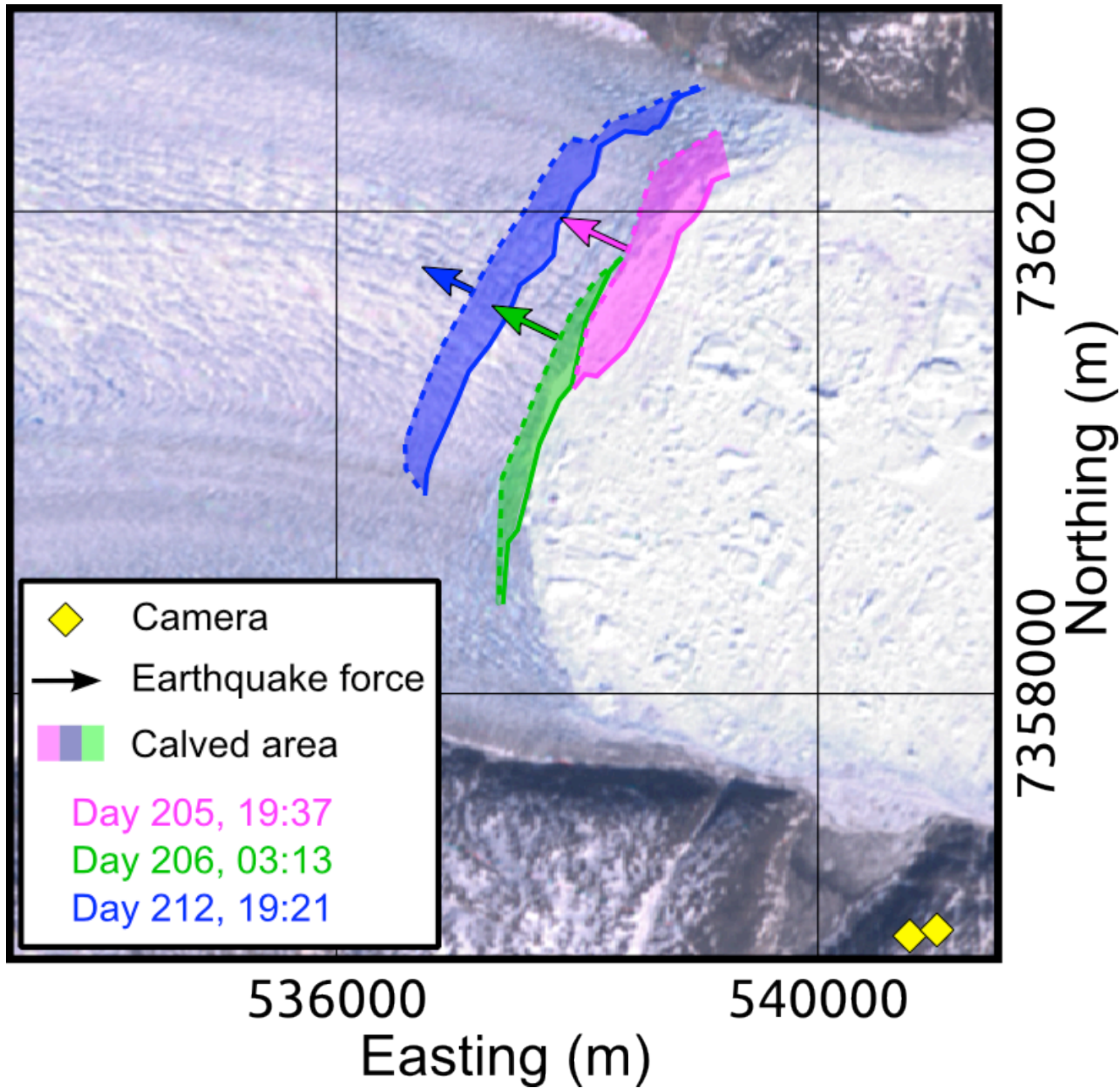




# Earthquakes occur when calving front is nearly grounded

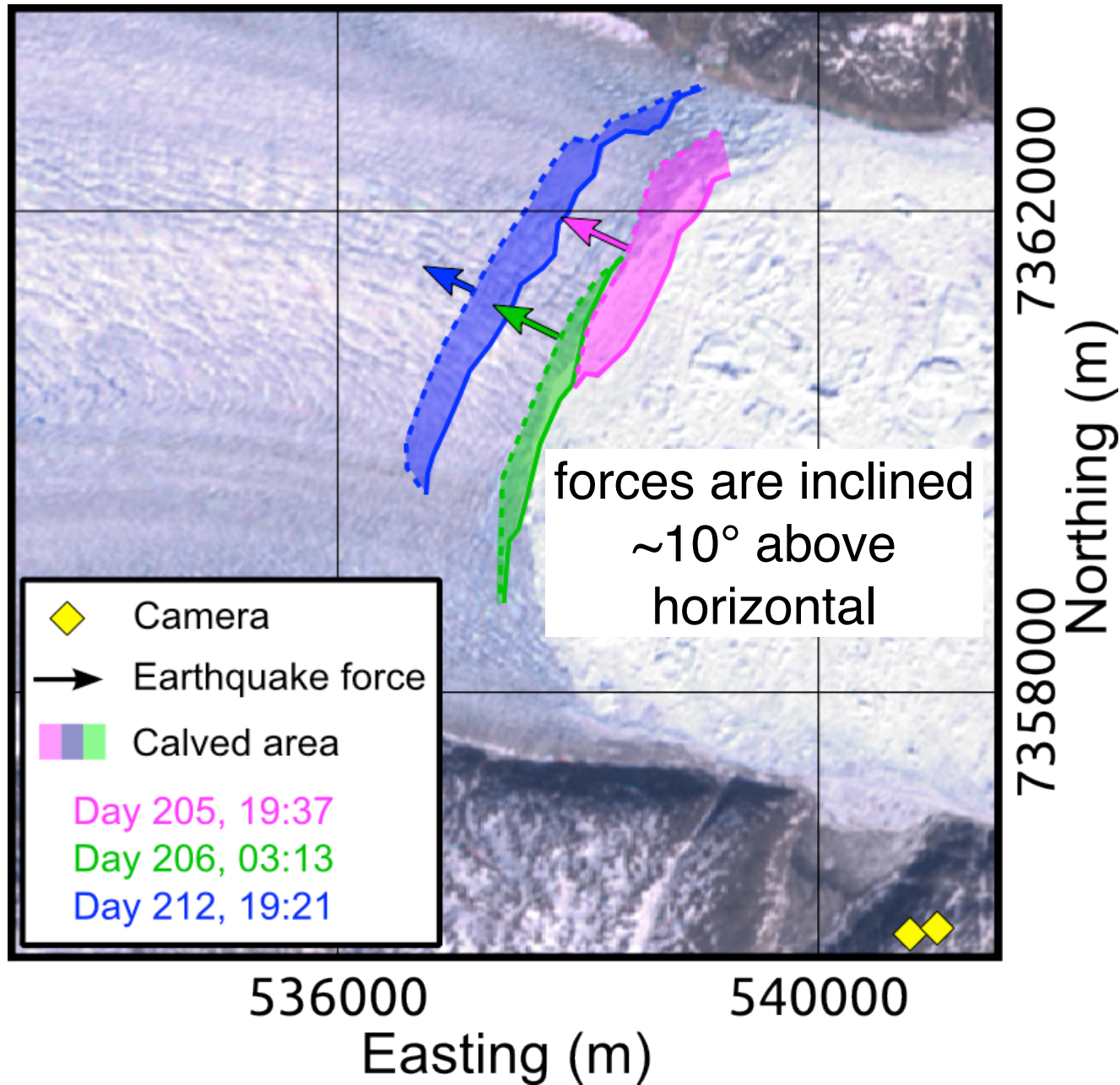


# Forces in detail: geometry of calved blocks from stereo camera imagery



*thanks to T. Murray, N. Selmes, T.D. James*

