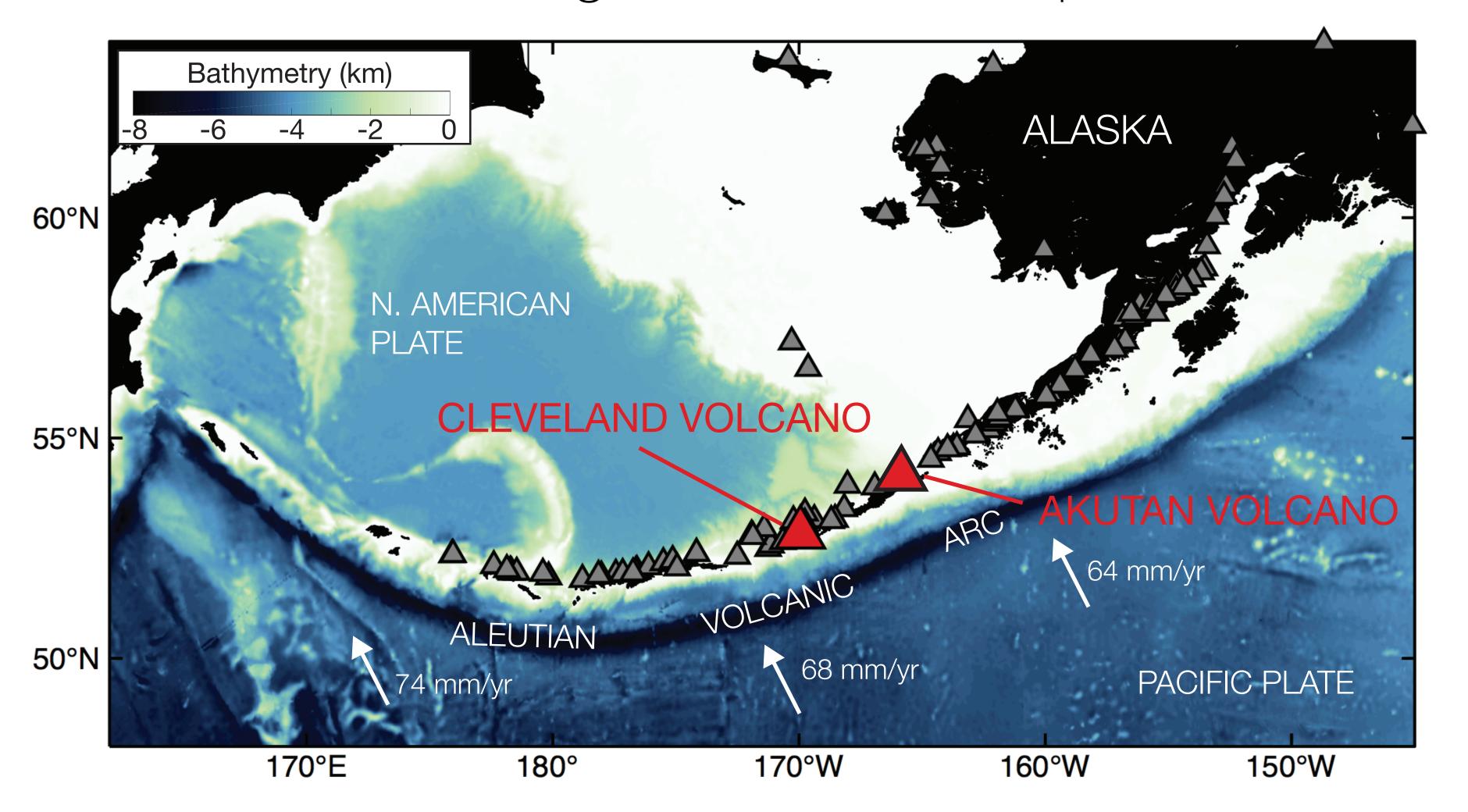
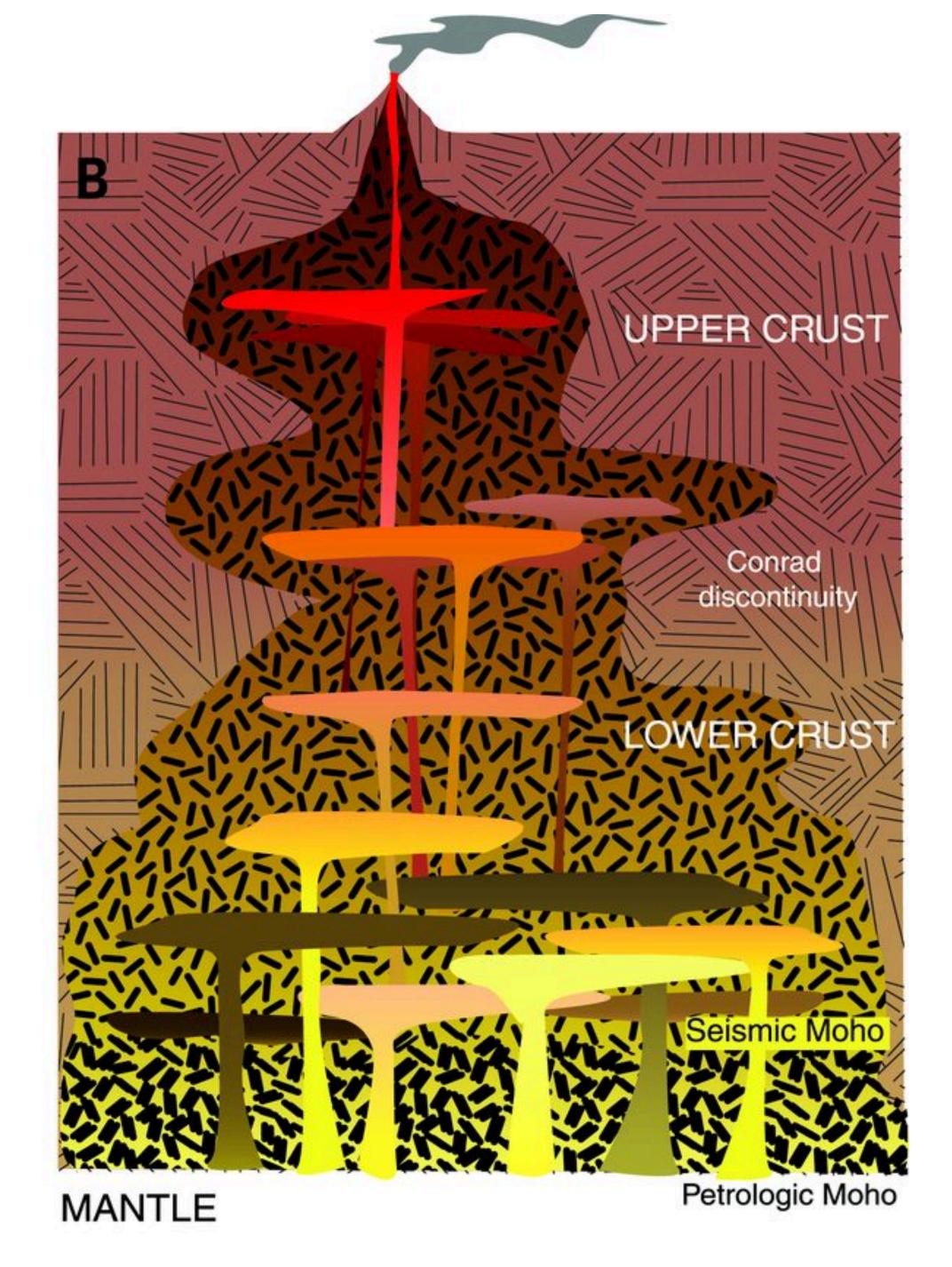
Imaging Trans-Crustal Magmatic Systems with Receiver Functions A Path Towards Linking Tectonic and Eruptive Processes



Helen A. Janiszewski | hjaniszewski@carnegiescience.edu | @helenjanisz Lara S. Wagner, Diana C. Roman SAGE/GAGE, 10 October, 2019



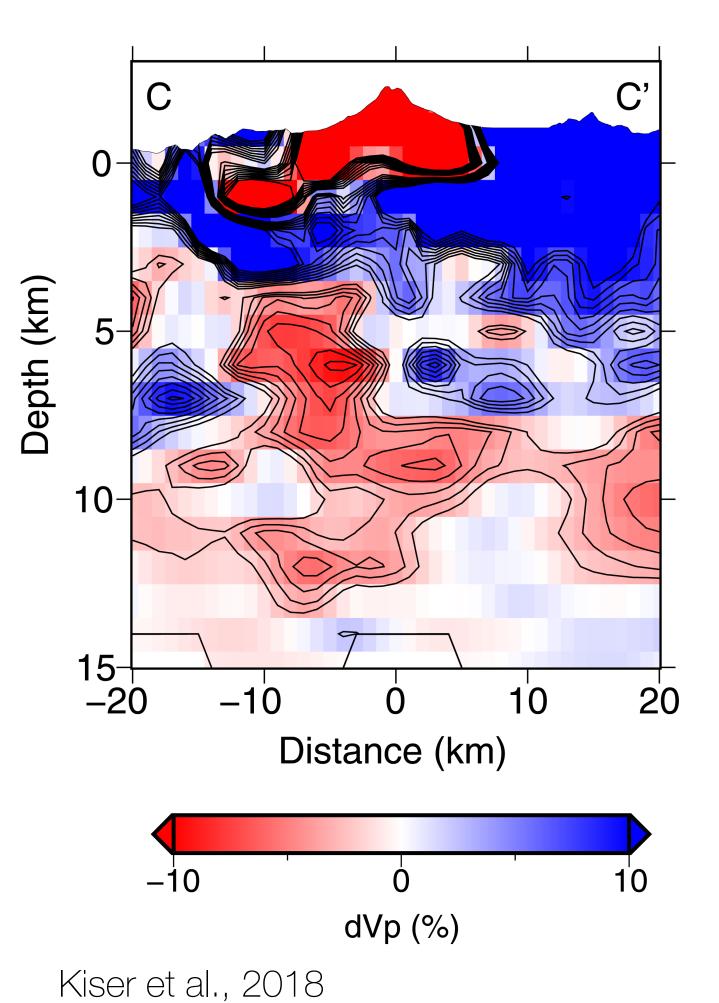
What is the crustal magmatic architecture beneath individual volcanoes?



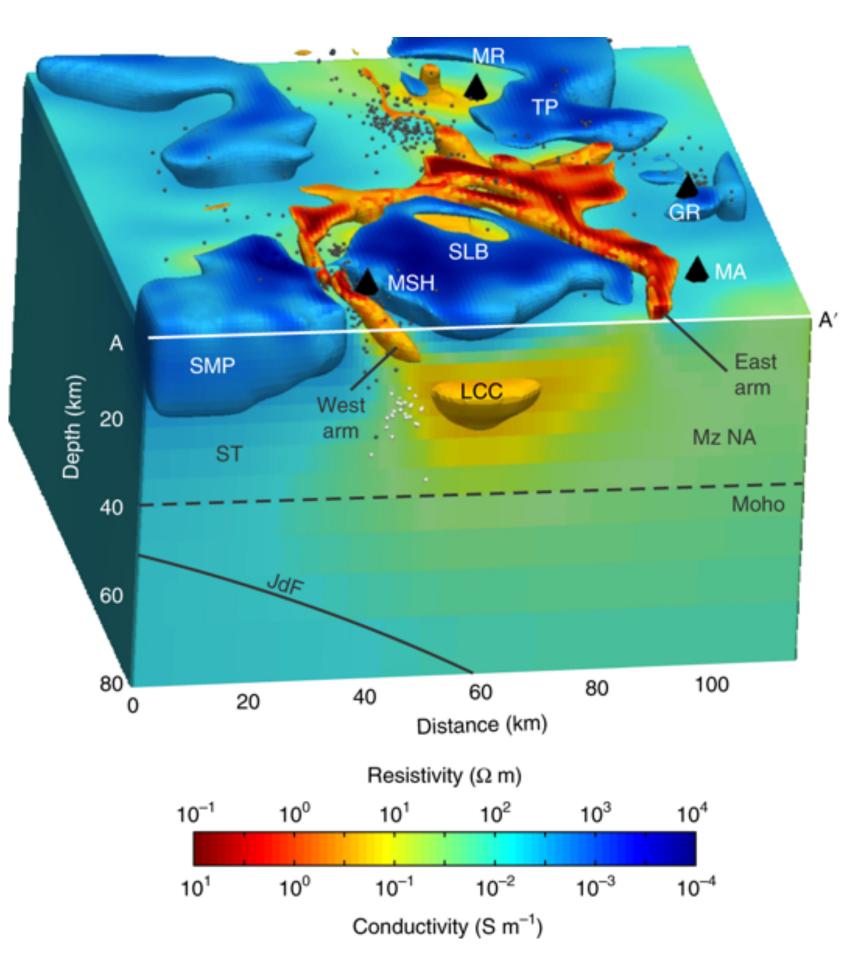
Cashman et al., 2017

What is the crustal magmatic architecture beneath individual volcanoes? Examples from iMUSH



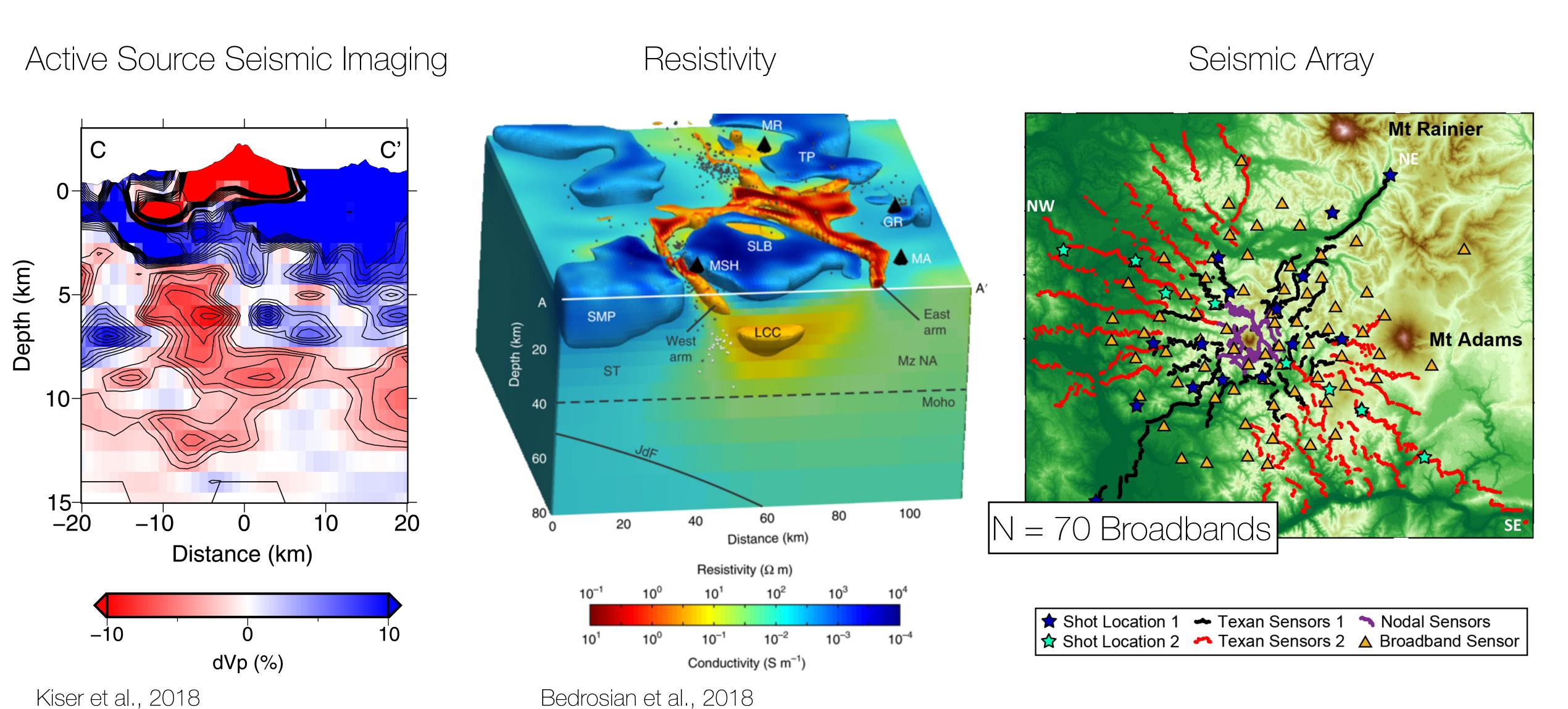


Resistivity

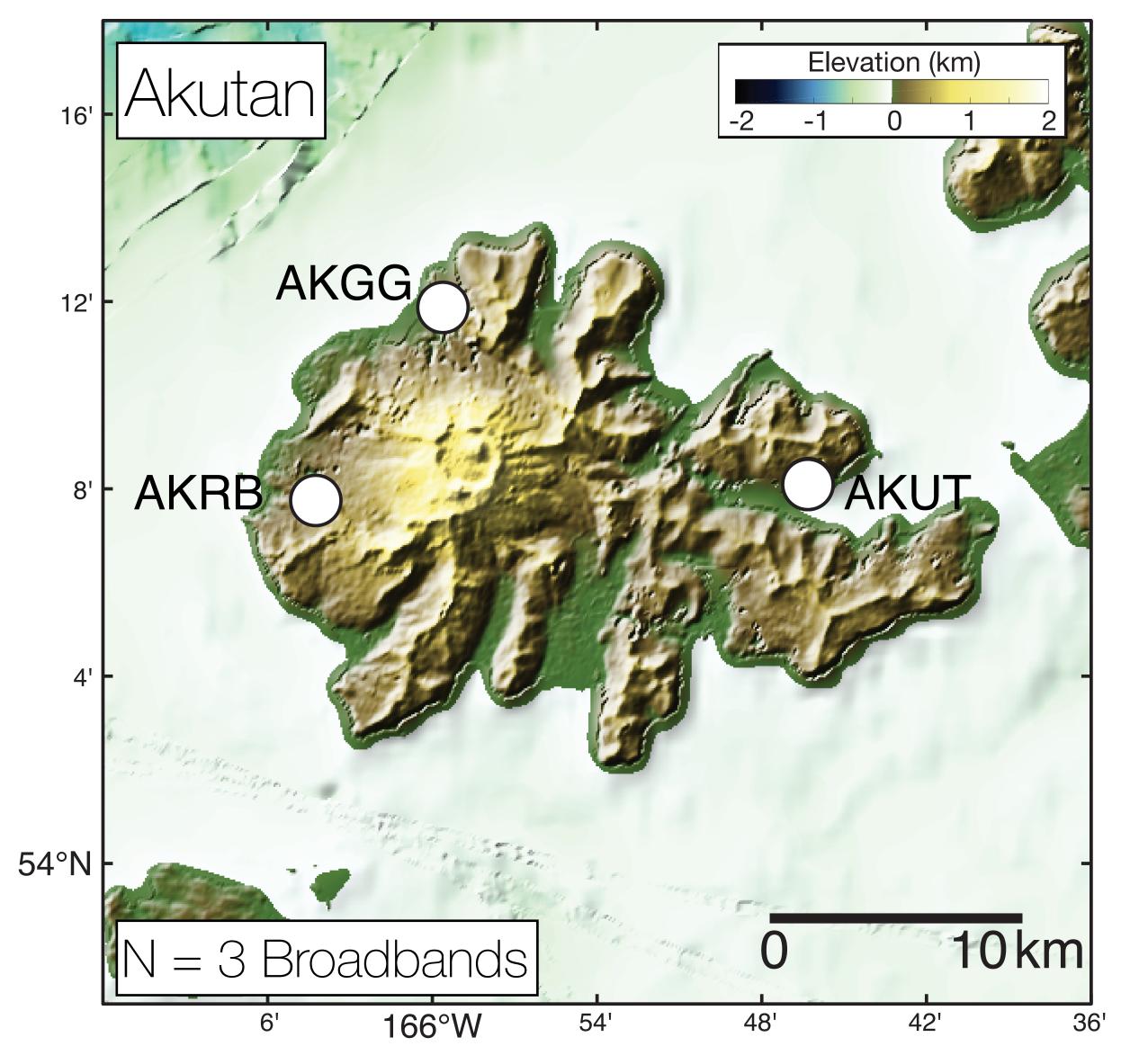


Bedrosian et al., 2018

What is the crustal magmatic architecture beneath individual volcanoes? Examples from iMUSH



What can we do with fewer resources?



Receiver function techniques can image deep crustal magmatic structure with only a few broadband instruments.

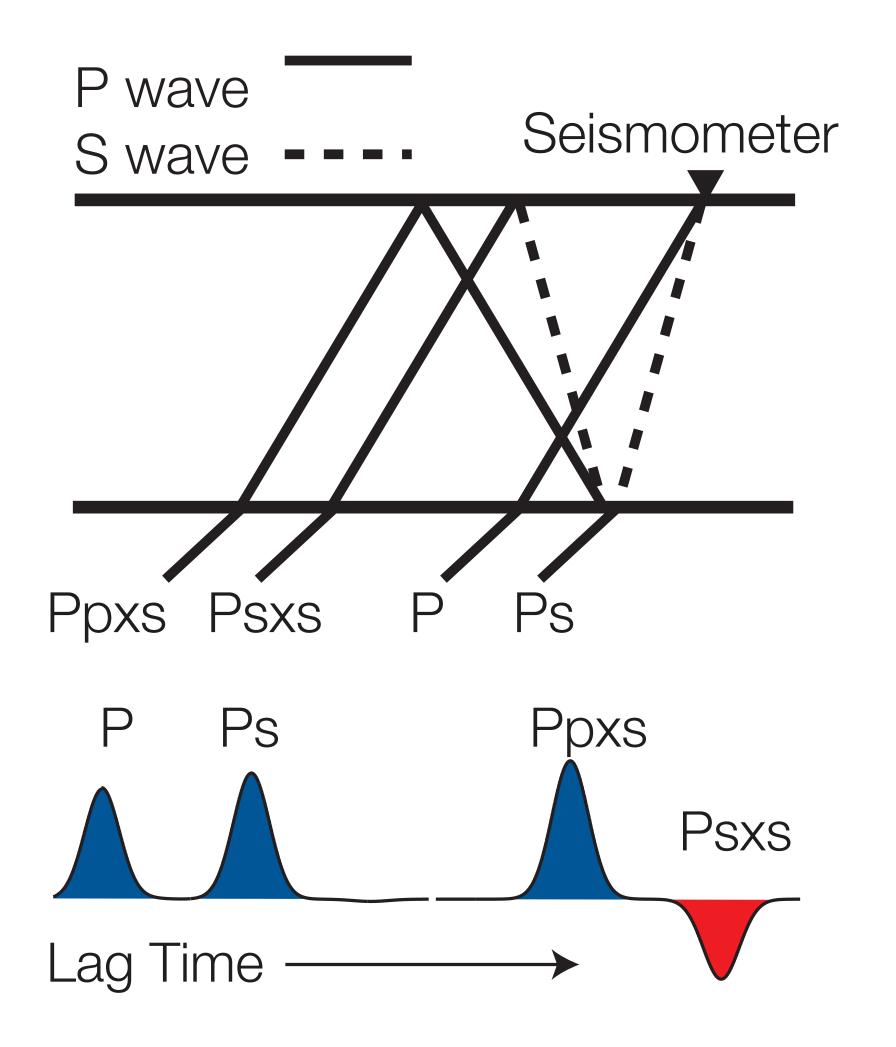
Comparable to a monitoring network.

Receiver Functions

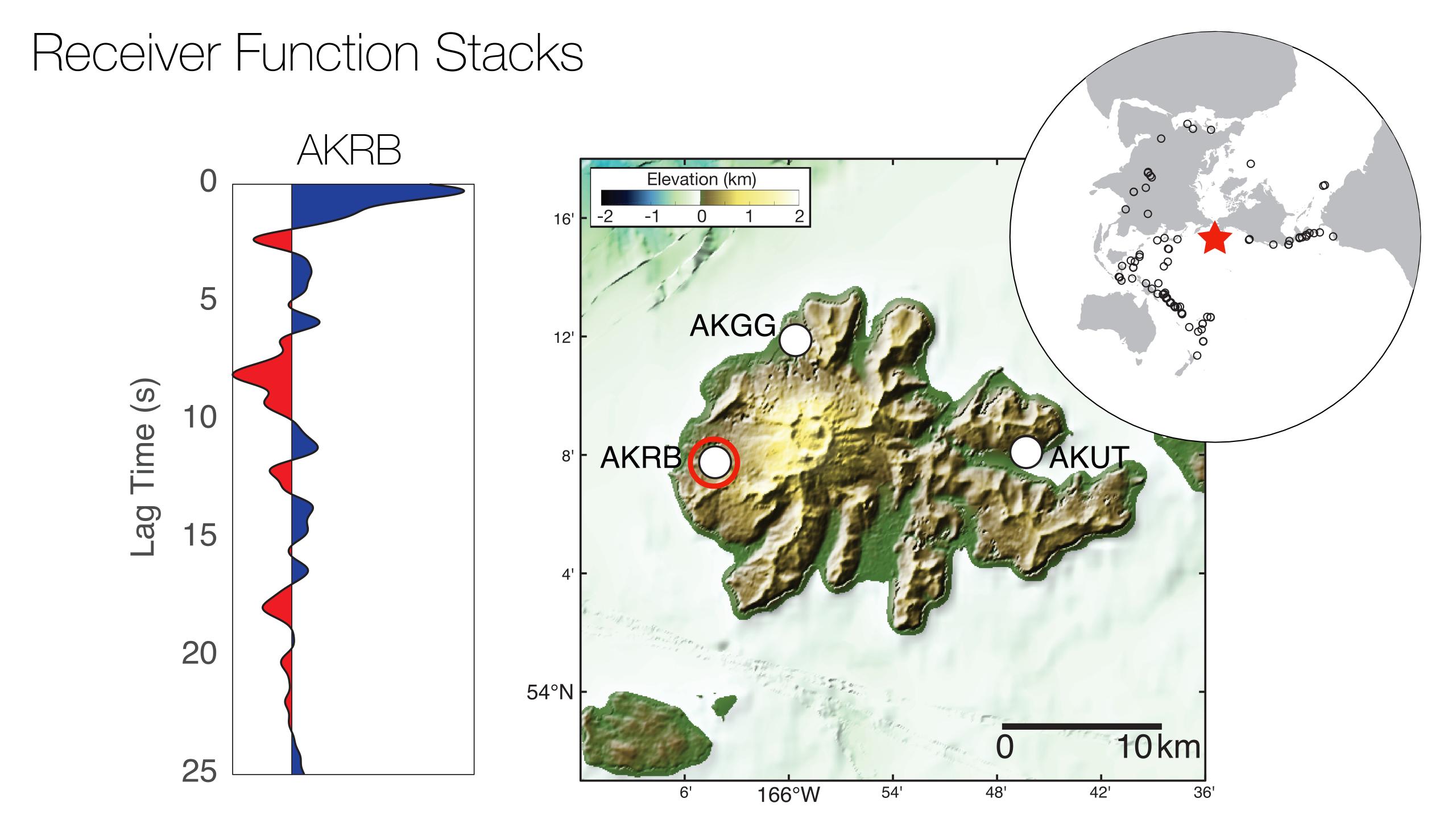
Receiver functions are sensitive to abrupt seismic velocity boundaries.

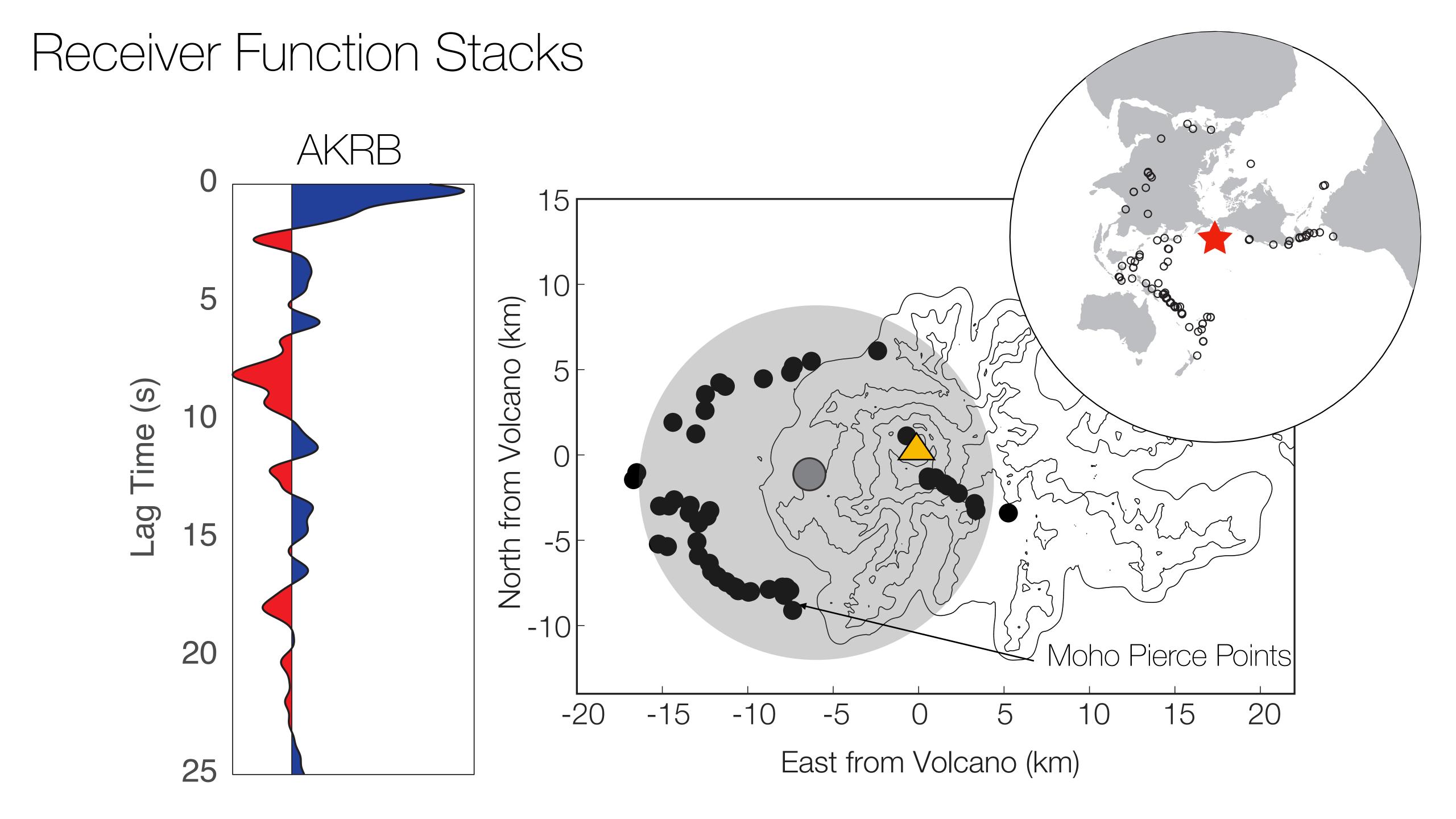
Traditionally used to image the Moho, Lithosphere-Asthenosphere Boundary, Transition Zone.