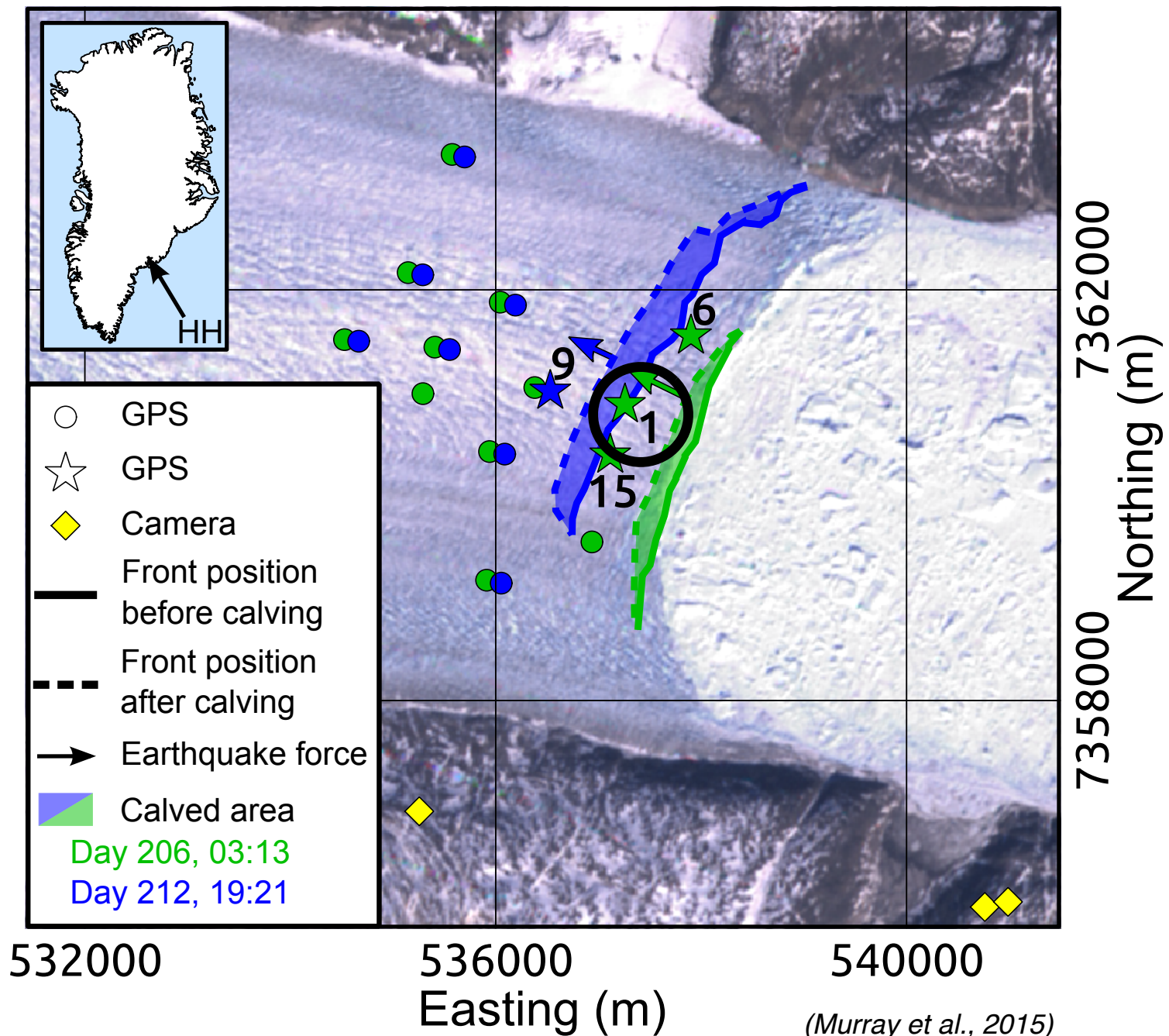
# Forces in detail: geometry of calved blocks from stereo camera imagery



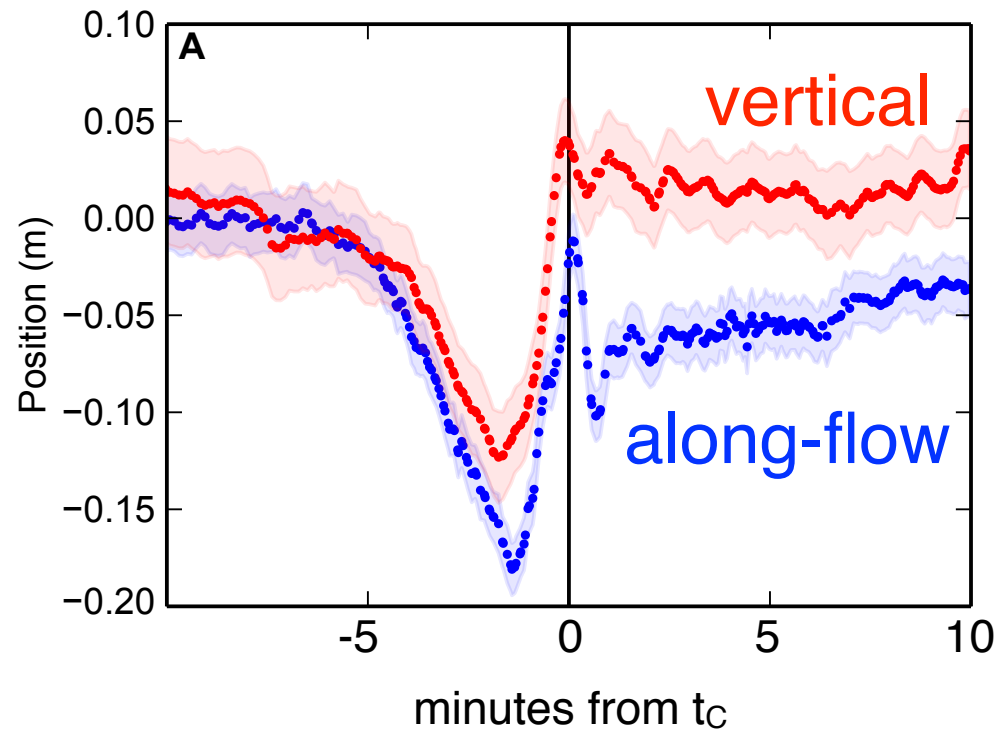
*thanks to T. Murray, N. Selmes, T.D. James*



# GPS receivers record glacier motion



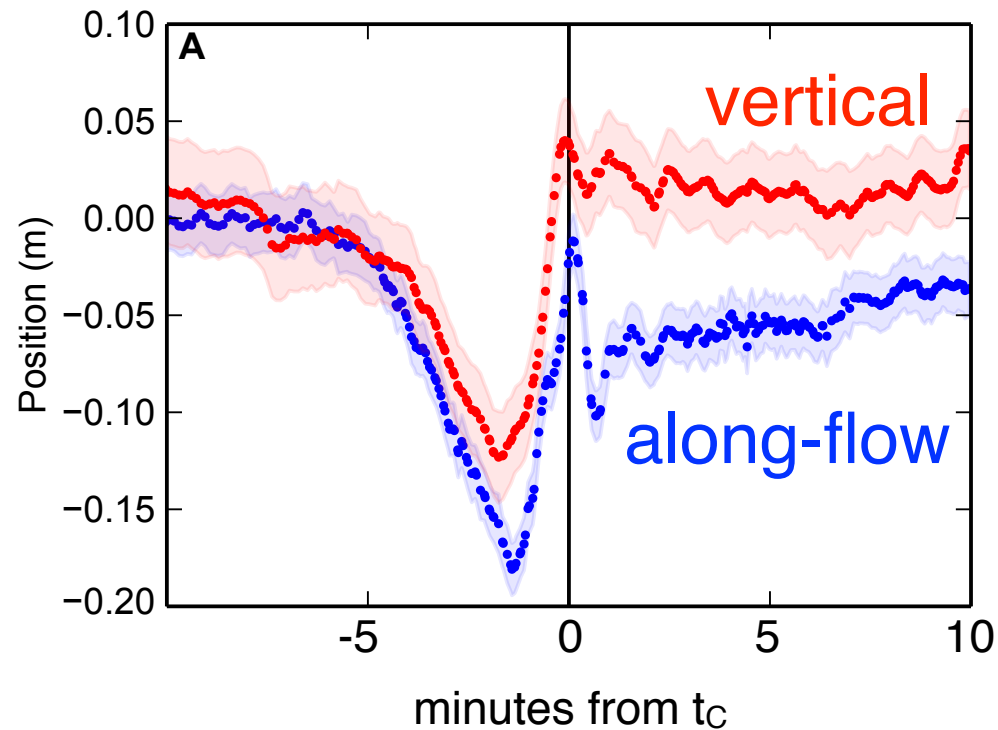
## GPS sensor 1, doy 206



glacier displacement  
direction *reverses* during  
earthquake

↑  
earthquake centroid time

## GPS sensor 1, doy 206

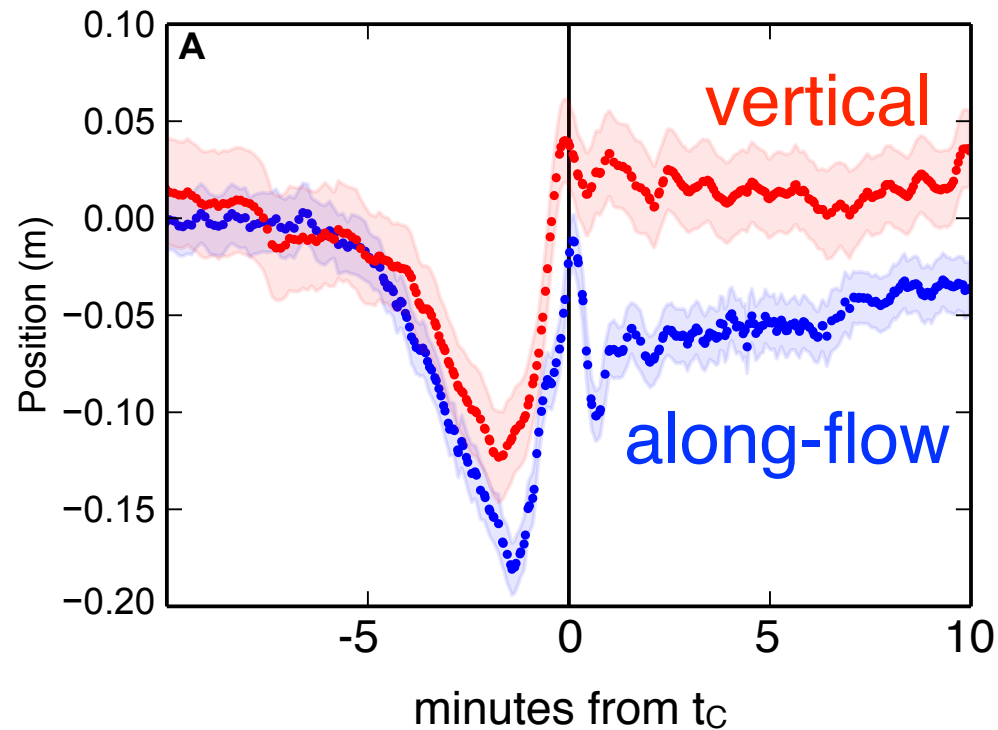


glacier displacement  
direction *reverses* during  
earthquake, and  
terminus is drawn down