Trade off between velocity and thickness of layer.

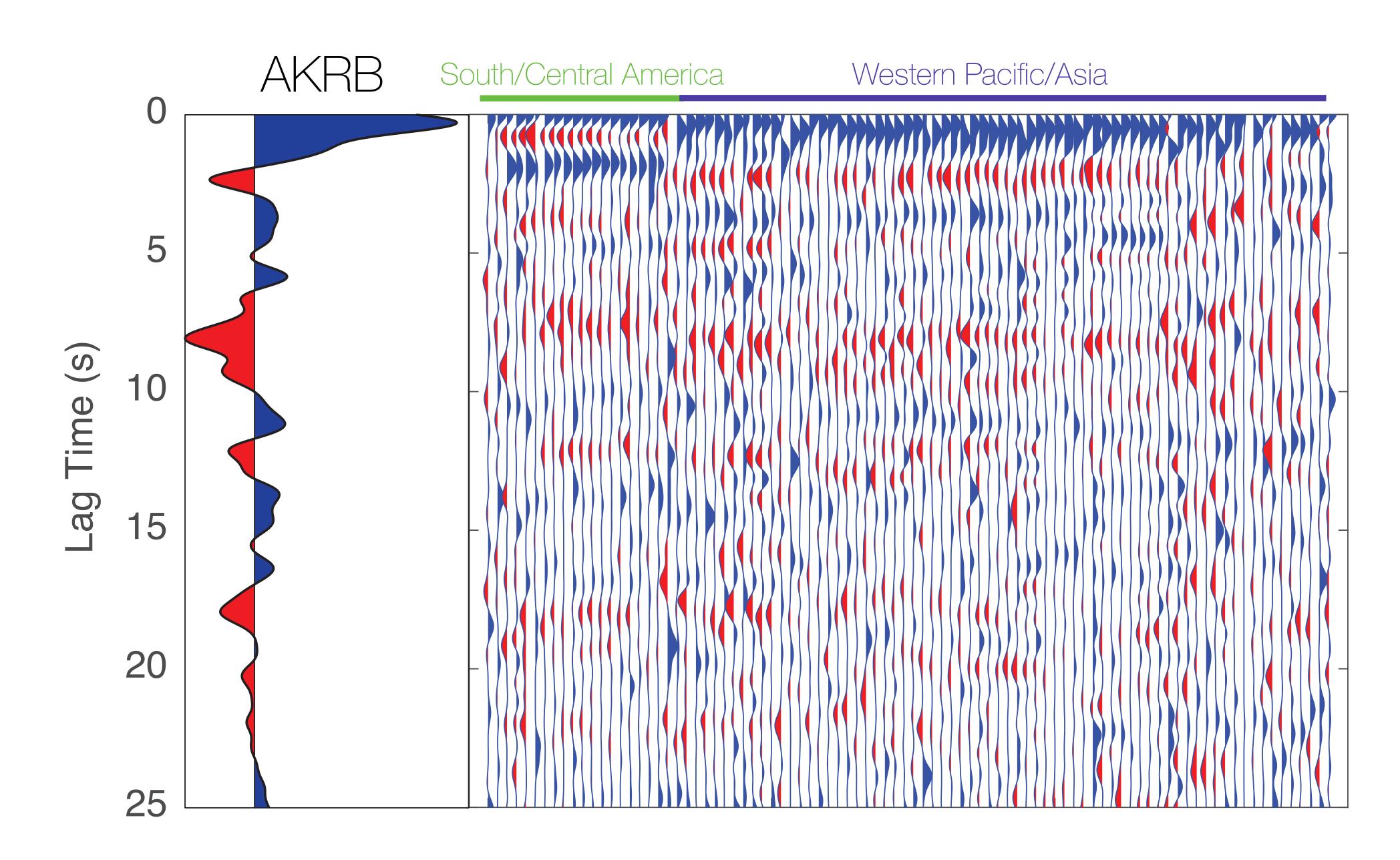


Arrival times relate to depth and velocity to boundaries

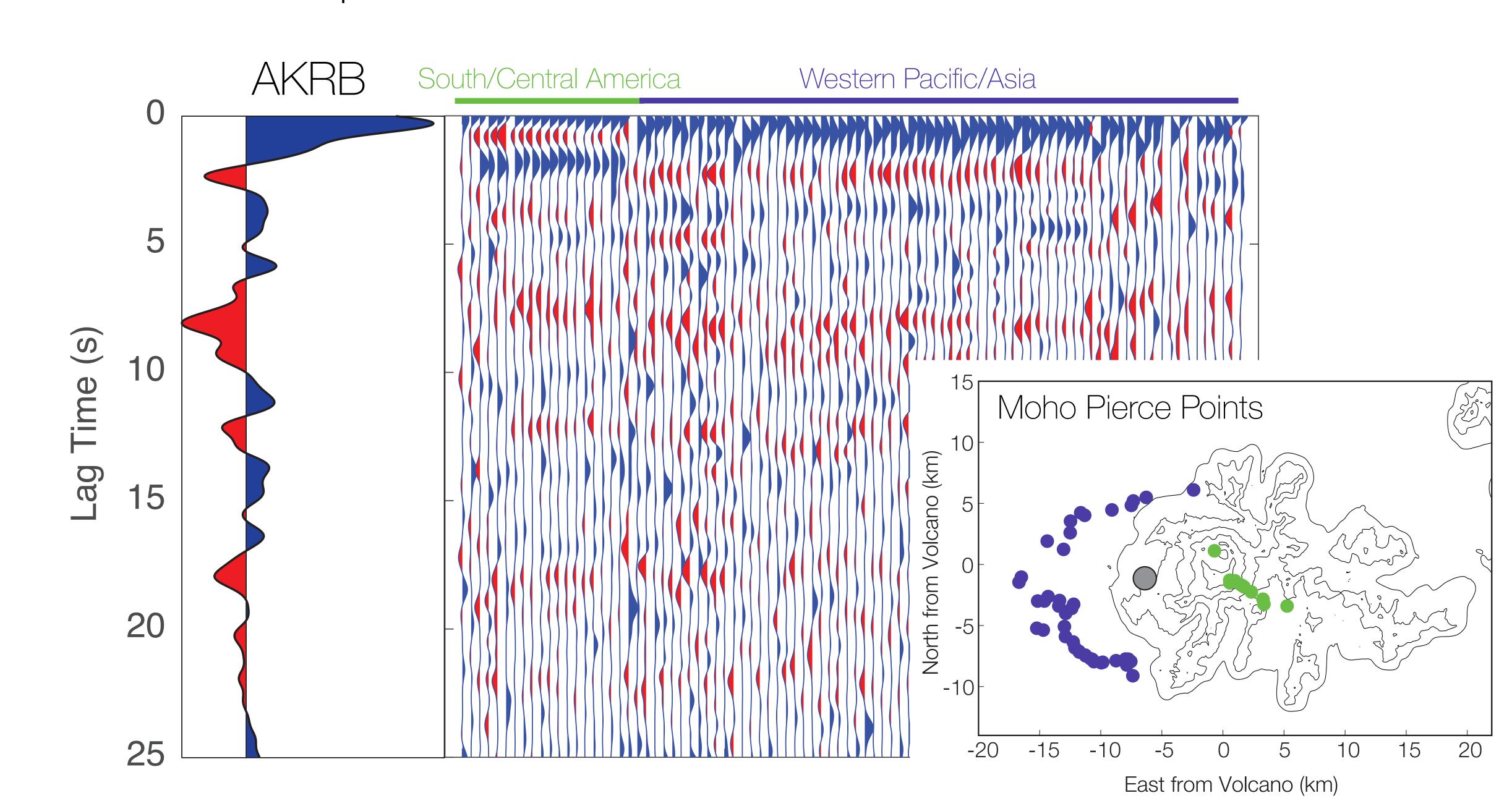




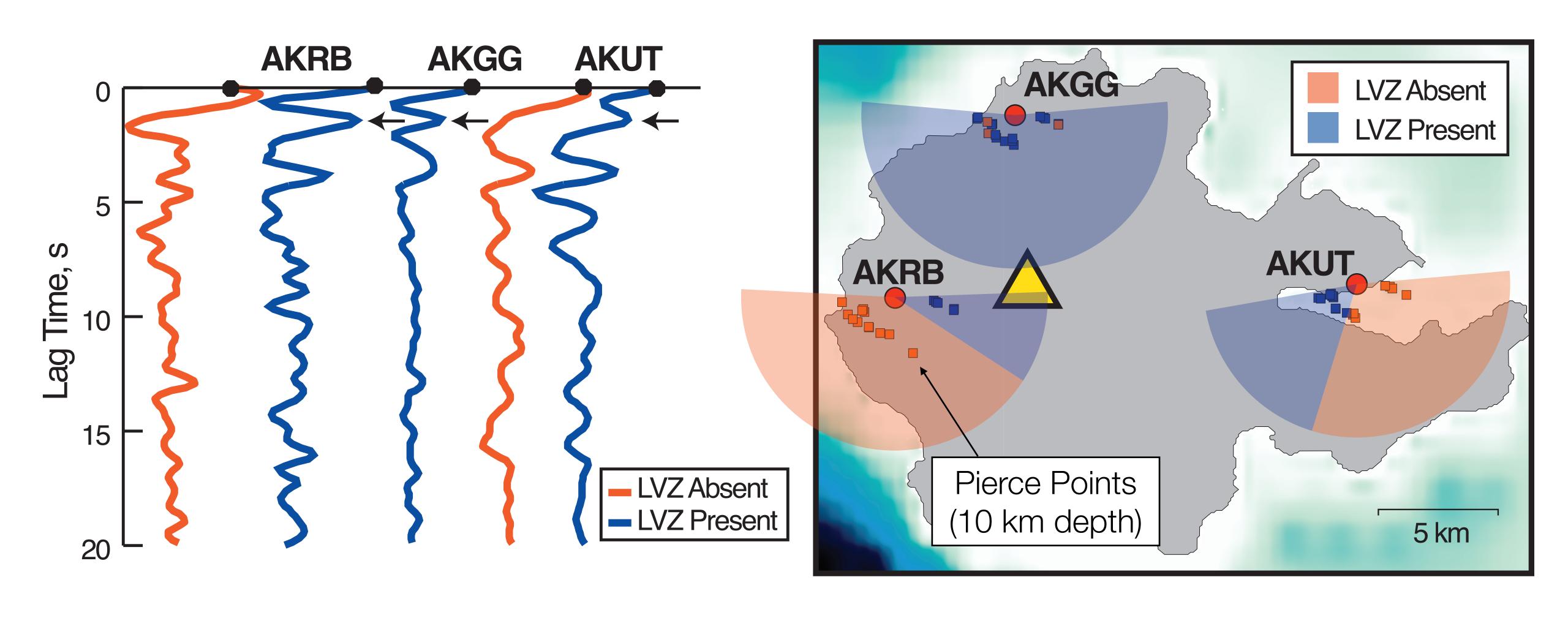
Individual Earthquake Receiver Functions