↑  
earthquake centroid time

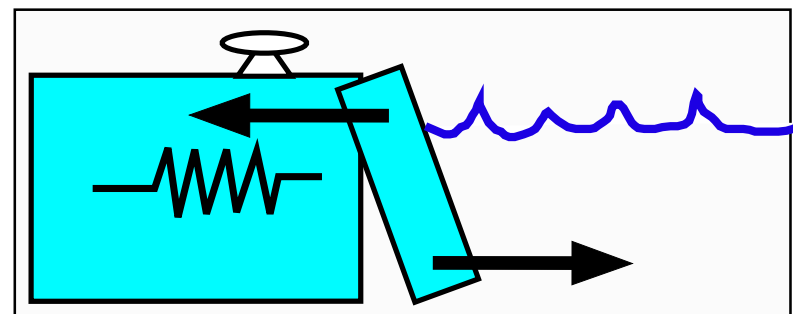


## GPS sensor 1, doy 206

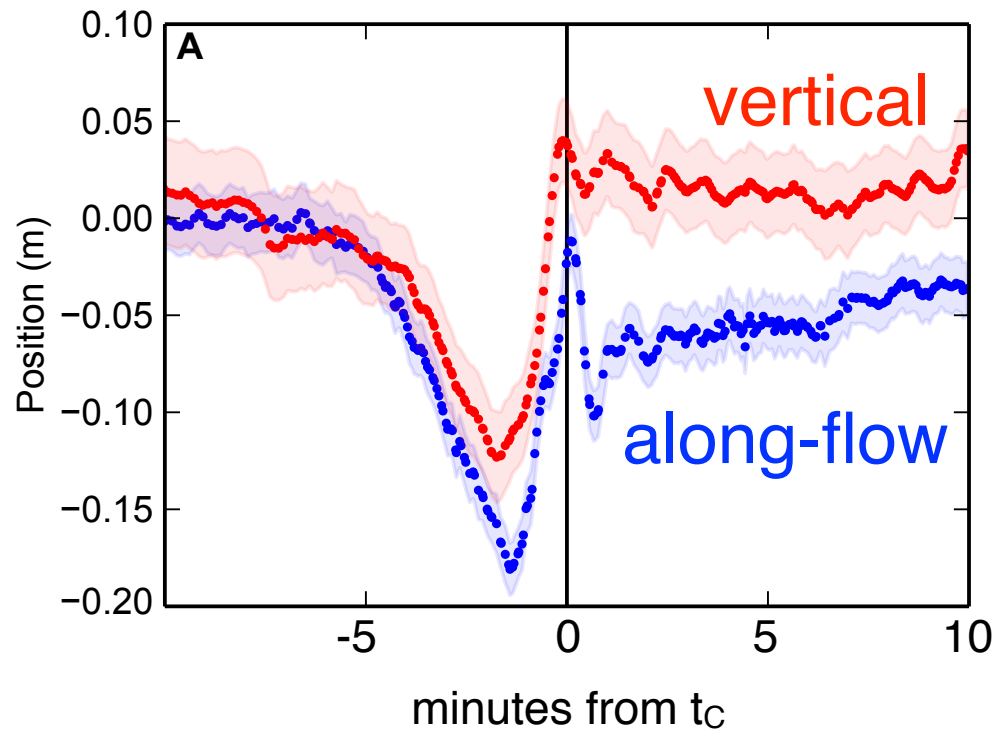


earthquake centroid time

glacier displacement  
direction *reverses* during  
earthquake

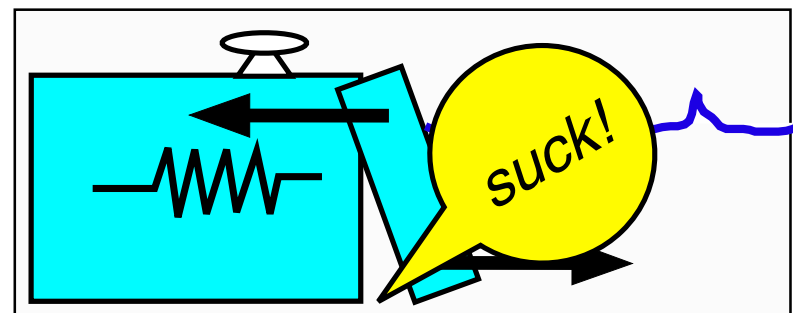


## GPS sensor 1, doy 206

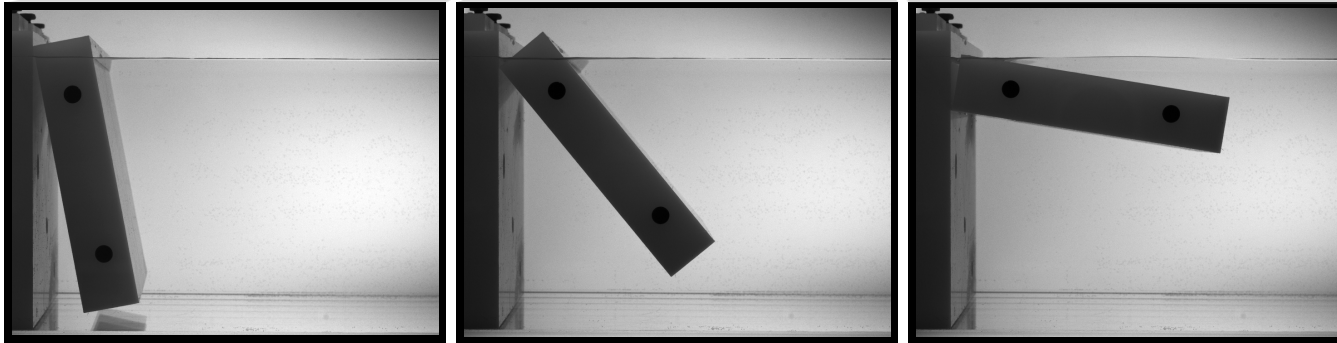


earthquake centroid time

glacier displacement  
direction *reverses* during  
earthquake, and  
terminus is drawn down



## Tank experiments: plastic icebergs capsizing in water



measure force on “glacier”, pressure at “terminus”

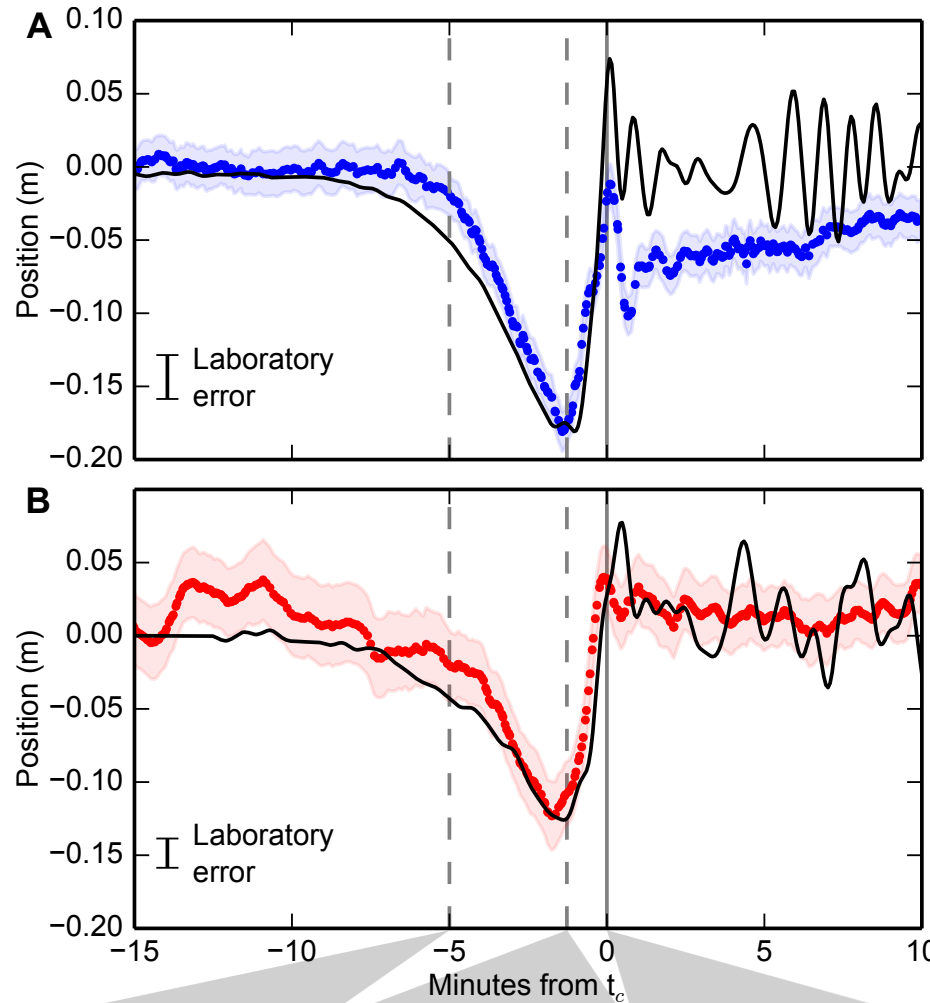
*(Murray et al., 2015)*  
*\*Mac Cathles, Justin Burton*



# Tank experiments: scaled results

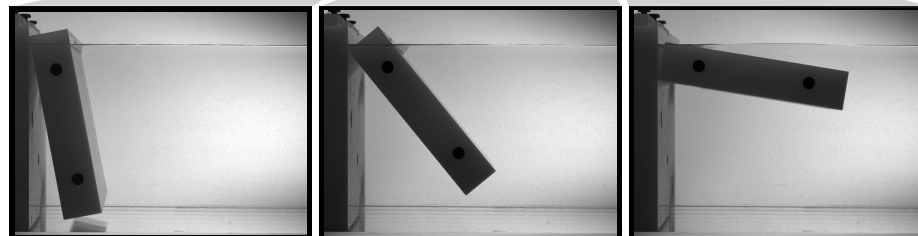
elastic  
compression of  
glacier front

drawdown of  
terminus by  
dynamic  
pressure drop



GPS along-flow

GPS vertical

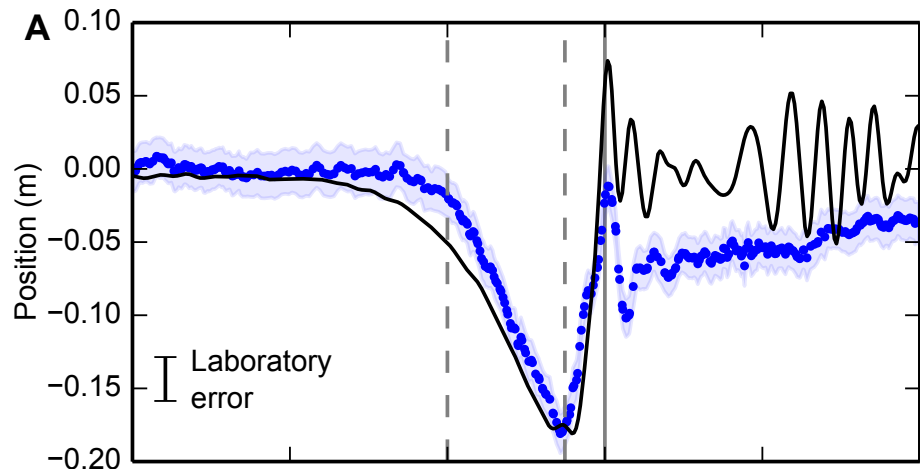


(Murray et al., 2015)

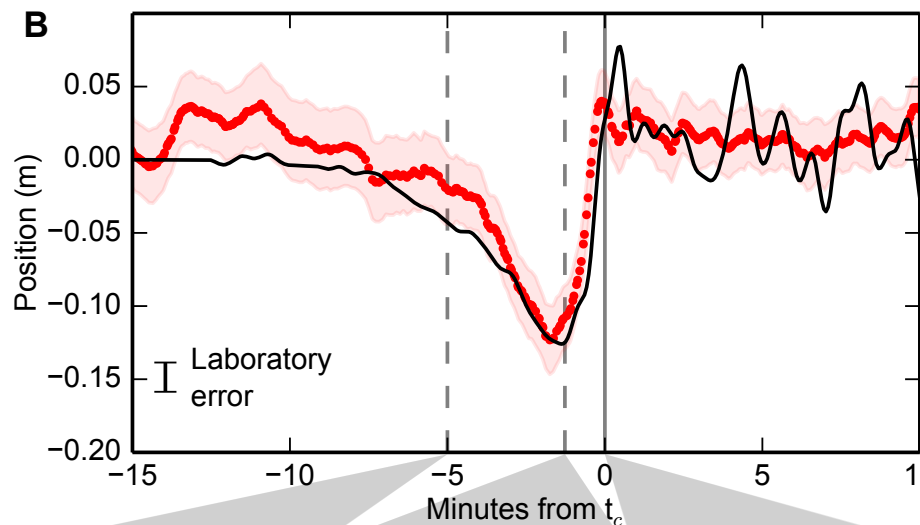
# Tank experiments: scaled results

elastic  
compression of  
glacier front

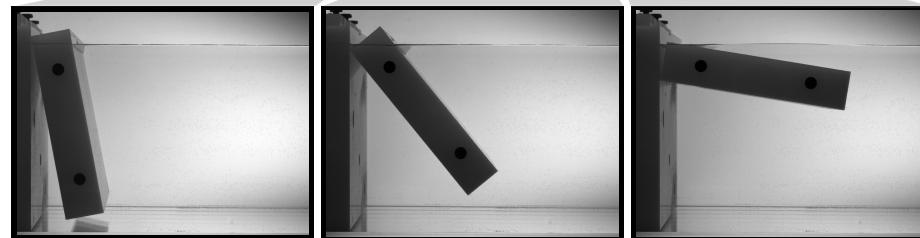
drawdown of  
terminus by  
dynamic  
pressure drop  
(force *up* on  
solid Earth)