Individual Earthquake Receiver Functions

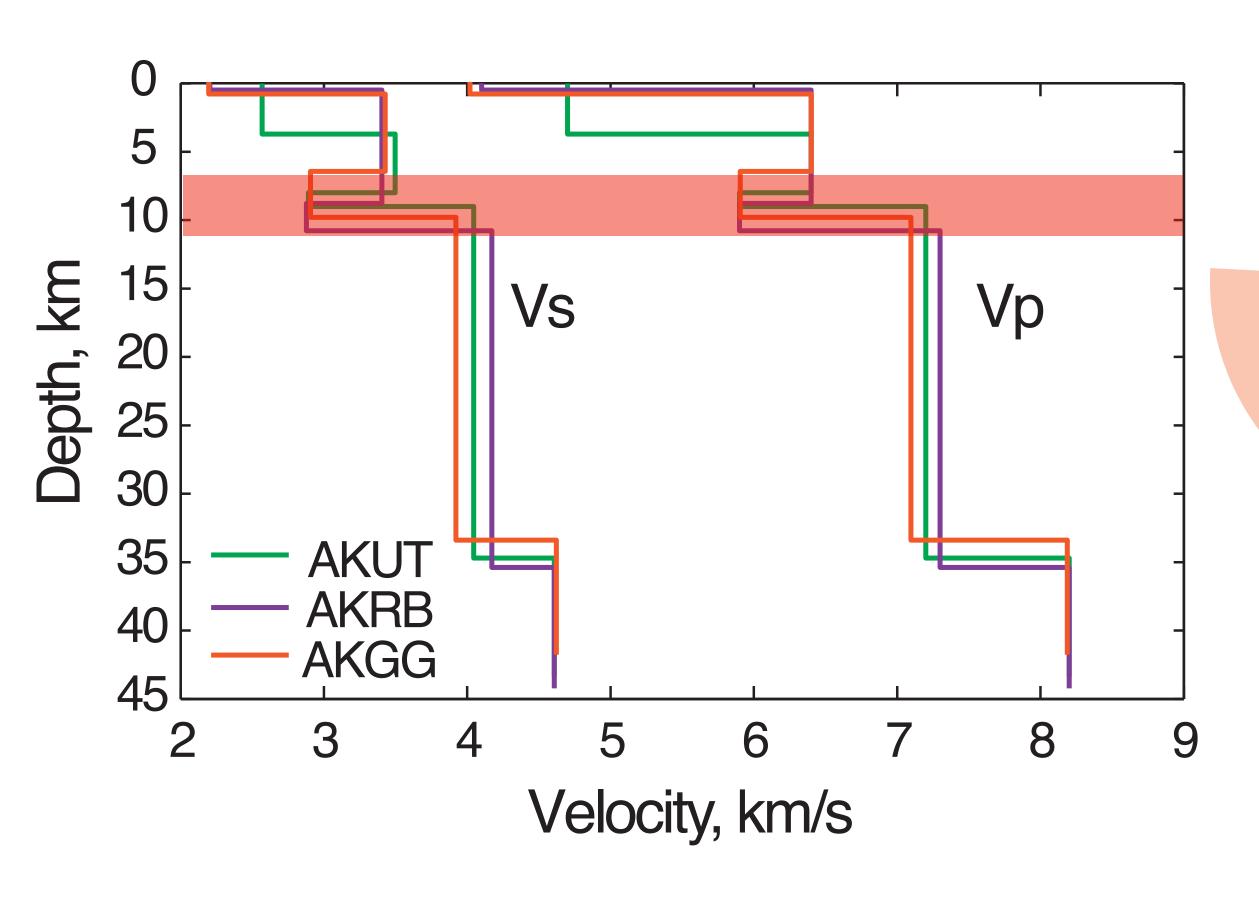


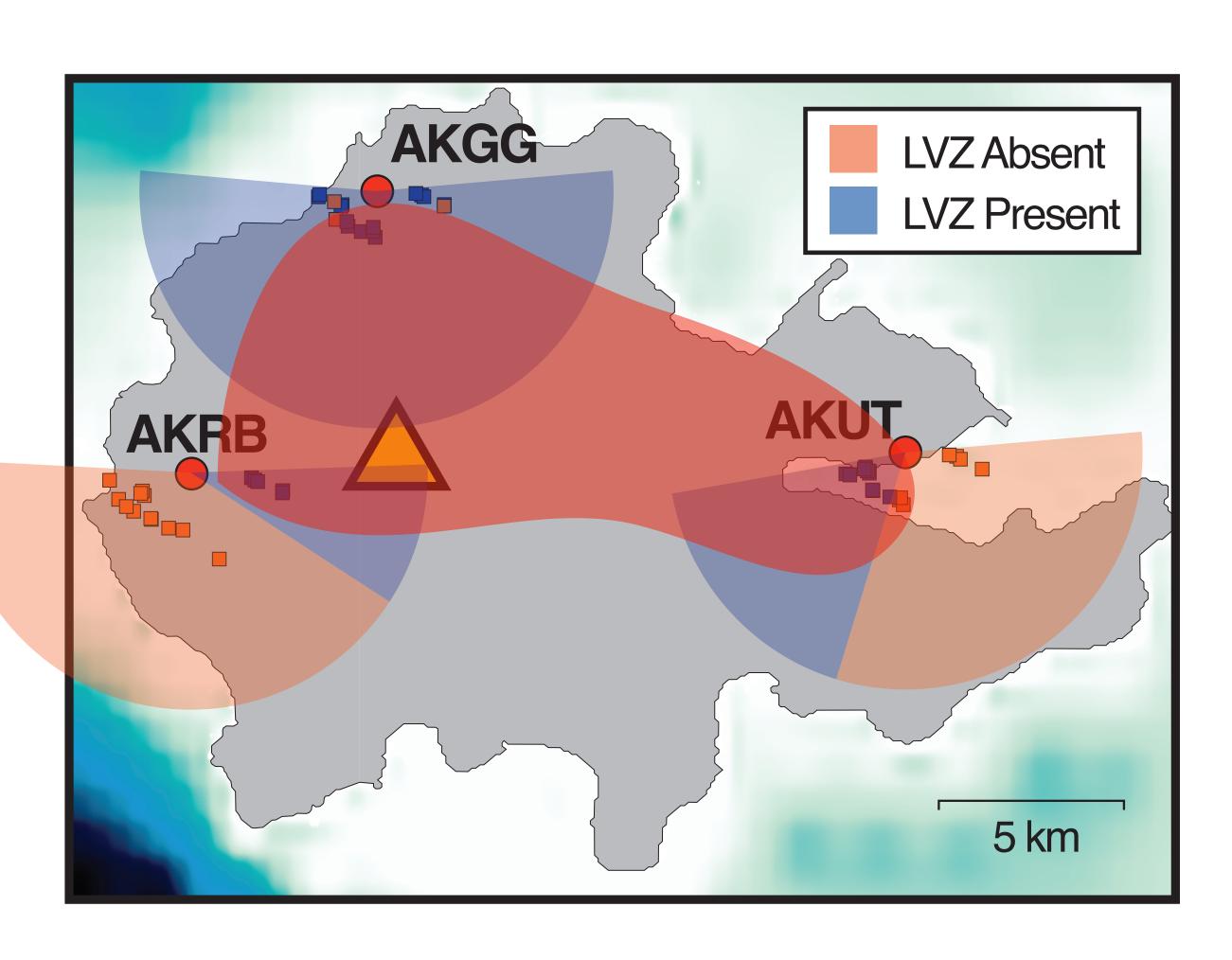
Akutan Volcano



Akutan Volcano

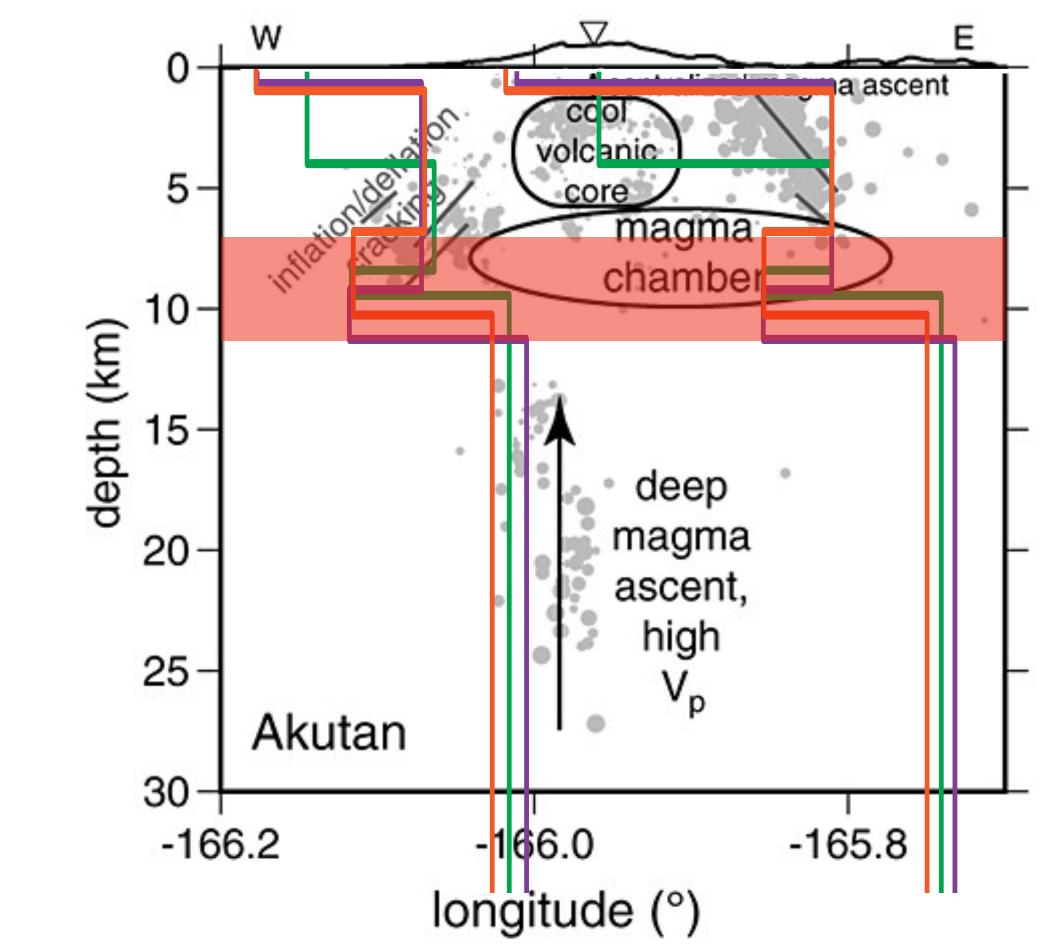
Mid-crustal magmatic region (7 - 11 km), widespread under island.

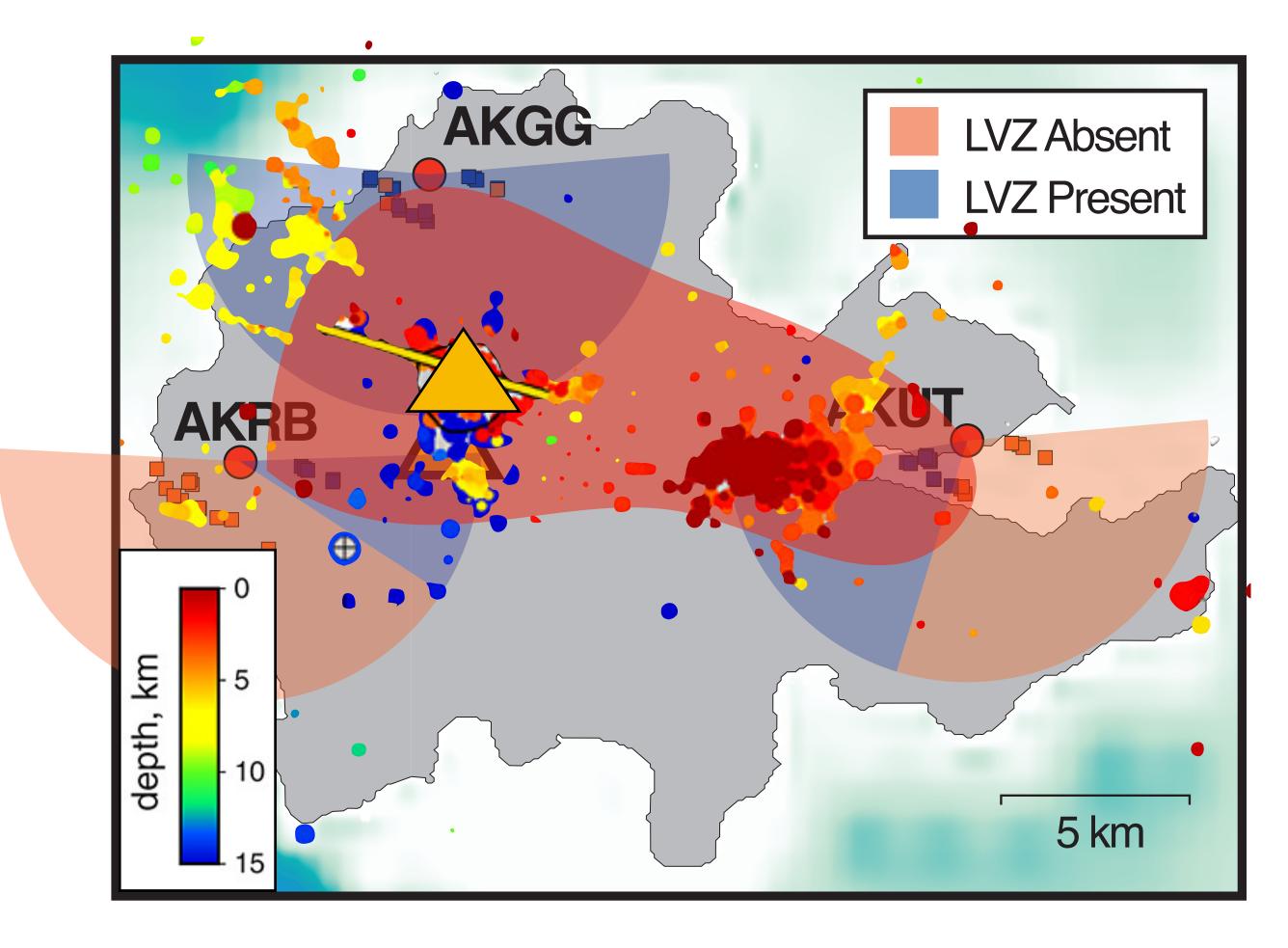




Comparison with Seismicity

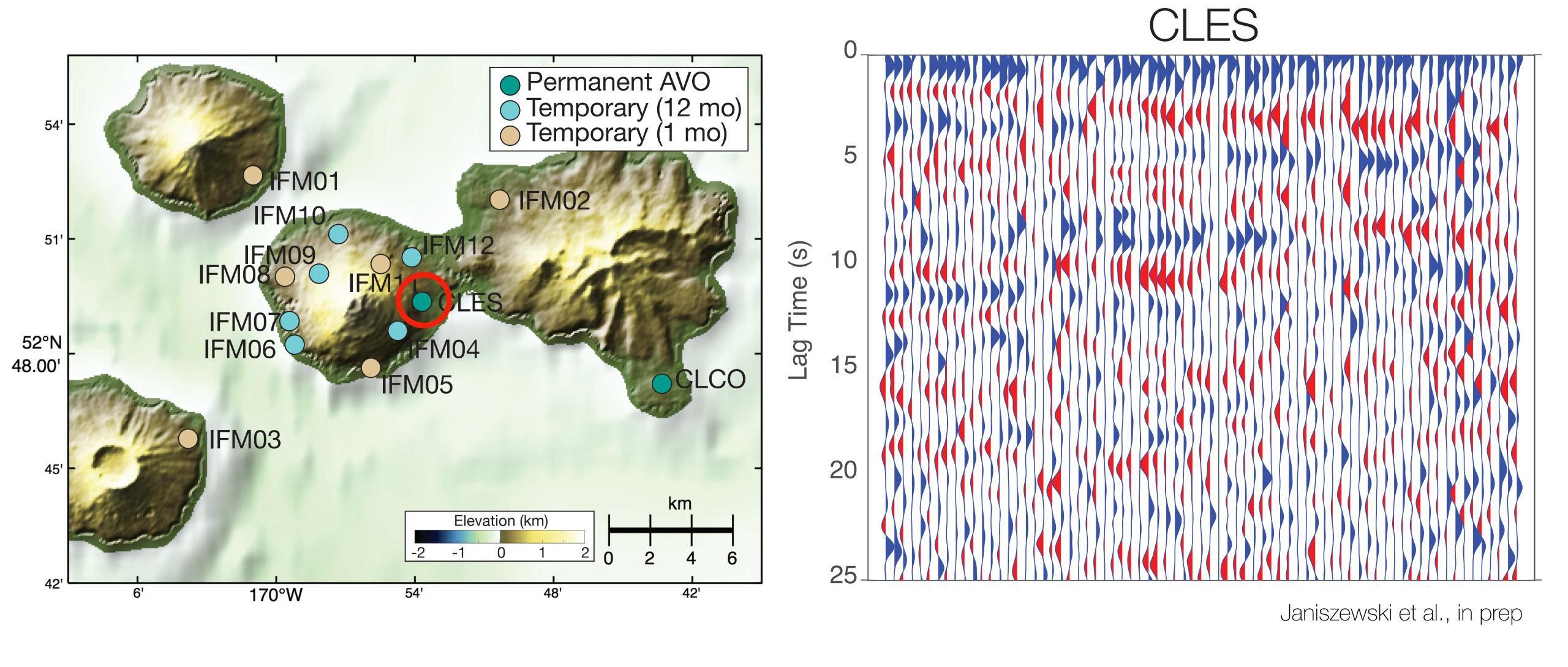
Similar depth and spatial extent as shallow seismicity related to inflation.





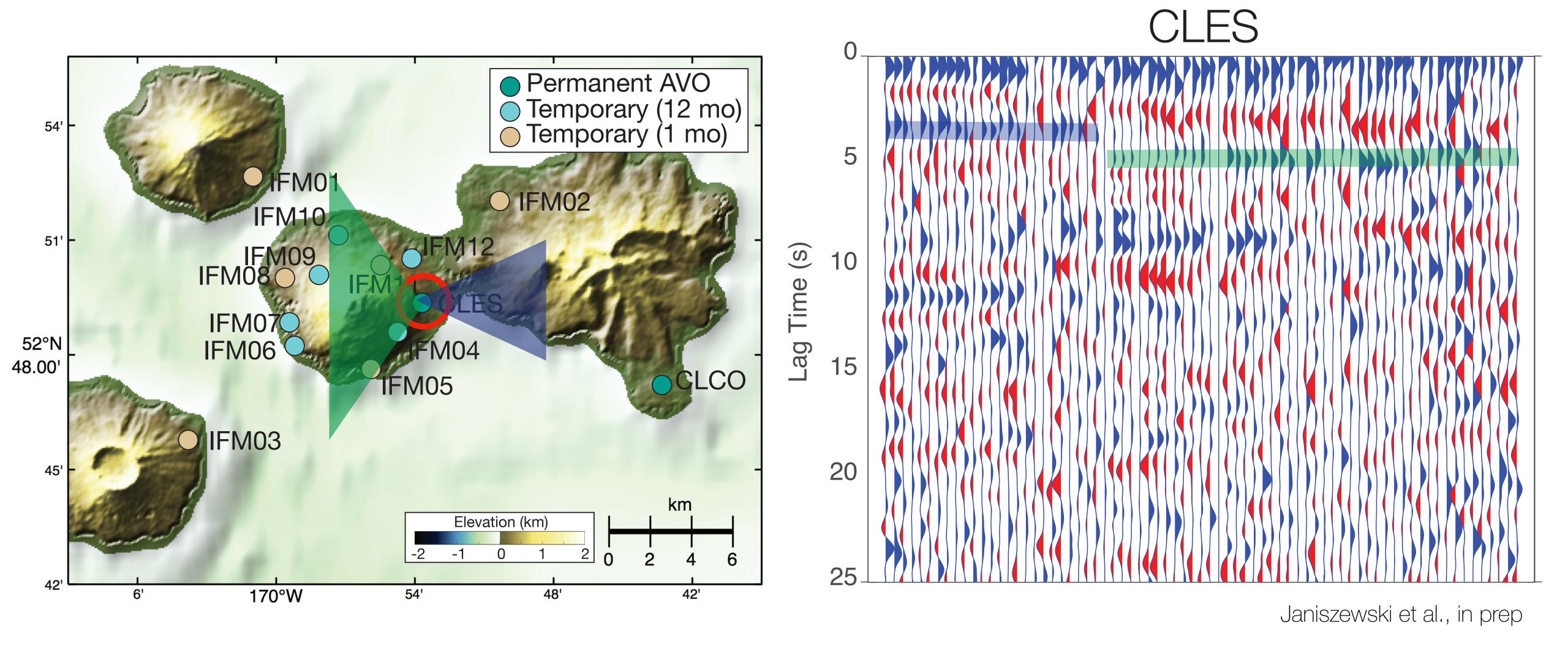
Janiszewski et al., 2013

Cleveland Volcano



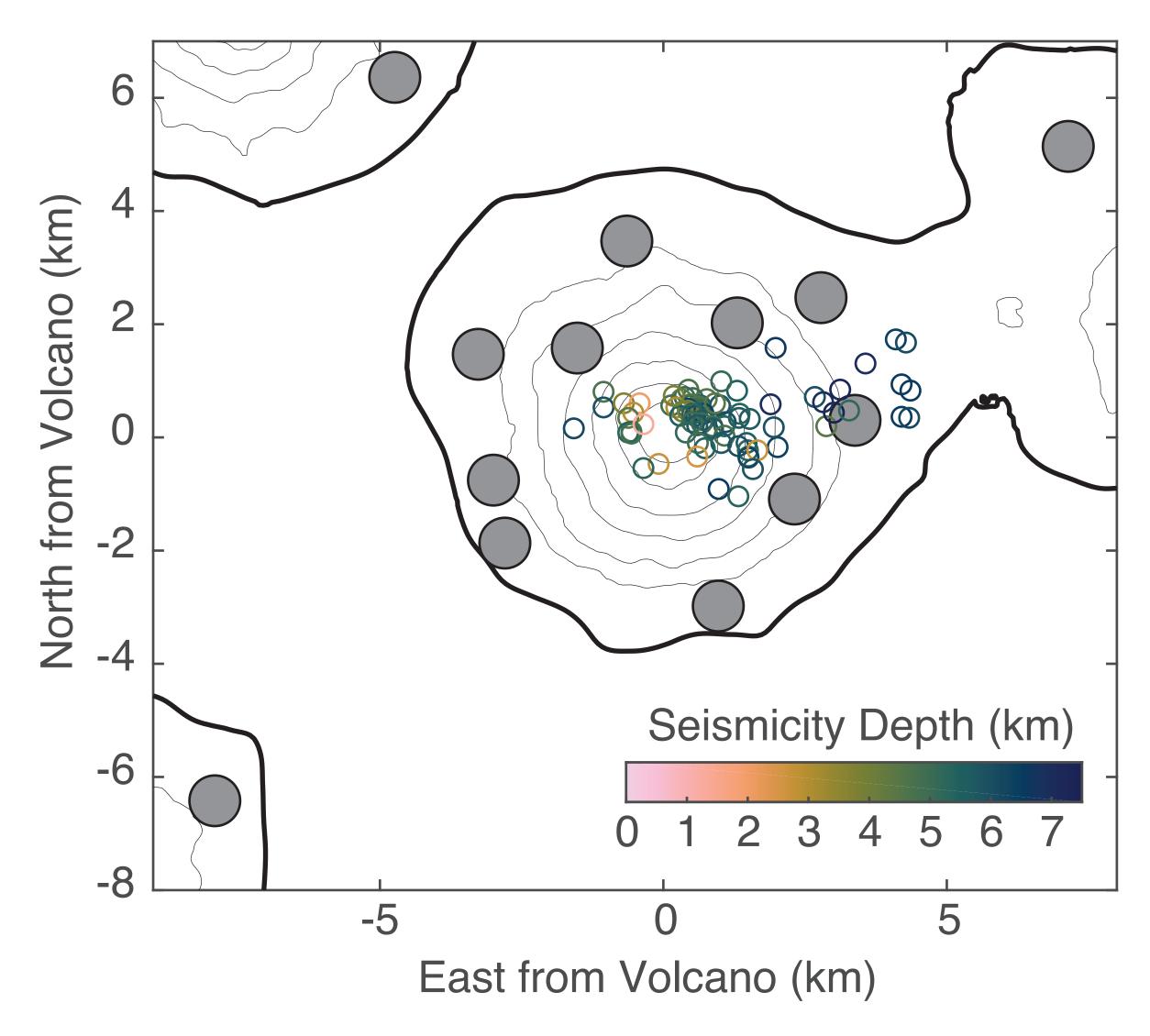
Do we see a low velocity zone underneath Cleveland, and does it have a similar relationship to seismicity as Akutan?

Cleveland Volcano



We do see local variation in the receiver function data, but no distinct additional arrival. What structures causes this?

Evidence for shallow magma storage at Cleveland



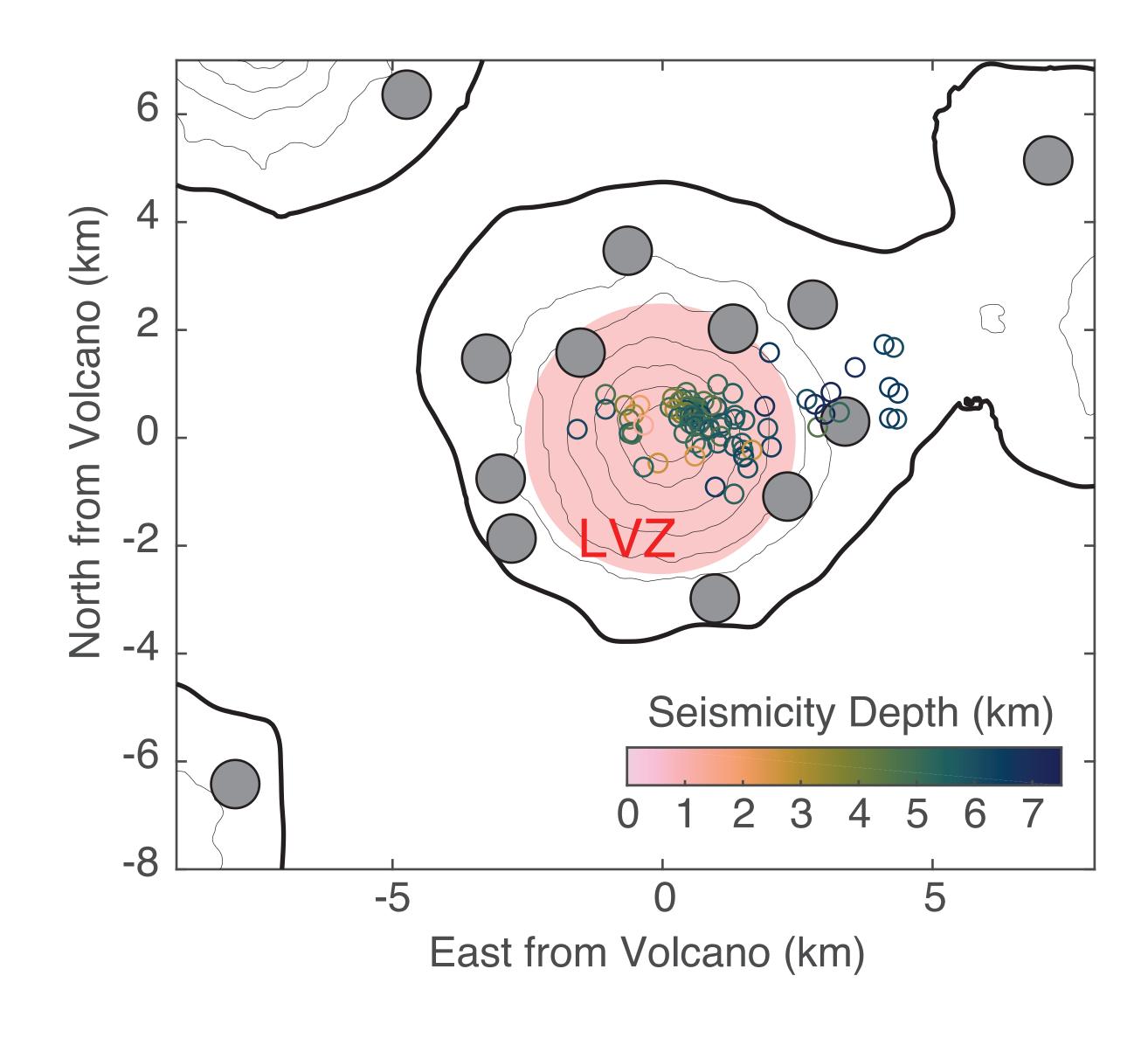
Seismicity < 7 km depth.

Magma/fluid transport at these depths.

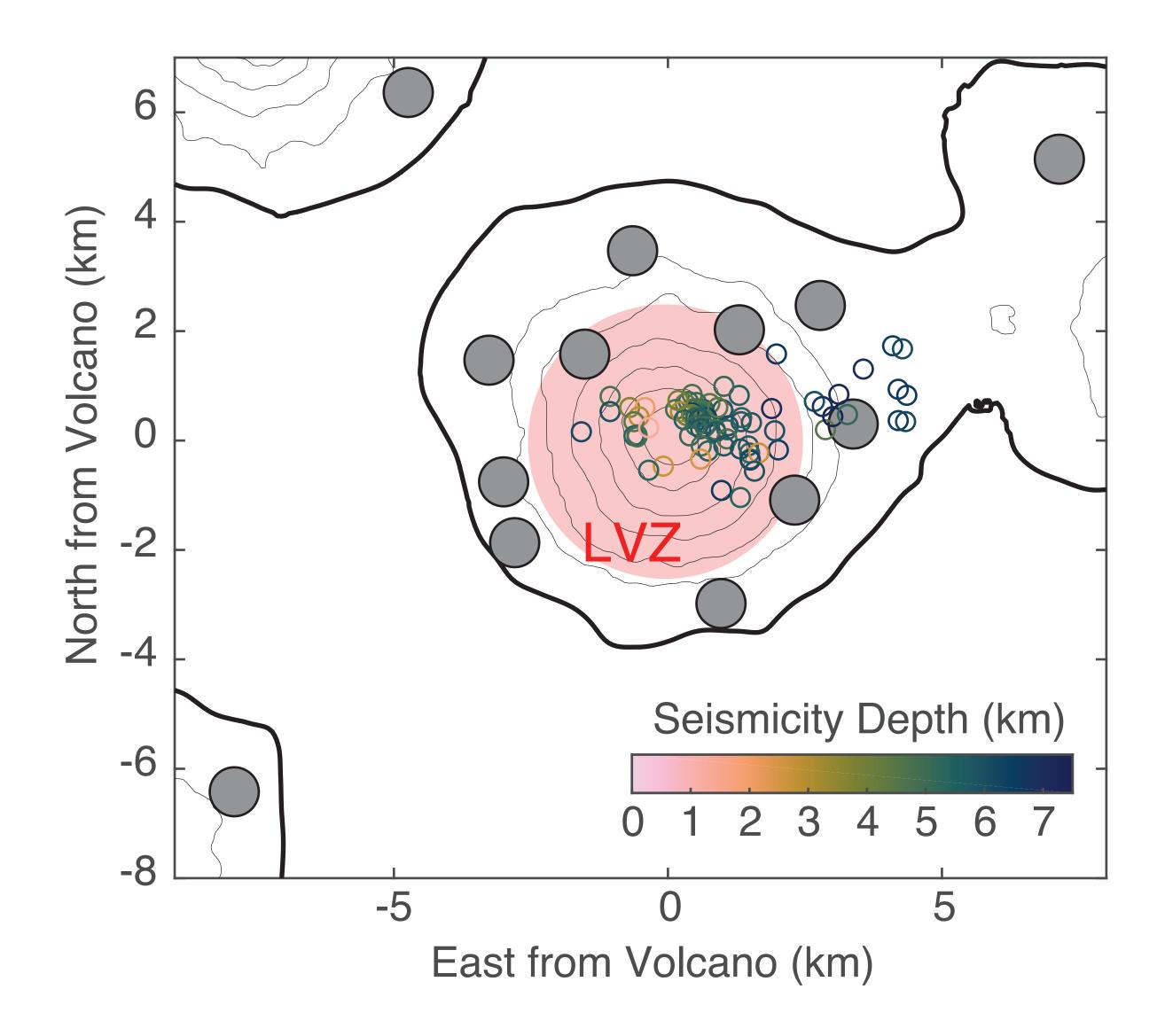
Melt inclusions suggest storage at < 5 km