GPS along-flow



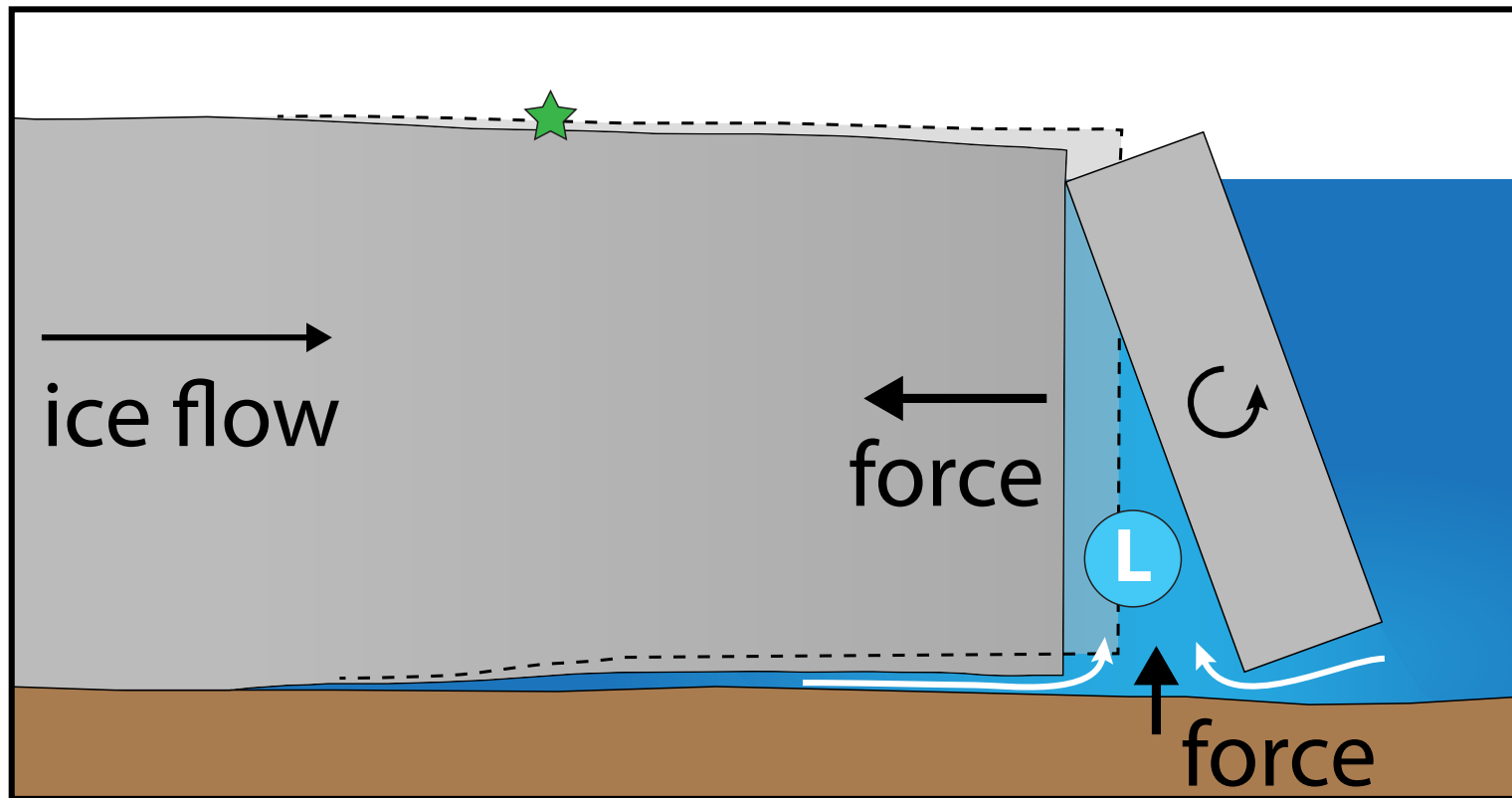
GPS vertical



(Murray et al., 2015)

## Earthquake source:

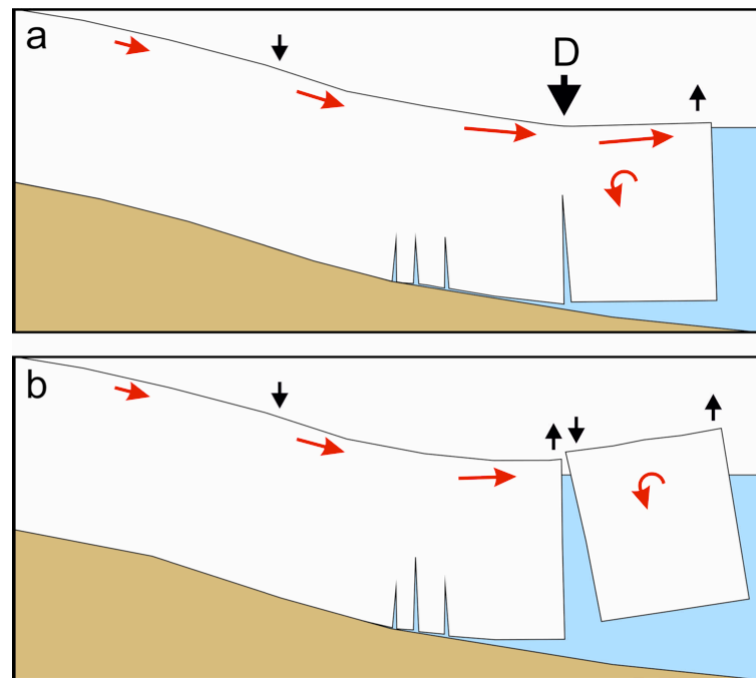
- horizontal force - reaction force due to iceberg acceleration away from glacier;
- vertical force - upward force due to dynamic pressure drop behind capsizing iceberg



(Murray et al., 2015)