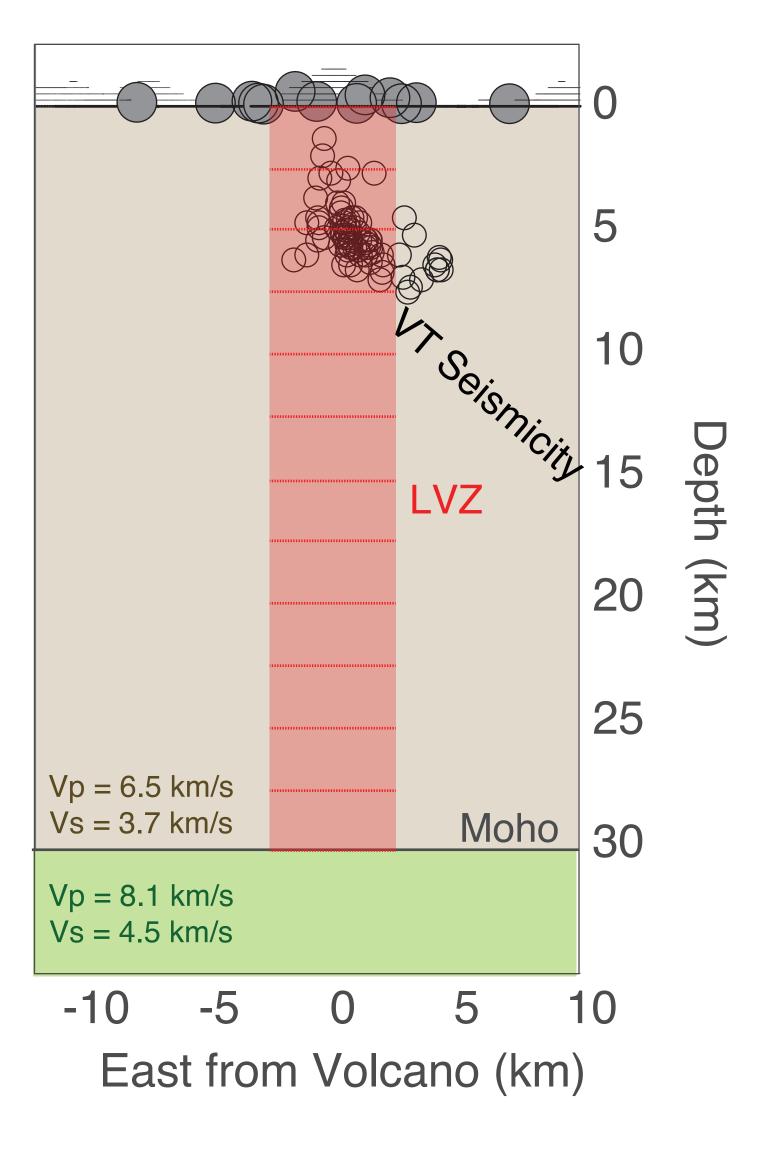
Can this shallow storage region explain the receiver functions, or are we seeing something new?

Modeling depth constraints from receiver functions

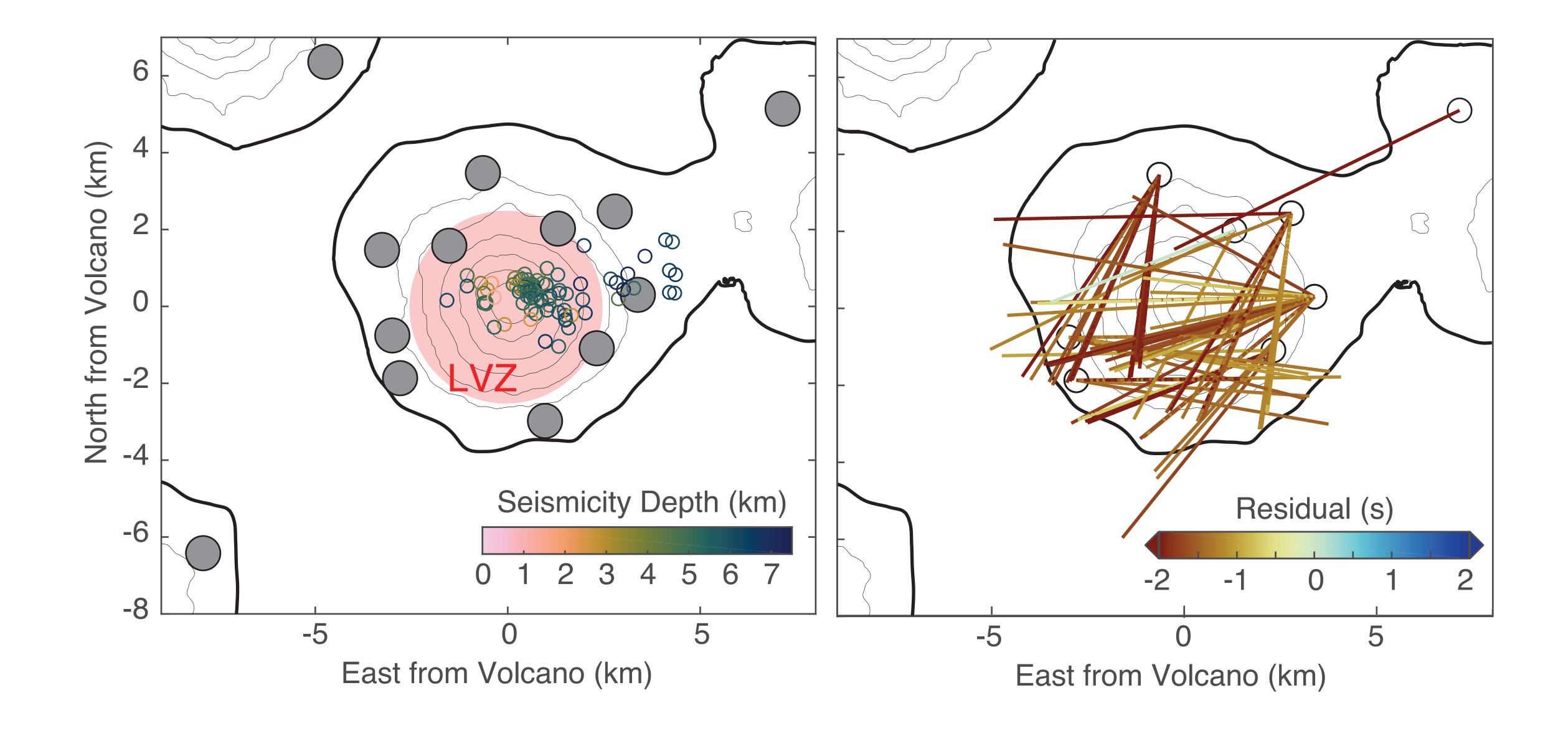


Modeling depth constraints from receiver functions





Goal: Reduce misfit in Ps lag times from the Moho

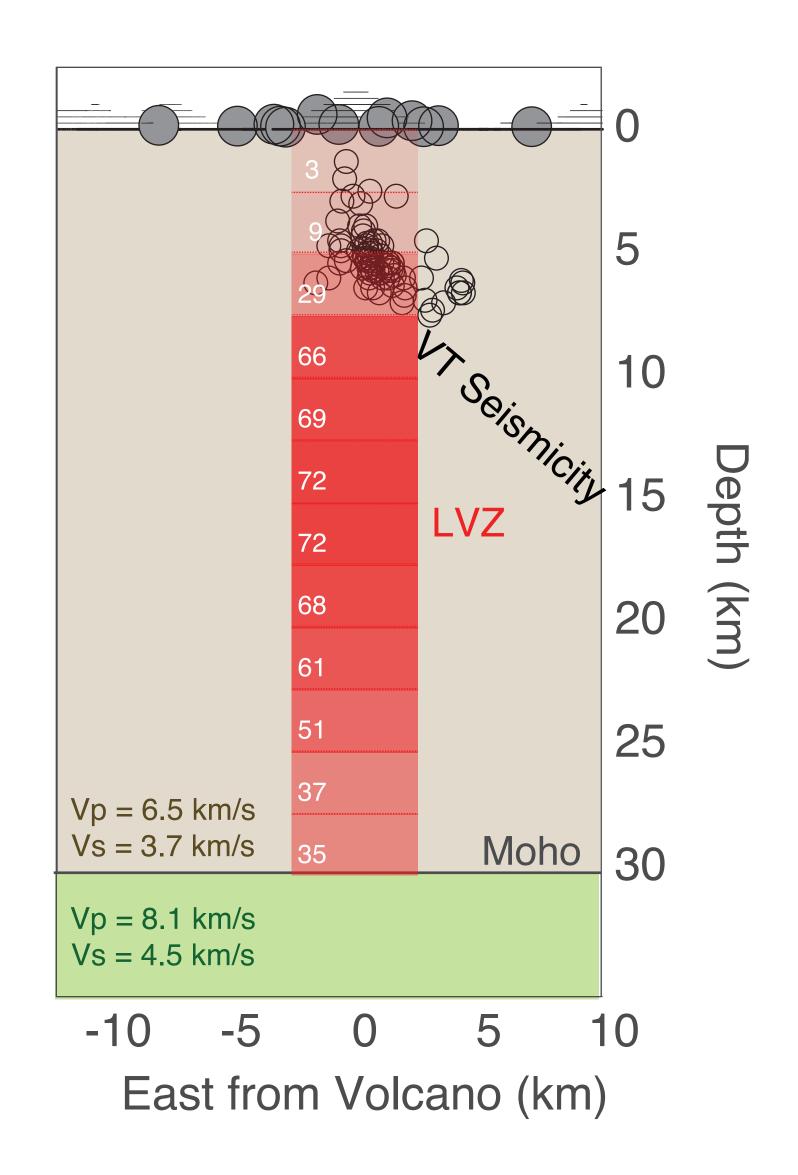


Depth Resolution

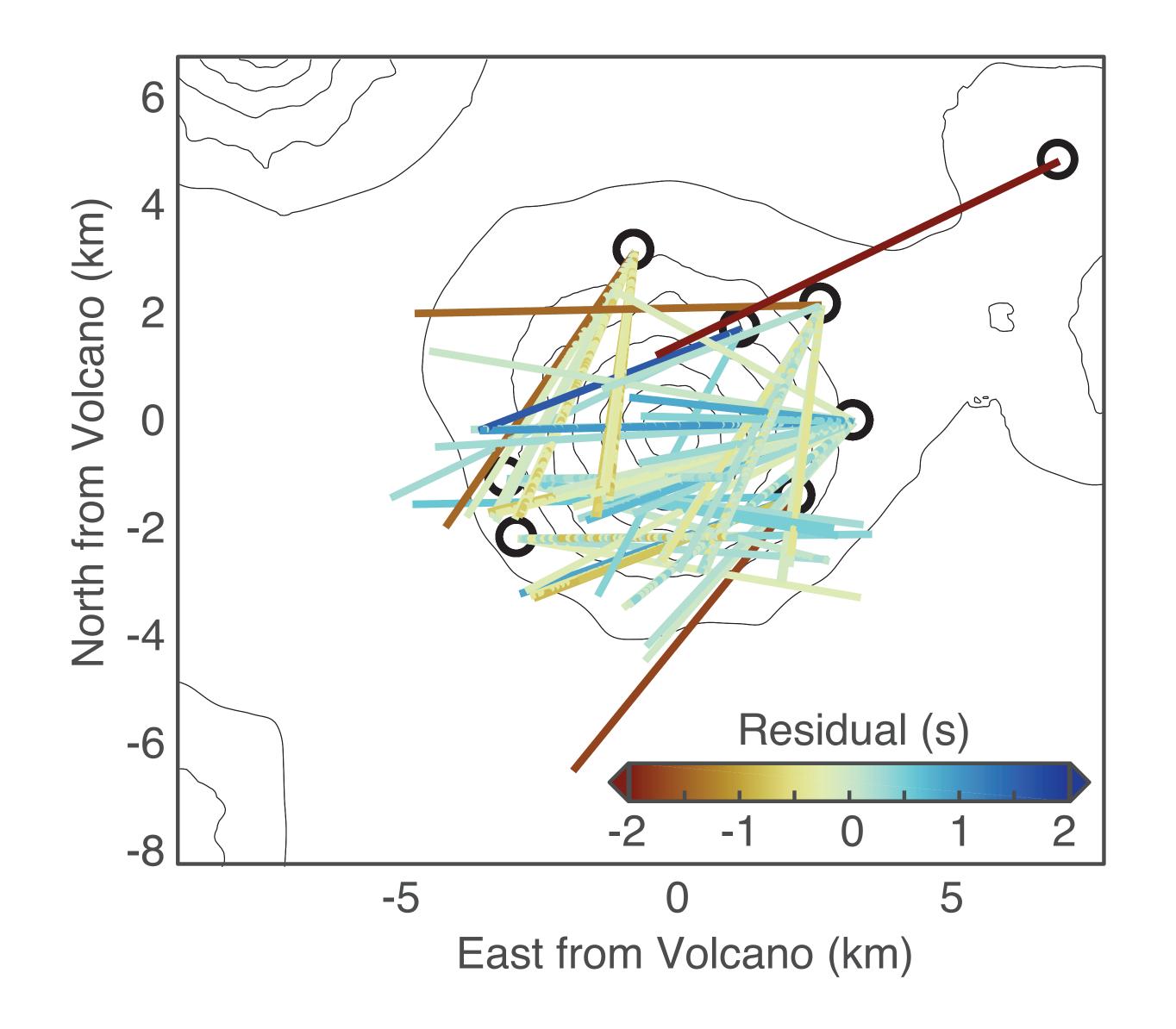
Few ray paths traverse the shallow crust.

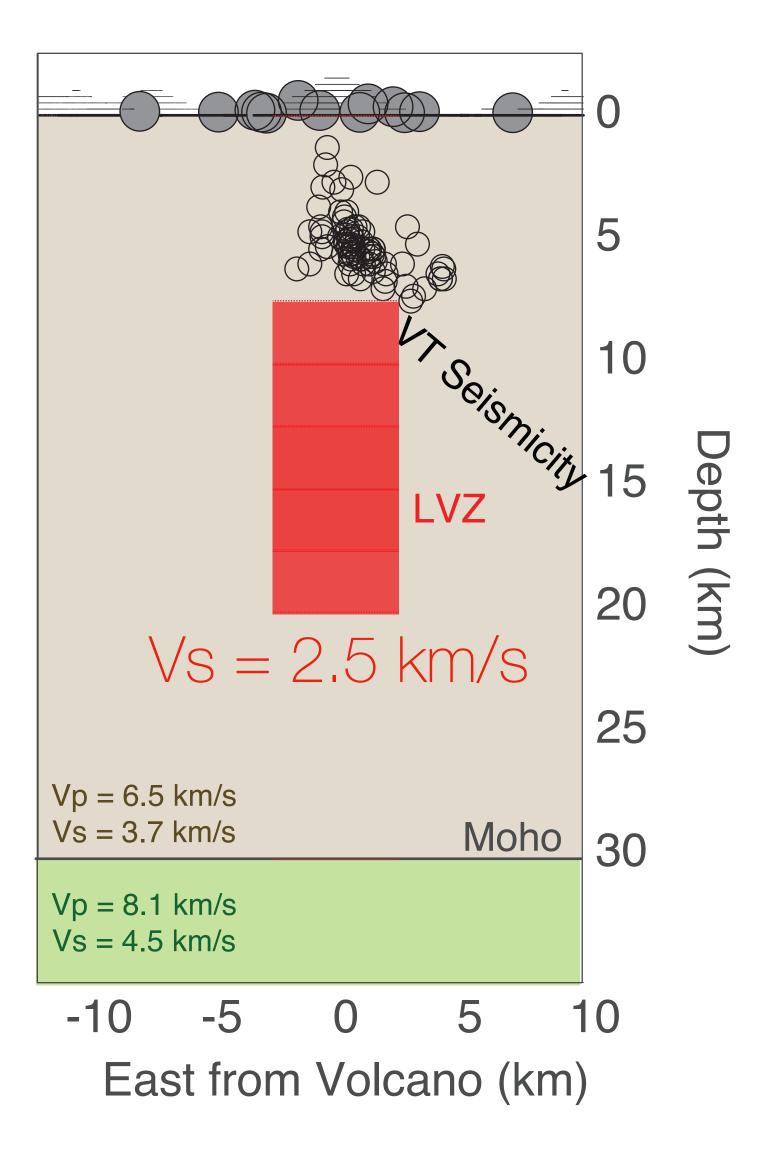
These ray paths are not sensitive to shallow crustal structure directly beneath the edifice.

Simplest model for our observations of slow ray paths crossing this central column is an LVZ in the mid crust.

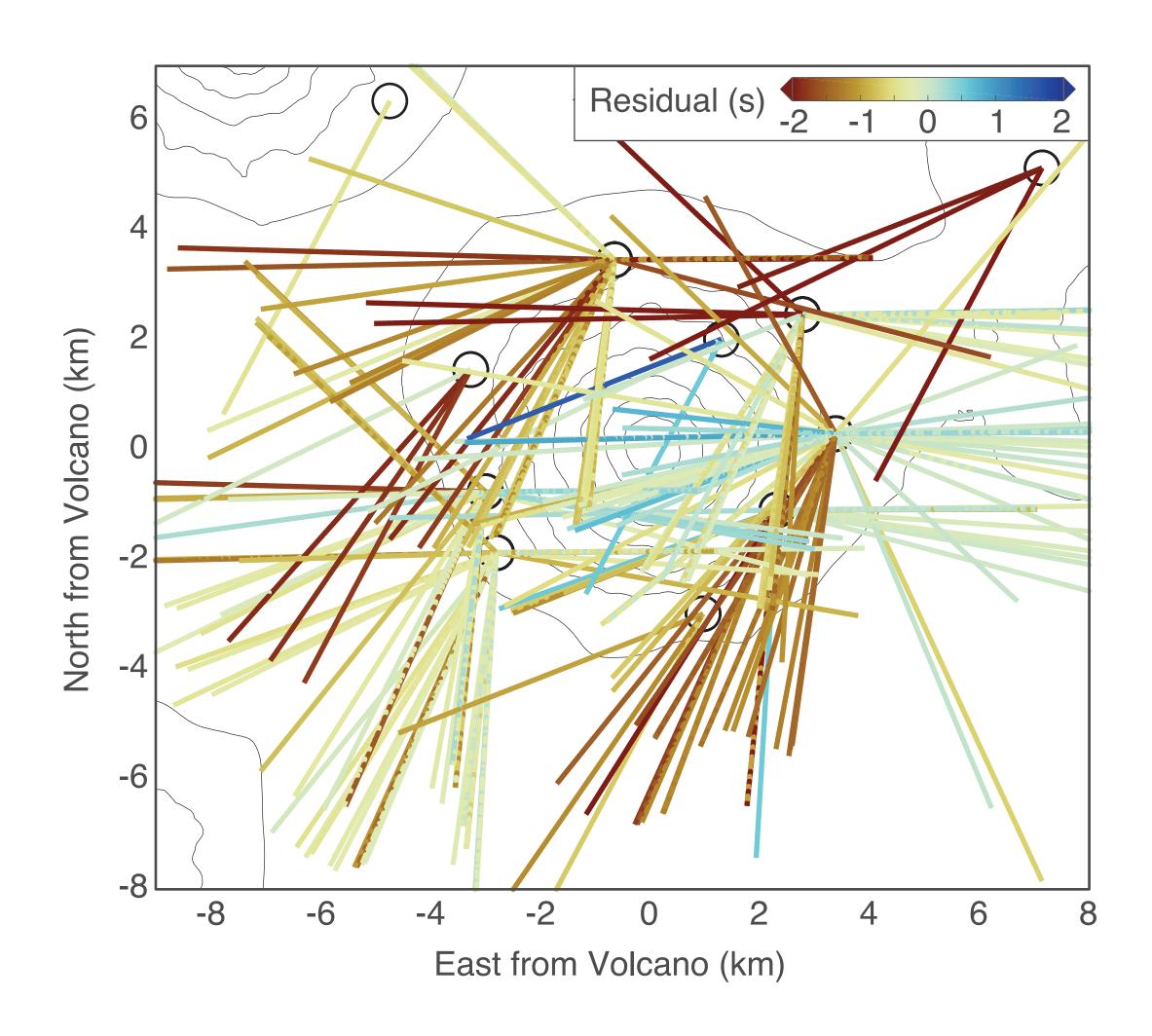


Mid-Crustal Magma Storage





Complex Magmatic Geometries

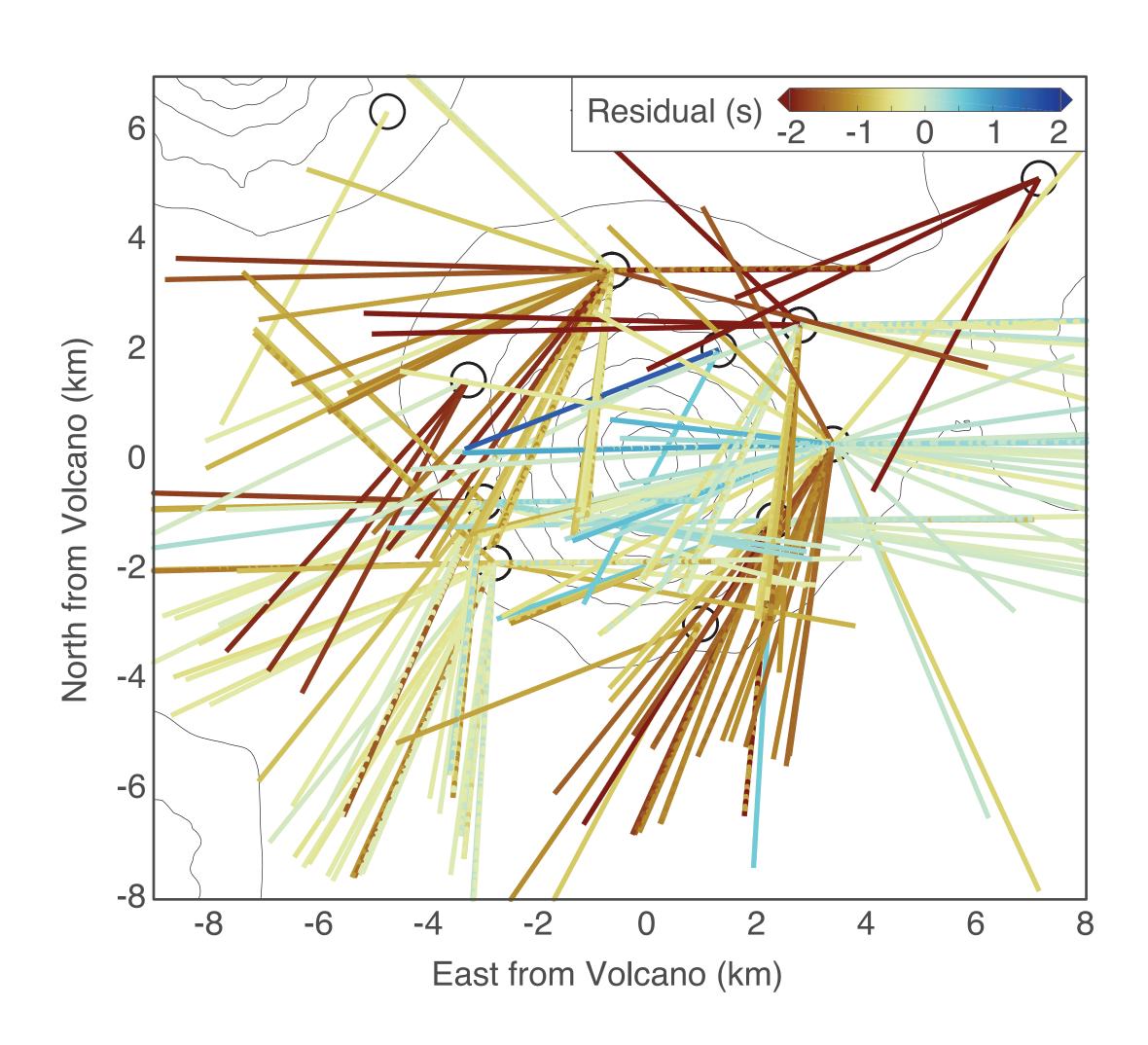


Under the main edifice, receiver functions point to a relatively thick LVZ indicating a region of melt/fluids in the mid-crust.

This doesn't explain all observations beneath Cleveland volcano.

Suggests a more complex 3-D geometry of velocity anomalies - potential for more detailed analysis.

Main Takeaways



Receiver function techniques are useful for determining basic mid- to deep-crustal magmatic architecture with only a few seismic instruments (monitoring scale).

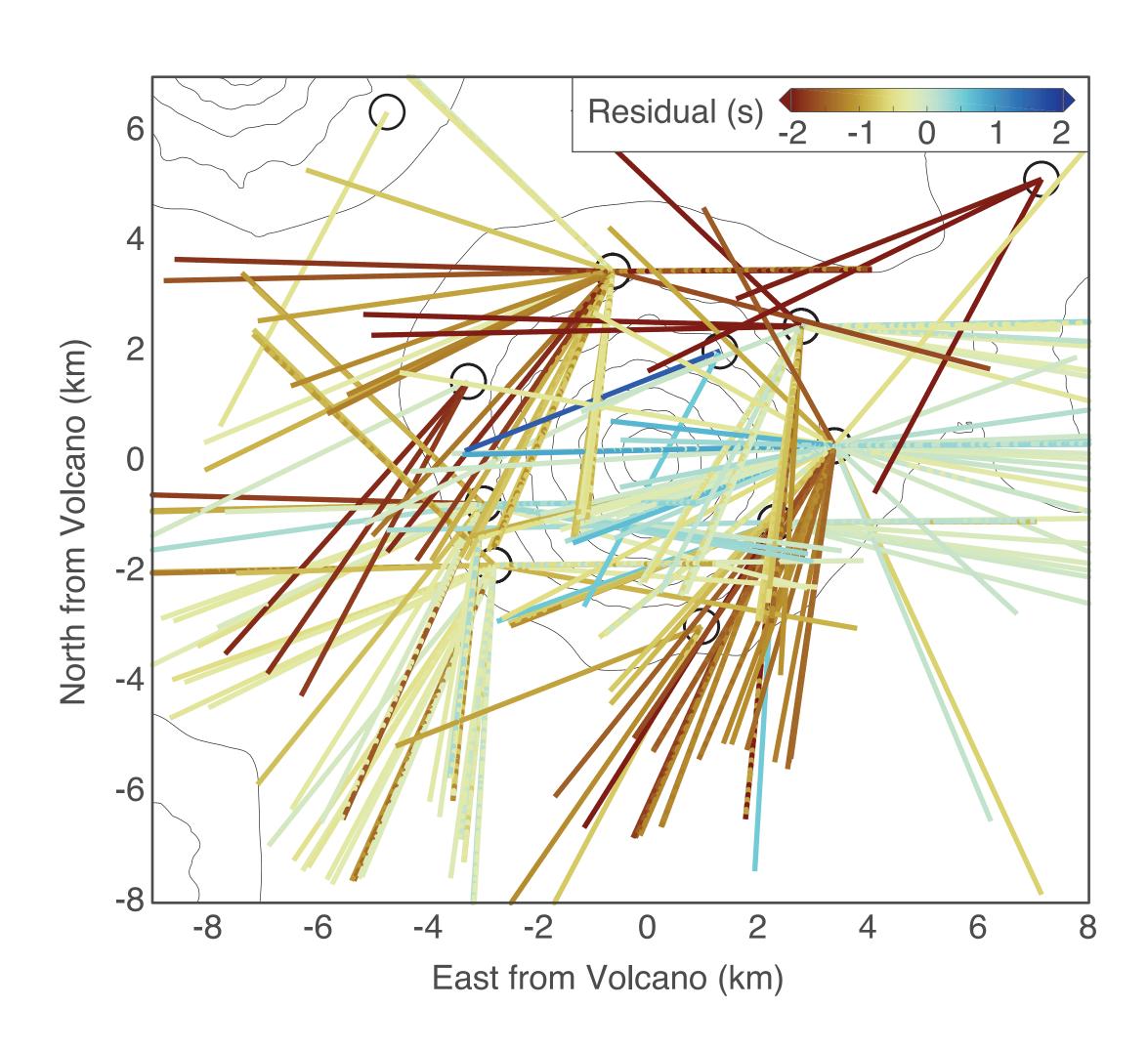
Unlike at Akutan, slow velocities wider depth range and likely extend much deeper than seismicity.

Evidence that we can discriminate between different "types" of magmatic architecture - sharp sill vs. gradual.

Useful for characterizing a difficult to constrain piece of the volcanic system with few instruments. Complements typical volcanology techniques.

Potentially useful in planning future dense deployments around volcanoes.

Main Takeaways



Receiver function techniques are useful for determining basic mid- to deep-crustal magmatic architecture with only a few seismic instruments (monitoring scale).

Unlike at Akutan, slow velocities wider depth range and likely extend much deeper than seismicity.