# Calving of (nearly) grounded ice: buoyancy-driven calving

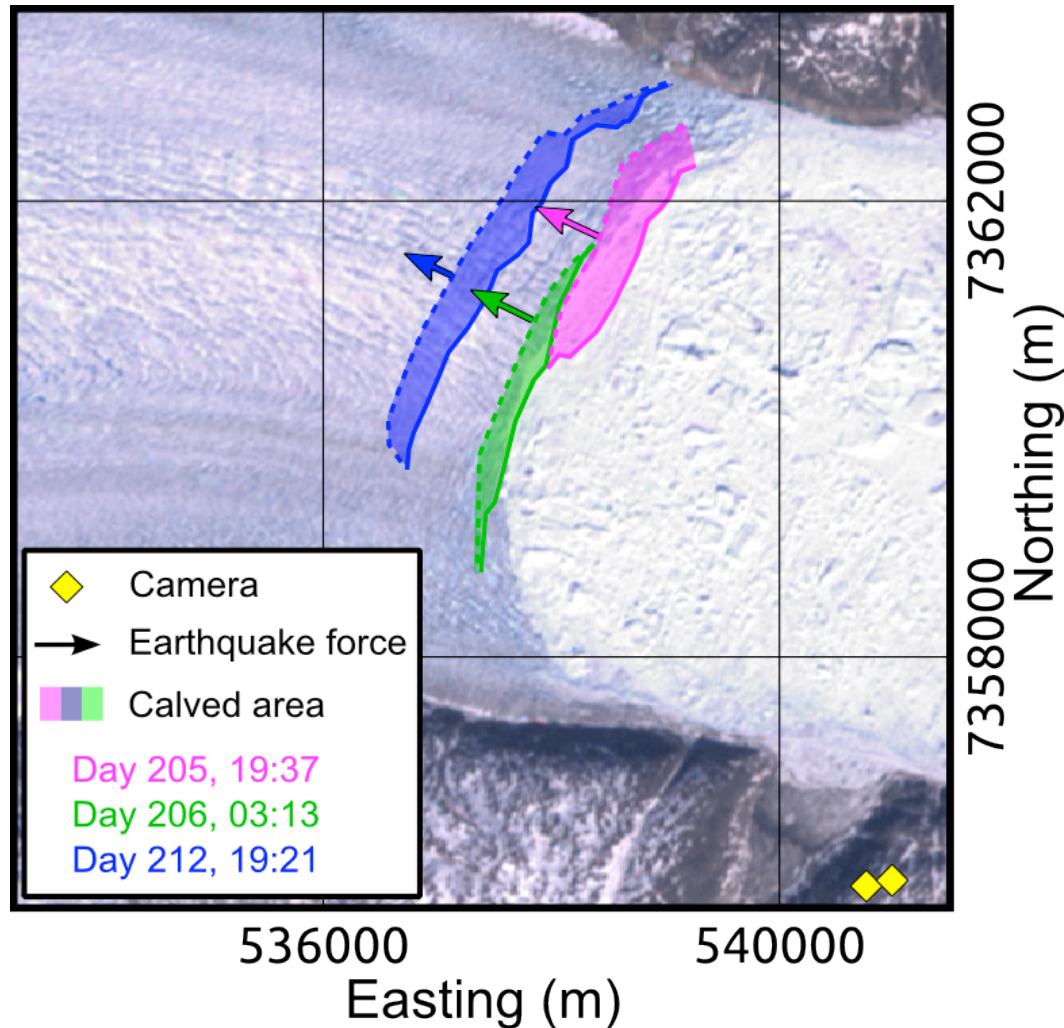


(Murray et al., 2014)

How many calving events are detected as glacial earthquakes?

How many glacial earthquakes correspond to calving events?

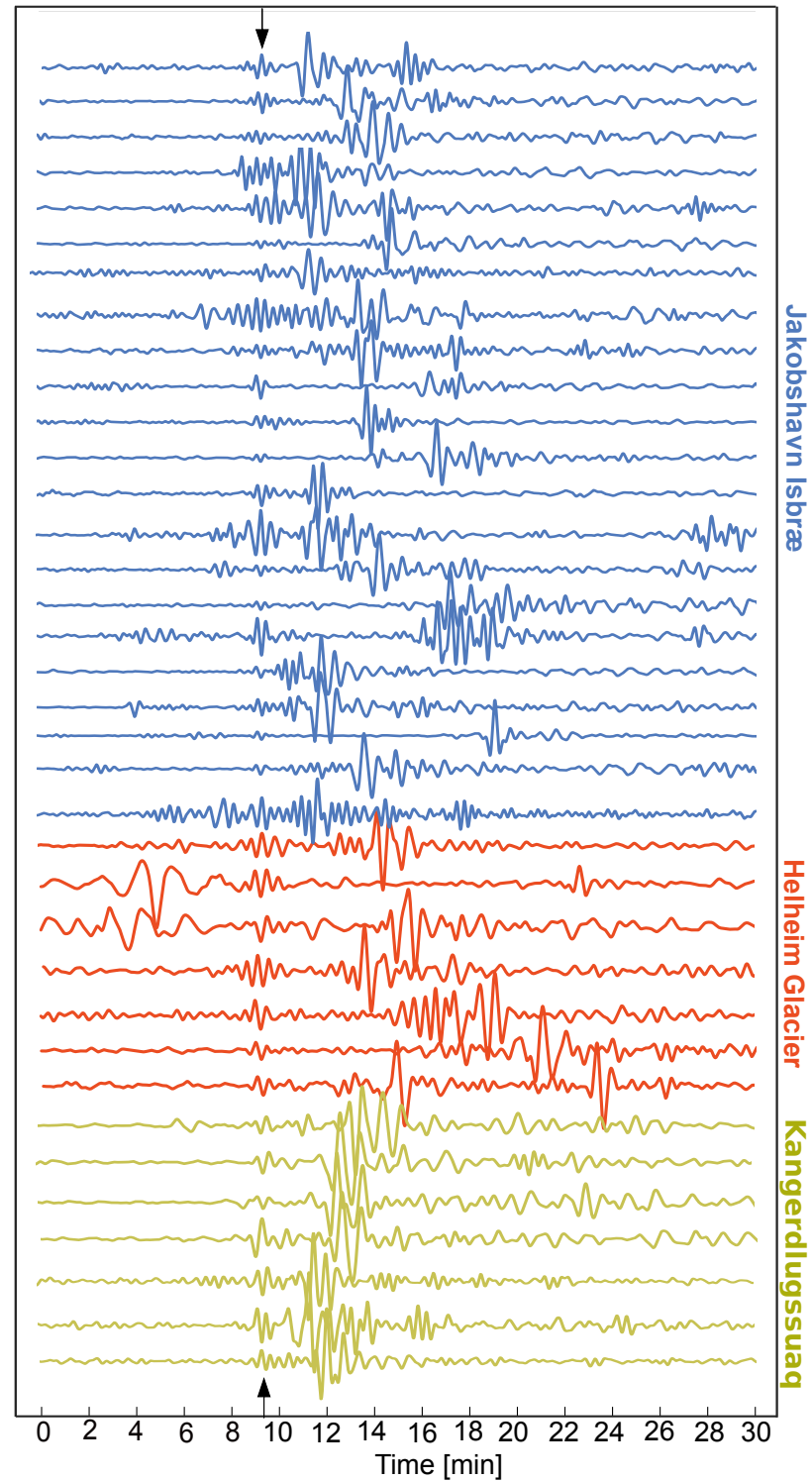
Excellent agreement in geometry, and one-to-one correspondence between visually identified calving events and earthquakes