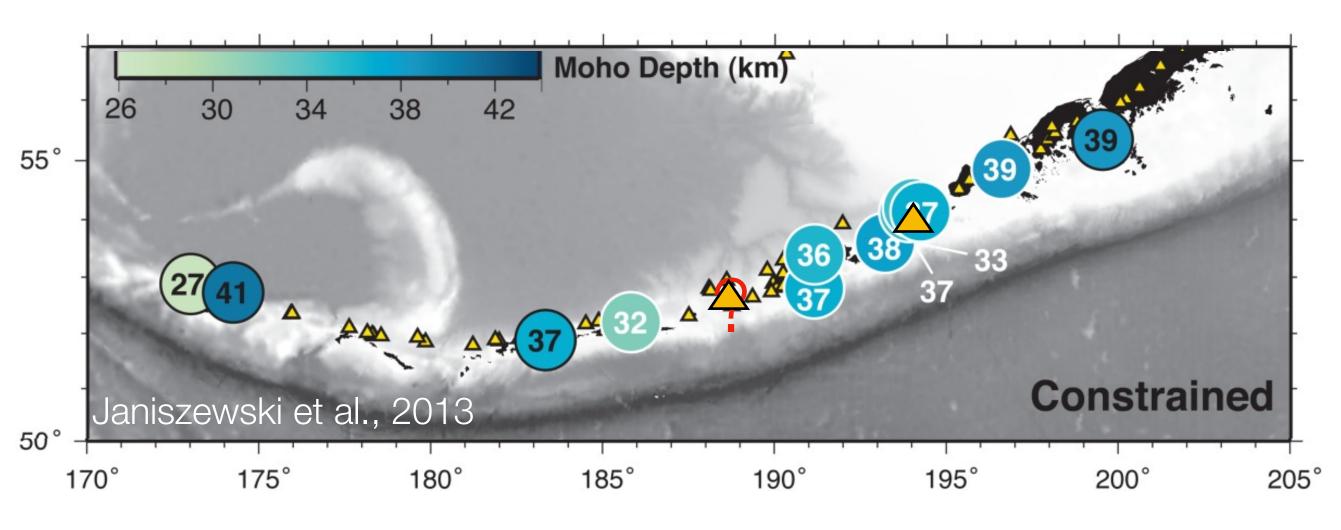
Evidence that we can discriminate between different "types" of magmatic architecture - sharp sill vs. gradual.

Useful for characterizing a difficult to constrain piece of the volcanic system with few instruments. Complements typical volcanology techniques.

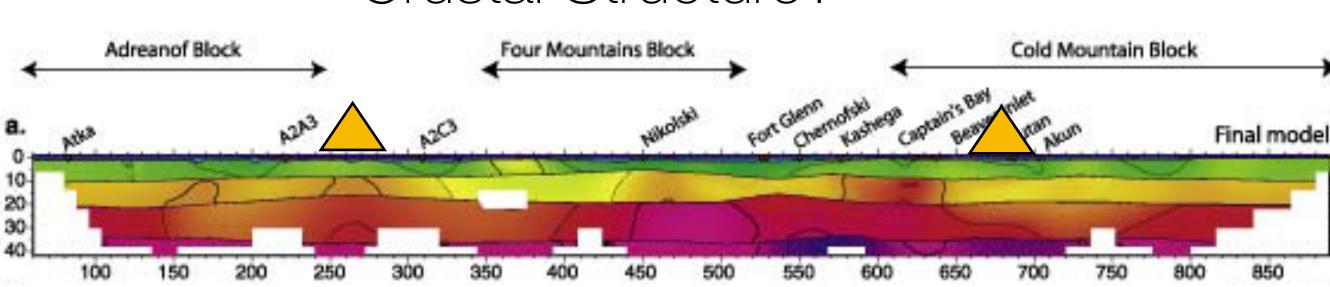
Potentially useful in planning future dense deployments around volcanoes.

Why do two volcanoes part of the same arc have such different magmatic structure?

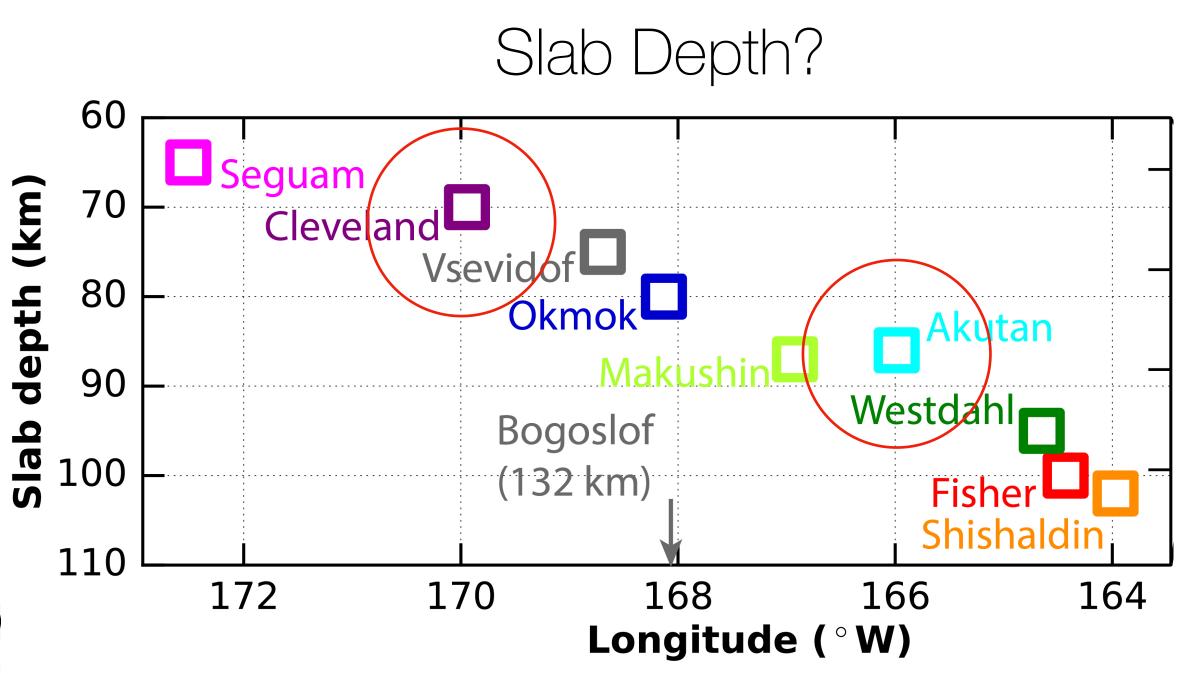
Crustal Thickness?



Crustal Structure?

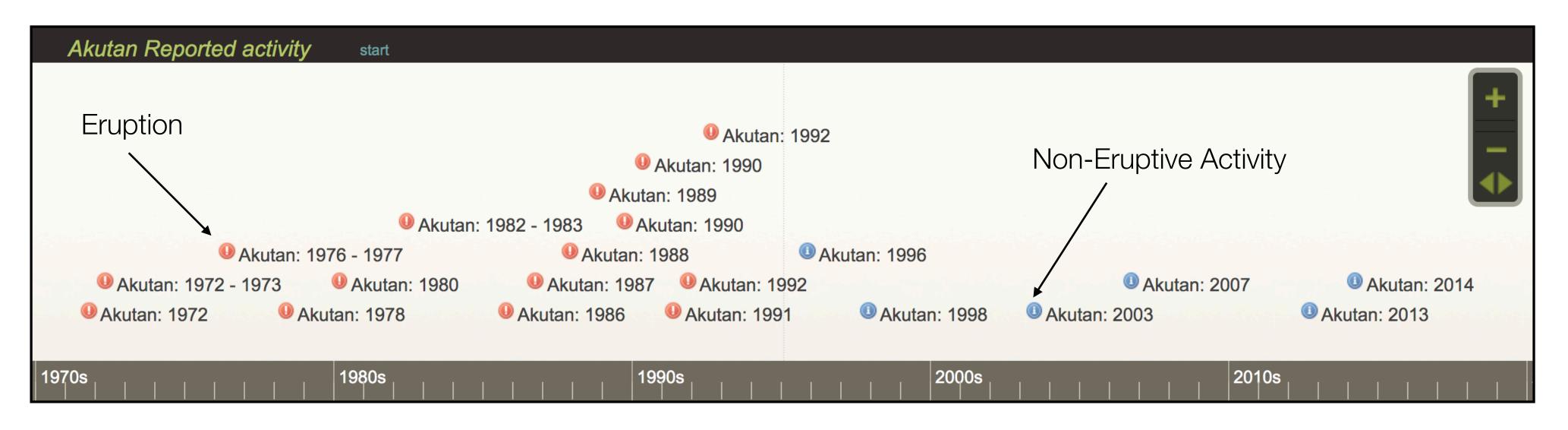


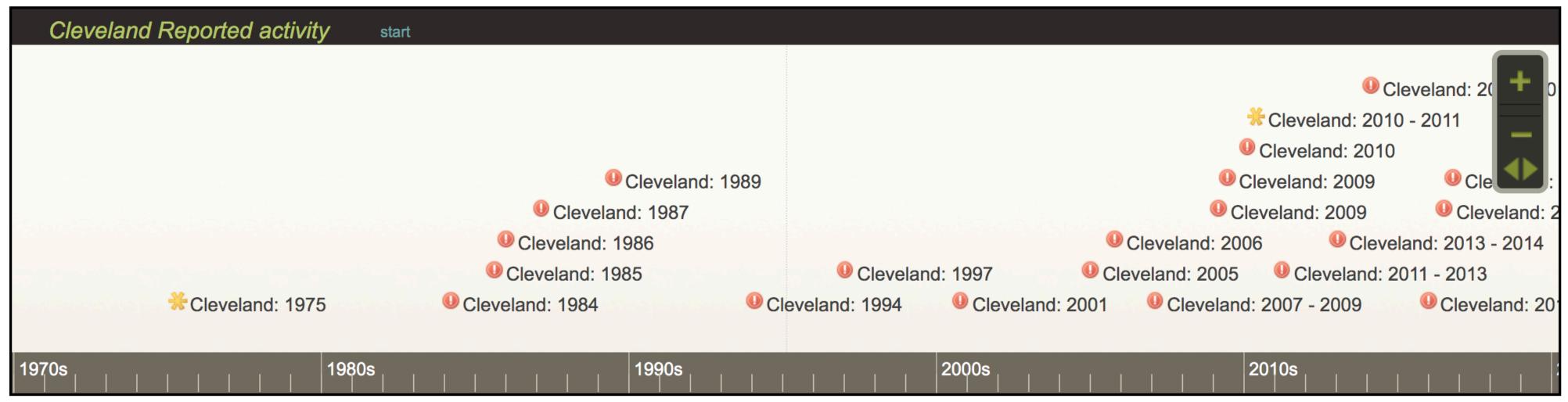
Velocity (km/s)



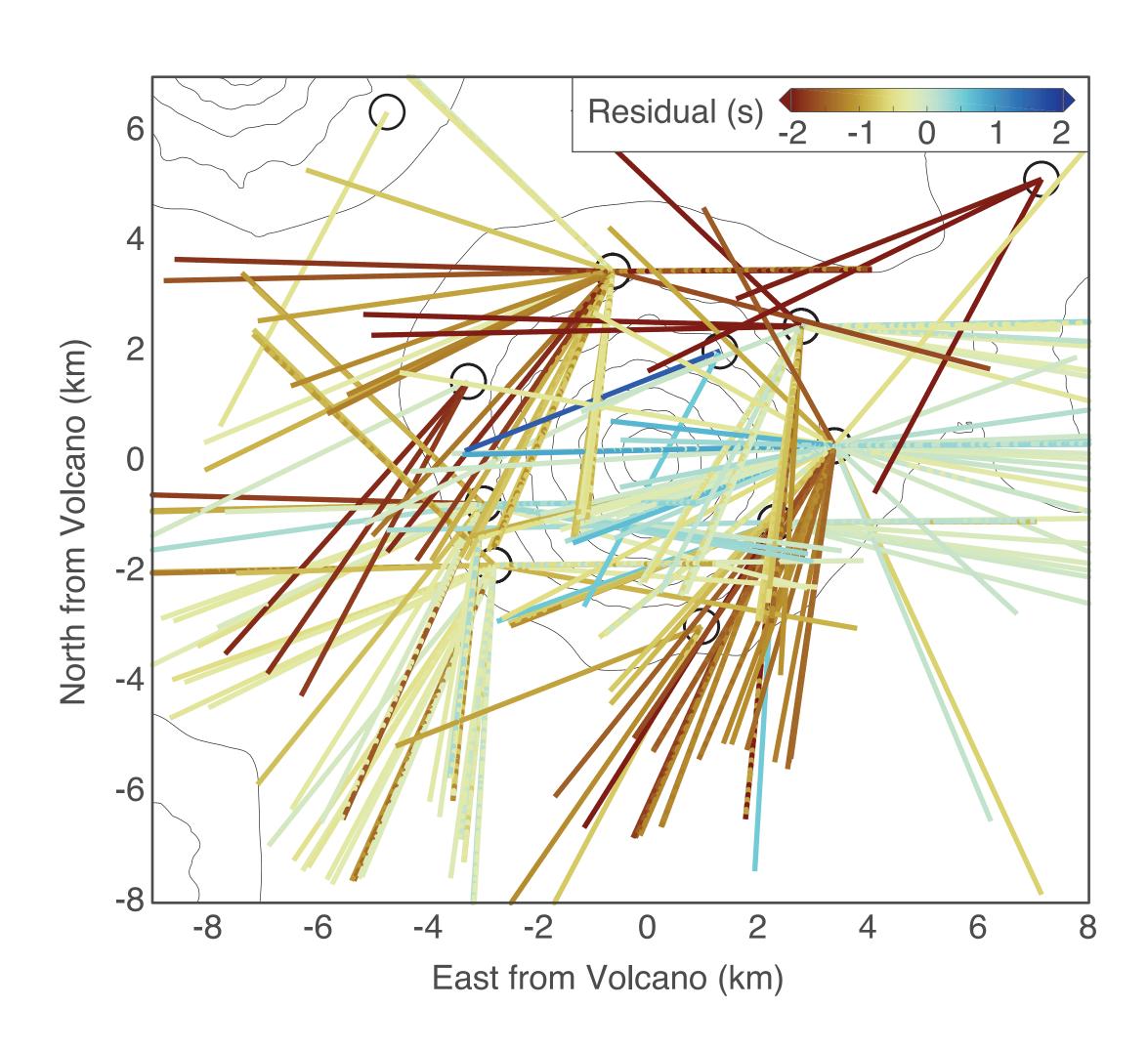
Rasmussen et al., in prep

(How) Do differences in deep crustal magmatic structure influence variations in eruptions?





Main Takeaways



Receiver function techniques are useful for determining basic mid- to deep-crustal magmatic architecture with only a few seismic instruments (monitoring scale).

Unlike at Akutan, slow velocities wider depth range and likely extend much deeper than seismicity.

Evidence that we can discriminate between different "types" of magmatic architecture - sharp sill vs. gradual.

Useful for characterizing a difficult to constrain piece of the volcanic system with few instruments. Complements typical volcanology techniques.

Potentially useful in planning future dense deployments around volcanoes.