*thanks to T.D. James, N. Selmes, T. Murray*

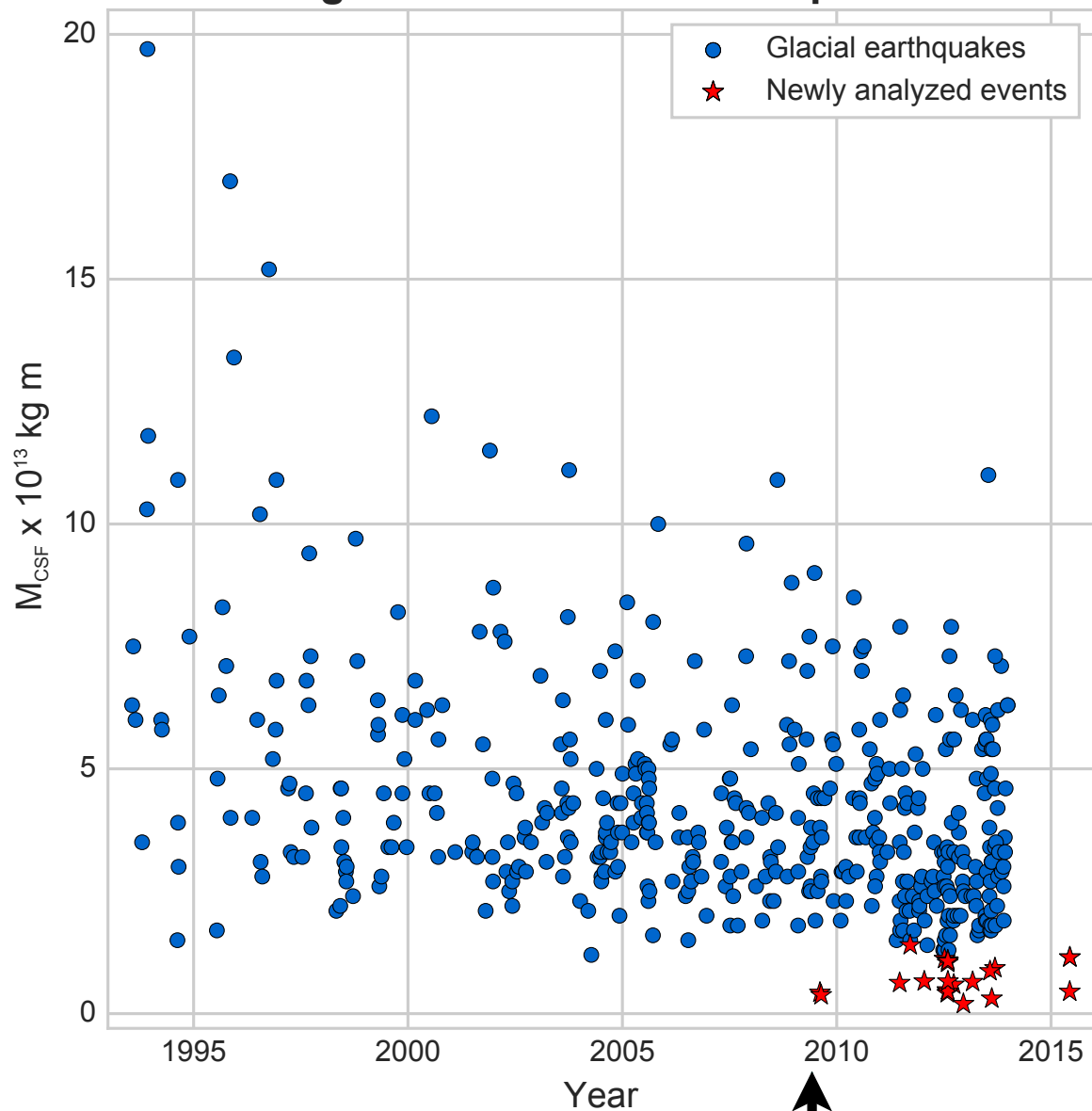


# Glacial-earthquake “precursors”, from Kira Olsen



*(ms in prep)*

# Magnitude of Glacial Earthquakes



glacial-earthquake  
precursors:  
smaller calving events

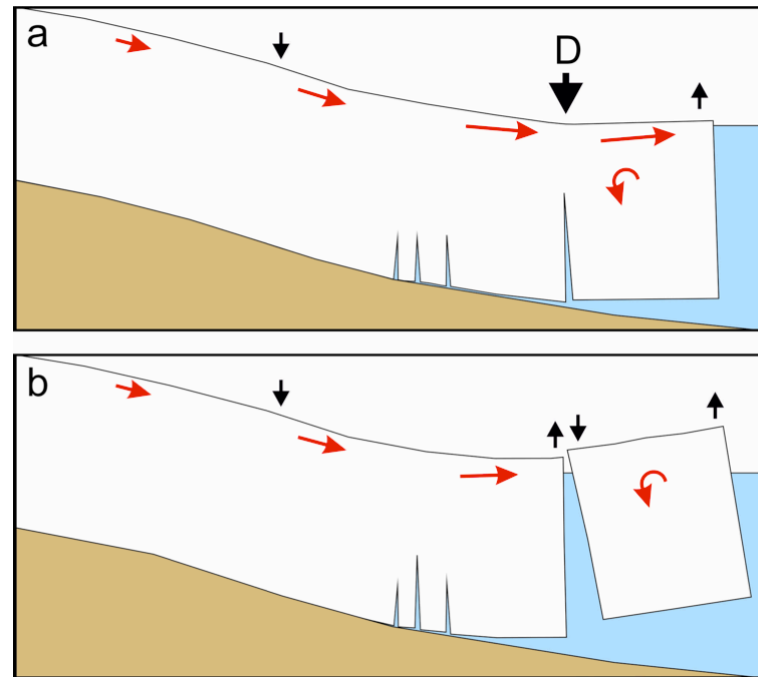
July, 2009

First GLISN station installed

(K. Olsen)

*(ms in prep)*

Buoyancy-driven calving likely accounts for almost all dynamic mass loss during the calving (~summer) season at nearly grounded glaciers



(Murray et al., 2014)



# Summary

- Solid Earth, hydrology, and ice dynamics are linked
- Seismology has a contribution to make at these boundaries
- Requires: high-quality data; high-quality data archiving and curation; interdisciplinary data and collaborations